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53			arming trend in daily temperature extremes was projected for much of Asia (medium			
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most of Asian region including the Tibetan Plateau. Annual mean precipitation trends are characterized by strong
 variability, with both increasing and decreasing trends observed throughout Asian regions [Table 24-2].

3

4 Water scarcity is expected to a major challenge for most of the region due to increase of water demand and 5 soaring water supply and lack of good management (medium confidence). Freshwater availability in Central, 6 South, East and South-East Asia, particularly in large river basins, is projected to decrease due to climate change 7 which, along with population growth and increasing demand arising from higher standards of living, could adversely 8 affect more than 1 billion people by the 2050s. Shrinking of glaciers in Central Asia and the Himalayas is projected 9 to affect water resources positively in the near future but negatively in the long term perspective. Better water 10 management strategies are needed to ease water scarcity. Water saving technologies and changing of crops into 11 drought tolerant crops are found to be successful adaptation options in the region. 12

13 The impacts of climate change on food production and food security in Asia will vary by region with many

regions experiencing a decline in productivity (medium confidence). This is evident in the case of rice

production would be generally negative in many regions. Most models using a range of GCMs and SRES scenarios show that higher temperatures will lead to lower rice yields as a result of shorter growing period. There are a number

of regions that are already near the critical temperature threshold. However, with CO2 fertilization, rice yield could

increase with climate change. This is also true for other crops. In Central Asia, some areas could be gainers (cereal

19 production in northern and eastern Kazakhstan could benefit from the longer growing season, warmer winters and

slight increase in winter precipitation), while others can could be losers (western Turkmenistan and Uzbekistan,

where frequent droughts will could negatively affect cotton production, increase water demands for irrigation, and

exacerbate desertification). In the Indo-Gangetic Plains (IGPs) of South Asia, there could be up 50% decrease in the

most favorable and high yielding wheat area due to heat stress at 2x CO2. There are many potential adaptation

strategies such as crop breeding but research on their effectiveness is limited [24.4.4].

25

26 Terrestrial and marine ecosystems are increasingly under pressure from both climatic and non-climatic

drivers; the projected changes in climate will impact natural and semi-natural vegetation, permafrost
 degradation spread and widespread damage to coral reefs in Asia during 21st Century (high confidence)

29 [24.2.2, 24.4.2, 24.4.3]. The largest changes and the highest rates of change are expected in cold northern and high-

altitude areas, where boreal and subalpine trees will *likely* invade treeless arctic and alpine vegetation, and evergreen

31 conifers will *likely* invade deciduous larch forest. Large changes may also occur in arid and semi-arid areas, but

32 uncertainties in precipitation projections make these more difficult to predict. Rates of vegetation change in the more 33 densely populated parts of Asia will be constrained by the impact of vegetation fragmentation on seed dispersal. The

34 impacts of projected climate changes on the vegetation of the lowland tropics are currently poorly understood.

35 Permafrost degradation during the 21st century will spread from the southern and low-altitude margins, advancing

36 northwards and upwards (24.4.2.3.2.). Many models agree on the direction of change, but rates of change vary

37 greatly between different model projections. In the Asian Arctic, there is *high agreement* and *medium evidence* that

rising sea-levels will interact with projected changes in permafrost and the length of the ice-free season to cause

increased rates of coastal erosion (Section 24.4.3.3.). Widespread damage to coral reefs correlated with episodes of

40 high sea-surface temperature has been reported in recent decades and there is *high confidence* that such damage will

41 increase during the 21st century as a result of both warming and ocean acidification (Sections 24.4.3.2. and

- 42 24.4.3.3.). However the capacity of coral reefs to adjust by changes in species composition, or by the acclimation or43 adaptation of coral species, is not well understood.
- 44

45 It is very *likely* that mean sea level rise will contribute to upward trends in extreme coastal high water levels

46 **in the delta.** Even most of the major deltas in Asia are now sinking at rates many times faster than the global sea-

47 level is rising. Widespread impacts can be attributed with high confidence to climate change, however, for coral

reefs, where the temporal and spatial patterns of large-scale bleaching events generally correlate well with higher than normal sea surface temperatures. Coastal freshwater swamps and marshes will be vulnerable to saltwater

49 than normal sea surface temperatures. Coastal fr50 intrusion with rising sea-levels [24.4.3].

51

52 Extreme events will have greater impacts on sectors with closer links to climate, such as water, agriculture

53 and food security, forestry, health, and tourism. (high confidence) More frequent and intense heat-waves in Asia

54 will increase mortality and morbidity in vulnerable groups; in particular in urban environments (urban heat island

effect), in combination with air pollution (from wildfires, traffic, etc), or among outdoor workers in both urban and
 rural environments [24.4.6].

- Multiple stresses caused by rapid urbanization, industrialization and economic development *are likely* to be compounded by climate change. Climate change is also expected to adversely affect sustainable development capabilities of most Asian developing countries by aggravating pressure on natural resources and the environment. Development of sustainable cities in Asia with less fossil fuel driven vehicles (mitigation) and with more trees and greenery (carbon storage as well as adaptation to urban heat island effect) would have a number of co-benefits including public health [24.4, 24.5, 24.6, 24.7].
- 10 11

13

12 24.1. Introduction

14 Asia is defined here as the land and territories of 51 countries/regions (Figure 24-1). It can be broadly divided into 15 six sub-regions based on geographical position and coastal peripheries (Table 24-1). These are (in alphabetical 16 order) Central Asia (5 countries), East Asia (7 countries/regions), North Asia (2 countries), South Asia (8 countries), 17 Southeast Asia (12 countries) and West Asia (17 countries). Asia has a diversity of social, cultural and economic 18 characteristics. The population of Asia in 2009 was reported to be about 4,121 million, which is 60.3% of the world 19 population (UN, 2009). The population density is about 130 per square kilometer (PRB, 2010). The highest life 20 expectancy at birth is 82.7 (Japan) and the lowest is 43.8 (Afghanistan). In 2009, the GDP per capita ranged from 21 US\$492 (Timor-Leste) to US\$39,738 (Japan) (World Bank, 2011). About 40% of the population in the developing 22 countries of Asia lives below the poverty line, where their income is below US\$ 1.25 per day by 2005 prices (World

- 23 Bank, 2008).
- 24

25 [INSERT FIGURE 24-1 HERE

Figure 24-1: The land and territories of 51 countries/regions.]

28 [INSERT TABLE 24-1 HERE

Table 24-1: The 51 countries/regions in the six sub-regions of Asia.]

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24.2. Major Conclusions from Previous Assessments

34 24.2.1. Climate Change Impacts

36 *Climate change and variability.* The observed increases in surface temperature presented in The Fourth Assessment 37 Report (AR4) range between less than 1°C to 3°C per century, with most pronounced increases noted in North Asia 38 [AR4, Chapter 10, 10.2.2]. In addition, the interseasonal, interannual and spatial variability in rainfall trends has 39 been observed during the past few decades all across Asia [AR4, Chapter 10, 10.2.2]. Future projections show that 40 warming is least rapid in South-East Asia, stronger over South Asia and East Asia and greatest in the continental 41 interior, with most pronounced warming at high latitudes in North Asia [AR4, Chapter 10, 10.3.1]. Annual 42 precipitation projections indicate an increase in most of Asia during this century [AR4, Chapter 10, 10.3.1]. Also an 43 increase in extreme weather event occurrences (including heat waves and intense precipitation events) is projected 44 for South Asia, East Asia, and South-East Asia, along with an increase of intensity in tropical cyclones in the same 45 regions, due to a rise in sea-surface temperature [AR4, Chapter 10, 10.3.1]. A warming trend in daily temperature 46 extremes was projected for much of Asia (medium confidence) [SREX, Chapter 3, 3.3.1] No systematic spatially 47 coherent trends in heavy precipitation have been found in most of Asia, except for a weak increase of the frequency 48 of extreme precipitation that was observed in northern Mongolia (low to medium confidence) [SREX, Chapter 3, 49 3.3.2]. However, both positive and negative statistically significant trends have been found at sub-regional scales 50 throughout Asia (low to medium confidence) [SREX, Chapter 3, 3.3.2]. Future projections show that heavy precipitation is projected to increase in West and South Asia, as well as the Asian monsoon region, notably in 51 52 Bangladesh and in the Yangtze river basin [SREX, Chapter 3, 3.3.2]. A decreasing trend was observed in rainfall in 53 the South Asian and East Asian monsoons, due to a rise in sea-surface temperature [SREX, Chapter 3, 3.4.1]. 54 Increases in precipitation were projected for the Asian monsoon, while projection results for the south Asian

1 monsoon precipitation point out to both increases and decreases in precipitation (low confidence) [SREX, Chapter

2 3., 3.4.1]. The coastal areas of Asia have reported that the sea level rise is accelerated relative to the long-term 3 average and greater than the global average [AR4, Chapter 10, 10.3.1]. Greatest vulnerability in terms of inundation

4 of land area to a 1m sea level rise is located in East Asia and the Pacific, followed by South Asia (high confidence)

5 [SREX, Chapter 3, 3.5.5].

6

7 *Climate change impacts.* Changes in drought patters have been reported for the monsoon regions of Asia with 8 variations at the decadal time scale (low confidence) [SREX, Chapter 3, 3.5.1]. Studies on East Asia show 9 increasing dryness in the second half of the 20th century (medium confidence) [SREX, Chapter 3, 3.5.1]. Other research data projects a higher likelihood of hydrological drought by the end of the century, with a substantial 10 11 increase in the number of drought days in southern Asia from Indochina to southern China, while increases in 12 drought are projected for inland China and central Eurasia [SREX, Chapter 3, 3.5.1]. Flood observation results show 13 that there is an upward trend in the annual flood maxima of the lower Yangtze, increasing likelihood for extreme 14 floods in the Mekong river, and both upward and downward trends in four selected river basins of the northwestern 15 Himalaya (low confidence) [SREX, Chapter 3, 3.5.2]. Projections point out to an increase in the risk of floods in 16 most humid Asian monsoon regions (low confidence) [SREX, Chapter 3, 3.5.2].

17 18

19 24.2.2. Vulnerabilities and Adaptive Strategies 20

21 Vulnerable sectors. Crop yields in the past few decades has declined in many parts of Asia due to increasing water 22 stress arising partly from increasing temperature, increasing frequency of El Niño and reduction in the number of 23 rainy days (medium confidence) [AR4, Chapter 10, 10.2.4.1; Chapter 10, Executive Summary]. Studies suggest that 24 in the future as well substantial decreases are probable not only in cereal production potential (medium confidence) 25 [Chapter 10, Executive Summary], but also in livestock, fishery, and aquaculture net primary productivity [AR4, 26 Chapter 10, 10.4.1.1, 10.4.1.3]. Most projections suggest that increasing urbanization and population in Asia could 27 result in increased food demand and reduced food supply due to limited availability of cropland area and yield 28 declines [AR4, Chapter 10, 10.4.1.4]. Food insecurity and loss of livelihood would be further exacerbated by the 29 loss of cultivated land and nursery areas for fisheries by inundation and coastal erosion in tropical Asia [AR4. 30 Chapter 10, 10.4.1.4]. Changes in the hydrological cycle, and therefore also changes in the water resources have 31 been observed with a noticeable regional variability in all of Asia [AR4, Chapter 10, 10.2.4.2]. One of the most 32 pressing environmental problems in South and South-East Asia will be the expansion of areas under severe water 33 stress as the number of people living under severe water stress is projected to increase substantially in absolute terms 34 [AR4, Chapter 10, 10.4.2.3]. Oceanic, coastal, and other natural ecosystems have suffered degradation as a result of 35 global warming, sea-level rise and changes in intensity and amount of precipitation [AR4, Chapter 10, 10.2.4.3; 36 10.2.4.4]. Projections show that all coastal areas in Asia are facing an increasing range of stresses and shocks, the 37 scale of which now poses a threat to the resilience of both human and environmental coastal systems, and could be 38 additionally exacerbated by climate change [AR4, Chapter 10, 10.4.3.1]. Many plant and animal species are at risk 39 to become extinct as a consequence of the combined effects of climate change and habitat fragmentation (medium 40 confidence) [AR4, Chapter 10, 10.2.4.5; Chapter 10, Executive Summary]. Central, East, South and South-East Asia 41 reported deaths and disorders from heat waves and outbreaks of infectious diseases linked to rising temperatures and 42 rainfall variability, particularly in low-income areas with poor water and sanitation safety (medium confidence) 43 [AR4, Chapter 10, 10.2.4.6; Chapter 10, Executive Summary]. Substantial direct impacts on public health and 44 livelihood can be expected also in the future due to possible increases in climate change related diseases, as well as 45 heat stress [AR4, Chapter 10, 10.4.5]. Climate change is also expected to adversely affect sustainable development 46 capabilities of most Asian developing countries by aggravating pressure on natural resources and the environment in 47 addition to factors such as rapid urbanization, industrialization and economic development (high confidence) [AR4, 48 Chapter 10, 10.7; Chapter 10, Executive Summary].

49

50 Vulnerable areas. Regions of South and South-East Asia were reported as vulnerable to climate change, due to the

- exposure of their population to severe water stress [AR4, Chapter 10, 10.4.2.3]. Furthermore, the same regions are 51
- 52 expected to experience higher endemic morbidity and mortality due to diarrheal disease related to climate change 53
- (high confidence) [AR4,Chapter 10, 10.4.5; Chapter 10, Executive Summary]. Increases in coastal water

1 10, Executive Summary]. Crop yields in South and West Asia could decrease by a third by the middle of this

2 century (medium confidence) [AR4, Chapter 10, 10.4.1.1, Chapter 10, Executive Summary]. Glaciers over Tibetan

3 Plateau are projected to shrink at an accelerated pace, thus possibly increasing the number and intensity of glacial

4 melt-related floods, slope destabilization and a decrease in river flows as glaciers recede (medium confidence)

- 5 [AR4, Chapter 10, 10.2.4.2, 10.4.2.1; Chapter 10, Executive Summary]. Projected sea-level rise would result in
- significant losses of coastal ecosystems, along with increased risk of flooding on the coasts of South and South-East
 Asia (high confidence) [AR4, Chapter 10, 10,4,3,1; Chapter 10, Executive Summary]. Sea-level rise and declining
- 8 river runoff, coupled with extreme events such as flooding and intensifying storm surges, would have adverse
- 9 impacts on human settlements, aquaculture industry and infrastructure of Asia's densely populated megadeltas (high
- 10 confidence) [AR4, Chapter 10, 10.4.3.2; SREX, Chapter 4, 4.4.3]. Stability of wetlands, mangroves and coral reefs
- around Asia is likely to be increasingly threatened (high confidence) [AR4, Chapter 10, 10.4.3.2, 10.6.1; Chapter 10,
- 12 Executive Summary].13

14 Adaptive strategies. Adaptive strategies for the agricultural sector that have been identified in AR4 are intended to 15 increase adaptive capacity by modifying farming practices, improving crops and livestock through breeding, 16 investing in new technologies and infrastructure, making changes in management philosophy, through education and 17 the provision of climate change-related information [AR4, Chapter 10, 10.5.1]. In the water sector, dealing with 18 water use inefficiency, and promotion of recycled water was found useful in many agricultural areas in Asia [AR4, 19 Chapter 10, 10.5.2]. Along the coast, protection, such as dike heightening and strengthening, is considered to be 20 important in responding to sea-level rise [AR4, Chapter 10, 10.5.3]. Most forests in Asia would benefit from 21 comprehensive inter-sectoral programs that combine measures to control deforestation and forest degradation [AR4, 22 10.5.4]. Implementation of monitoring and warning systems would be helpful in reducing the impacts of climate 23 change of human health [AR4, Chapter 10, 10.5.5]. Effective adaptation and adaptive capacity in Asia, particularly 24 in developing countries, will continue to be limited by several ecological, social and economic, technical and 25 political constraints [AR4, Chapter 10, 10.5.7]. These constraints also include alterations of the physical 26 environment, as well as the adaptive capacities of some ecosystems, spatial and temporal uncertainties associated 27 with forecasts of regional climate, limited national capacities in climate monitoring and forecasting, and lack of 28 coordination in the formulation of responses [AR4, Chapter 10, 10.5.7]. Countries of Asia facing serious domestic 29 conflicts, pervasive poverty, hunger, epidemics, terrorism and other urgent and pressing concerns may not view climate change and the need to implement adaptation as immediate priority [AR4, Chapter 10, 10.5.7]. Slow change 30 31 in political and institutional landscape, and existing legal and institutional framework remains inadequate to 32 facilitate implementation of comprehensive and integrated responses to climate change [AR4, Chapter 10, 10.5.7]. 33 In order to address such constraints the following measures would be of use. Improving access to high-quality information about the impacts of climate change, adaptation and vulnerability assessment by setting in place early 34 35 warning systems and information distribution systems to enhance disaster preparedness; reducing the vulnerability 36 of livelihoods and infrastructure to climate change; promoting good governance including responsible policy and 37 decision making; empowering communities and other local stakeholders so that they participate actively in 38 vulnerability assessment and implementation of adaptation; and mainstreaming climate change into development 39 planning at all scales, levels and sectors [AR4, Chapter 10, 10.5.7].

40 41

42 24.3. Observed and Projected Change43

44 24.3.1. Observed Climate Trends and Variability

45 46 Temperature. In accordance with the findings of AR4, increasing trends in annual mean temperatures have been 47 observed across most of Asian region including the Tibetan Plateau during the 20th century, with the warming trend 48 continuing into the new millennium (see Table 24-2). Several studies pointed out the contribution of urban heat 49 island to the increase in annual mean temperatures. Despite a limited amount of information, a stronger upward 50 trend is observed for the winter mean temperatures, as compared to the summer mean in East Asia, as well as in 51 Bangladesh, Nepal, and over eastern Khengay and across Khentey Mountains, Mongolia. On the other hand, decreasing trends were observed for the summer diurnal temperature range in North-Western part of Kashmir, India 52 53 (Roy and Balling, 2005), and the mean minimum temperature in Karachi, Pakistan (Sajjad et al., 2009).

54

1 [INSERT TABLE 24-2 HERE

2 Table 24-2: Summary of key observed past and present climate trends and variability.]

3 4 Precipitation. Annual mean precipitation trends are characterized by strong variability, with both increasing and 5 decreasing trends observed throughout Asian regions (see Table 24-2). In India, Japan, and Kazakhstan no clear 6 national trend was observed, however, on a subnational level both positive and negative trends were observed. The 7 amount of summer total precipitation shows an increasing trend in South-East and North-West China and a 8 decreasing trend over Central China (Yao et al., 2008).

9 10

11 24.3.2. Observed Changes in Extreme Climate Events 12

13 Temperature extremes. Increasing tendencies are observed in temperature extremes (see Table 24-3). Mean 14 maximum temperatures show increasing trend in the number of warm days, and a decreasing trend in cold days has 15 been observed throughout Asia during the late 20th century (medium confidence) (SREX, Chapter 3, Table 3-2]) 16 Mean minimum temperatures show an increasing tendency on the continental scale, as observed in overall increase 17 of warm nights and decrease in the number of cold nights (medium confidence) (SREX, Chapter 3, Table 3-2).

19 **[INSERT TABLE 24-3 HERE**

20 Table 24-3: Summary of observed changes in extreme events and severe climate anomalies.]

21 22

18

Heat waves. Trends in heat waves displayed noticeable regional variability (see Table 24-3). Increases in the warm

23 spell duration index were observed overall in North Asia, in few parts of Central Asia, West Asia, and northern 24 China, while decreasing trends were recorded in southern China, and a few areas of North Asia (medium

25 confidence) (SREX, Chapter 3, Table 3-2).

26

27 *Heavy precipitation.* Regionally and sub-regionally varying trends were observed in heavy precipitation over the 28 Asian continent, however, there is insufficient evidence or inconsistent trending for the South and South-East Asian 29 region, as well as the Tibetan Plateau (low confidence) (SREX, Chapter 3, Table 3-2). Decreasing trends were 30 observed for the West Asian region (medium confidence) (SREX, Chapter 3, Table 3.2).

31

32 Dryness. Spatially varying trends in dryness, indicated by different measures (Consecutive Dry Days, Soil Moisture 33 Anomalies, Palmer-Drought Severity Index), were observed within most Asian regions (low confidence) (SREX, 34 Chapter 3, Table 3-2). Overall tendency for increased dryness was reported in East Asia, with just a few areas 35 showing opposite trends (medium confidence) (SREX, Chapter 3, Table 3.2).

36 37

38 24.3.3. Socio-Economic Scenarios for Climate Modeling 39

40 Since the AR4 was published, high-resolution (approx. range between 20-40km) GCMs or RCMs have been 41 examined in accordance with the SRES, and the future scenarios for tropical-cyclone outbreaks and monsoon-related 42 changes in precipitation were reported based on the GCMs/RCMs. As mentioned in Working Group I Chapter XX 43 and Chapter XII of AR5, new climate scenarios were developed by inputting RCP data.

44

45 Under the process of assessing climate change for the purposes of AR5, scenarios of Representative Concentration 46 Pathways (RCPs) were developed, in which the wider range of potential future radiative forcing pathways were

- 47 presented. Subsequently, socio-economic and climate scenarios have been developed in parallel by utilizing the RCPs (Moss et al., 2010).
- 48
- 49

50 As noted in Working Group III Chapter VI, the purpose of developing the four RCP scenarios was to compare the

- differences of climate change, climate change impacts, and emission pathways under different stabilization targets 51
- 52 (Moss et al., 2010). In addition, Shared Socio-economic Pathways (SSPs) and Shared Climate Policy Assumptions
- 53 (SPAs) have also been developed to provide the scenario elements such as Economic Growth, Globalization,

Distribution/ Equity, Environmental Ethics and Values, Institutions and Governance, Technological Change and Access, Population and Demographics.

24.3.4. Projected Climate Change

The projected changes for a variety of climate parameters such as temperature, precipitation, temperature extremes,
heavy precipitation, dryness, and sea levels do not show explicit regional trends (see Tables 24-4 and 24-5). Further
information is being collated for specific sub-regions.

11 [INSERT TABLE 24-4 HERE

12 Table 24-4: Summary of projected changes for a variety of climate parameters.]

14 [INSERT TABLE 24-5 HERE

15 Table 24-5: Description of climate parameter abbreviations used in Tables 24-2 to 24-4.]

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24.4. Observed and Projected Impacts, Vulnerabilities, and Adaptation

- 20 24.4.1 Freshwater Resources
- 22 24.4.1.1. Sub-Regional Diversity

23 24 The water sector in Asia is significantly vulnerable to shifts in climate, due to the dependence of its huge 25 agricultural sector on precipitation, river runoff, and groundwater (see Table 24-6). Among the countries of Asia, 26 twenty have renewable annual per capita water resources in excess of 3,000 m3, eleven are between 1,000 and 3,000 27 m3, and six are below 1,000 m3 (there are no data from the remaining six countries). Hence, adequate water supply 28 is one of the major challenges in Asia, particularly Central Asia (Vorosmarty et al., 2010). Growing demand for 29 water is driven by soaring population, increasing urbanization, and thriving economic growth. Arid countries of the 30 Middle East and Central Asia face major challenges in ensuring fresh water supply, which will continue to decline 31 with the decrease in precipitation, groundwater recharge and surface runoff. Mismanagement of water resources is 32 increasing tension among five Central Asian states of the former Soviet Union - Kazakhstan, Kyrgyzstan, 33 Turkmenistan, Uzbekistan, and Tajikistan (Lioubimtseva and Henebry, 2009; Siegfried et al., 2010). 34

- 35 [INSERT TABLE 24-6 HERE
- 36 Table 24-6: Summary of observed and projected impacts in the water sector.]
- 37 38
- 39 24.4.1.2. Observed Impacts
- 40

40 Climate change have impacts on the water availability in arid and semi-arid areas (Brutsaert and Sugita, 2008), in

42 South China (Jiang et al., 2008), in Northwest Himalaya (Bhutiyani et al., 2008). The surface water resources of

- 43 Central Asia are primarily generated in mountain glaciers. Increased runoff from shrinkage of glacier is observed in
- Himalayas (Zhang et al., 2011) and Central Asia mountains due to increased temperature and this currently has

45 positive impact on the water availability (Casassa, G., P. Lopez, et al., 2009; Shrestha and Aryal, 2011). Apart from

- 46 water availability, precipitation is highly correlated with surface water quality, represented by dissolved oxygen, pH,
- 47 conductivity (Prathumratana *et al.*, 2008 in Delpla *et al.*, 2009), dissolved salt content (Huang et al., 2009),
- 48 concentrations of phosphorus related to agricultural activities (Park et al., 2010), carbon and nutrients (Zhang *et al.*,
- 49 2007b; Goldsmith *et al.*, 2008), which may increase health risk (Tornqvist et al., 2011). It is also noticeable that the
- 50 water quality in groundwater is also related to climate change (Thakue and Ojha, 2010; Winkel et al., 2011; Fendorf
- 51 et al., 2010). Increased frequency of flooding and droughts are also observed recently related to the climate warming.
- 52 Water crisis in Asian countries is also caused by poor management (Biswas and Seetharam, 2008).
- 53 54

24.4.1.3. Projected Impacts

Projected impacts of future climate change on water availability (considering the future demand) in Asia differ substantially among river basins shown by 5 GCMs for A1B scenario (Immerzeel et al., 2010). The water demand in most Asian countries is gradually increasing because of increases in population, irrigated agriculture (Lal, 2011) and growth in the industrial sectors. Tropical Asia will experience severe dry and wet spells that will reduce water supply reliability and increase chances of flooding. Even though precipitation in Northern and Temperate Asia is expected to increase overall (Park et al., 2010), socio-economic development will pose a challenge to freshwater resources. Projections (A2 scenario from multi-GCMs) suggest that throughout much of Russia a warmer climate would decrease water availability due to the increase of evaporation, but on the other hand precipitation would increase which tends to increase water availability (Alcamo et al., 2007). In China, the projection (A2 scenario from PRECIS) suggests that there will be insufficient water for agriculture in China in 2020s and 2040s due to the increases in water demand for non-agricultural uses although positive trends of precipitation (Xiong et al., 2010). In a study of the Mahanadi River Basin, the future water availability projection (A2 from CGCM2) indicated an escalating trend in excess river runoff (runoff after meeting water demand) thereby increasing the future possibility of floods for the month of September, yet the outcomes for April indicate an accelerating water scarcity (Asokan and Dutta, 2008). In the Ganges effects of climate change could become large enough to offset the large increases in demand in a +4°C world, due to a projected large increase in rainfall (2°C and 4°C temperature increase from ensemble GCMs; Fung et al., 2011). Given the already very high level of water stress in many parts of Central Asia, projected temperature increases and precipitation decreases (SRES scenarios from IPCC AR4 23 models) in the western part of Kazakhstan, Uzbekistan, and Turkmenistan could exacerbate the problems of water shortage and distribution (Lioubimtseva and Henebry, 2009). Considering the dependence of Uzbekistan's economy to its irrigation agriculture, which is consuming more than 90% of the available water resources of the Amu Darya basin, climate change related impacts on river flows would also strongly affect the economy (Schlüteret al., 2010). Coastal fresh water resources might reduce over the next century in Asia except for South-East Asia and the vulnerable areas

fresh water resources might reduce over the next century in Asia except for South-East Asia and the vulnerable area include South India and Bangladesh region and China while Japan stands in a good place due to its higher availability of fresh groundwater and lower population density (A2 scenario from HadCM3, Ranjan *et al.*, 2009).

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24.4.1.4. Vulnerabilities to Key Drivers

It is likely that river discharge will be influenced by rainfall change, rapid melting of snow and frozen soil in the watershed (Tachibana et al., 2008) associated with the Asian monsoon change (Jian et al, 2009). Water management in river basin needs to be coordinated among countries, for example water management in the Syr Darya river basin relates to Kyrgyzstan, Tajikistan, Uzbekistan, Turkmenistan, Kazakhstan (Siegfried et al., 2010).

38 24.4.1.5. Adaptation Options

39 40 Asia is by far the largest user of irrigation water in terms of volume. During the second half of the 20th century, 41 Asia has built many reservoirs and almost tripled its surface water withdrawals for irrigation. Reservoirs partly 42 mitigate the seasonal differences and increase water availability for irrigation (Tyler and Fajber, 2009). However, 43 they might not be able to continue the same supply because of a change in reservoir inflow due to effects of climate 44 and socioeconomic change. On the other hand, reservoirs might have an increasing role in meeting future water 45 requirements in regions where water stress is an issue of distribution rather than of absolute shortage (Biemans et 46 al., 2011). To adapt the climate change impact on water resource, many Asian countries apply water saving 47 technologies in irrigation (Ngoundo et al., 2007; Tischbein et al., 2011) and other consumptive purposes (Fleskens et 48 al., 2007), changes of crop types to drought tolerant crops (Thomas, 2008; Zhao et al., 2010), increasing water 49 supply (Sadoff and Muller, 2009), and improved management (Kranz et al., 2010). It is found that in monsoonal 50 Asia, development of water control systems have contributed to improved rice harvests (Hatcho et al., 2010). 51 52 For dealing with flooding, four strategies (a new flood map, an early warning system, a relief programme, and more

53 community education) are developed in the Sarawak River system in Malaysia to reduce the excessive flood lose

54 (Mah *et al.*, 2011). Hazard mapping could help both decision-makers and local communities to understand the

1 current situation and, through this, it would be possible to anticipate or assess the flexibility to adapt to future

2 changes through proper planning and technical design. Examples include mapping in the Himalayan region 3 (Eriksson et al., 2009) and proposed investment in river regulation and storage in Nepal to control floods and to

4 augment low-season flows in India and Bangladesh in the Ganges River Basin (Sadoff and Muller, 2009). 5

6 The equitable sharing of water and the drought proofing of rural livelihoods will require an increasing physical 7 capacity to store water (van der Zaag and Gupta, 2008). Moreover, policy processes in the current water 8 management regime are strongly shaped by informal institutions and the lack of enforcement of formal regulations. 9 The high degree of centralization of the management regime (Webster and McElwee, 2009) and the lack of vertical 10 integration are possible explanations for the rather low adaptive capacity (Schlüter et al., 2010). Legal aspects in 11 water management are also suggested to be considered in South Aisa (Uprety and Salman, 2011; D'Agostino and 12 Sovacool, 2011).

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24.4.2. Terrestrial and Inland Water Systems

24.4.2.1. Sub-Regional Diversity

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19 Asia supports examples of all the major natural terrestrial ecosystem types on earth, with the predominant types 20 differing in sub-regions. North Asia is a region of tundra, boreal forests and grasslands, Central and West Asia are 21 dominated by desert and semi-desert ecosystems, and the Tibetan Plateau is covered in a variety of largely treeless 22 alpine ecosystems. These four sub-regions have relatively low human population densities in most areas, except for 23 parts of Central Asia, and are still largely covered in natural ecosystems, although some of these have been 24 extensively modified. In the three remaining sub-regions, in contrast, natural ecosystems have been completely 25 replaced over large areas by human-dominated landscapes. The major natural ecosystems of East Asia included 26 temperate deciduous and subtropical evergreen forests, giving way to boreal forest in the northeast and to grasslands 27 and deserts in the west. South Asia and Southeast Asia were largely covered in tropical forests, with deciduous and 28 semi-evergreen forests most extensive in South Asia and evergreen rain forests more important in Southeast Asia. 29 South Asia also has extensive semi-desert areas in the west and northwest, and a variety of alpine ecosystems in the 30 north, while Southeast Asia supports a small area of alpine vegetation above the treeline in New Guinea. Asia 31 includes several of the world's largest river systems (Ganga-Brahmaputra-Meghna, Yangtze, Ob, Amur, Lena, 32 Yenisei, Mekong) with their associated deltas, as well as the world's deepest and most biological diverse freshwater 33 lake, Lake Baikal, the semi-saline Caspian Sea, and the saline and now greatly shrunken Aral Sea. Other major 34 saline (endorheic) lakes in central and west Asia include Balkhash (SE Kazakhstan), Issyk-Kul (E Kyrgyzstan), 35 Urmia (NW Iran), and Qinghai Lake (China).

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38 24.4.2.2. Observed Impacts

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40 Temperatures have shown a largely consistent rise across Asia since 1970, but changes in precipitation have been 41 complex and varied (WGI AR5 ZOD). In general, observations of biological changes in terrestrial ecosystems 42 consistent with the impacts of climate change are more common in the cold and/or arid north and west of the region, 43 and at high altitudes, where rising temperature and, in some areas, increasing precipitation have relaxed constraints 44 on the growth of plants and the distributions of both plants and animals. In contrast, there have been very few 45 reports from the tropical lowlands of impacts and none that can be linked to recent climate change with high 46 confidence. Changes in inland water systems have also been widely reported, but the impacts of climate change have 47 been difficult to disentangle from natural variability and a wide variety of other, concurrent, human impacts (Bates 48 et al., 2008; Wang et al., 2011c; Zheng, 2011).

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50 Phenology. The most widely reported impacts attributed to the observed climate trends have been changes in the 51 timing of life-history events, including leafing, flowering, and leaf fall in plants, the breeding periods of animals, the

52 emergence of insects, and the arrival and departure of migrant birds (e.g. Soja et al., 2007; Doi, 2007; Doi and

- 53 Katano, 2008; Sokolov and Gordienko, 2008; Primack et al., 2009; Fujisawa and Kobayashi, 2010; Bai et al., 2010;
- 54 Choi et al., 2011; Ge et al., 2011; Ogawa-Onishi and Berry, in press; Section 28.2.3.2., Chapter 28 WG2 AR5 ZOD).

1 Trends in phenological timing are consistent with the impacts of regional warming are widespread in northern China

2 and Japan, including spring advances and autumn delays, particularly for plants. However, consistency is lower 3 elsewhere and also for animals, where spring delays in phenology have been reported for some species (e.g., barn

- 4 swallows in Korea; Lee et al., 2011).
- 5

6 Plant growth, greenness and NPP. Recent changes in the growth rates of plants have also been reported (e.g. 7 Feeley et al., 2007, Nock et al., 2011) and where long records are available from tree rings, these changes can be 8 more confidently attributed to recent climate change (e.g. Duan et al., 2010; Dulamsuren et al., 2010a; Sano et al., 9 2010; Yang et al., 2010; Shishov and Vaganov, 2010; Li et al., 2012). Changes in satellite-measured 'greenness' 10 (NDVI) reflect changes in plant growth over larger areas. For temperate East Asia (30-80°N), NDVI data show 11 growing season length increased by 9.5 days/decade in the period 1982-2000, with the biggest change at the 12 beginning of the season, but that part of this increase was reversed during 2000-2008 (Jeong et al., 2011). On the 13 Tibetan Plateau, warmer springs lead to an advance in greening while warmer winters cause a delay, leading to an 14 overall delay in recent spring phenology (Yu et al., 2010).

15

16 *Changes in the distributions of species and biomes.* Also widely reported are changes in species distributions:

17 generally upwards in elevation (e.g. Soja et al., 2007; Round and Gale, 2008; Bickford et al., 2010; Kharuk et al.,

2010 a, b, e; Moiseev et al., 2010; Chen et al., 2011; Jump et al., 2012) or polewards (e.g. Tougou et al., 2009; 18 19

Ogawa-Onishi and Berry, in press) in response to recent warming. Movements of dominant plant species can

20 eventually lead to changes in the distributions of major vegetation types (biomes). Evidence for biome shifts has so

21 far been reported only from the north of the region and at high altitudes, where it involves trees invading treeless 22

tundra, steppe or alpine meadows, or the invasion of the forest understory by species from adjacent biomes (Soja et 23 al., 2007; Kharuk et al., 2006; Bai et al., 2011; Singh et al., 2012; Ogawa-Onishi and Berry, in press). In

24 Uttarakhand, in the Indian Himalayas, the treeline has moved upwards into the alpine zone by an average of 388 m

25 between the 1970s and 2006 (Singh et al., 2012). The position of the ecotone between boreal forest and tundra is

26 controlled largely by air temperature during the growing season and annual precipitation, but forest fires can also 27 catalyze change (Soja et al., 2007). Soil moisture and light are the main factors governing the forest-steppe ecotone,

28 but competition between trees and grasses, as well as fires, are also important (Soja et al., 2007; Zeng et al., 2008;

29 Dulamsuren et al., 2010 a, b; Eichler et al., 2011).

30

31 Larch-dominated forest occupies about half the area of Siberia. Invasion of dark needle conifers (DNC, e.g. Siberian 32 pine, spruce and fir) and birch into the larch habitat over the last three decades has been observed (Kharuk et al.,

33 2010c). Siberian pine and spruce have high invasion potential both along the margin and in the centre of the larch-

34 dominated zone. This phenomenon could be attributed to increases in temperature and precipitation. Winter 35 temperature regime is important for the Siberian pine regeneration survival. The process is wildfire dependant. On

36 the western and southern margins of this zone, DNC regeneration has formed a second layer in the forest canopy.

37 Eventually, the larch in the overstorey could be replaced by these young DNC trees. In mixed stands, both larch and

- 38 fir growth have increased over time, but the fir growth increase has been larger which may presage a shift in
- 39 competitive balance between these species. Overall, it is *likely* that prevalence of evergreen conifers in areas

40 currently dominated by deciduous larch species is increasing (Kharuk et al., 2010c, d; Osawa et al., 2010; Lloyd et

41 al., 2011). At the same time, climate change has driven larch stand crown closure, and larch invasion into tundra at a

42 rate of 3–10 m/year was observed in the northern forest-tundra ecotone in Siberia in the last three decades of the

20th century (Kharuk et al., 2006). Shrub expansion in arctic tundra as result of an increase in shrub growth, 43

44 infilling of existing patches and the shrub line advancing into tundra is another change in the forest-tundra ecotone

45 of Northern Asia that has been attributed to climate change (Myers-Smith et al., 2011; Blok et al., 2011; Section 46 28.2.3.1., Ch 28 WG2 AR5).

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48 The forest-steppe ecotone in the western Khentey Mountains, northern Mongolia, has experienced a significant

49 increase in summer temperature and decrease in summer precipitation since 1961. Siberian larch tree-ring analysis

- shows a strongly decreasing annual increment since the 1940s (Dulamsuren et al., 2010a). Regeneration of larch 50
- decreased as well and is now virtually lacking in this forest. Studies on a wider scale show a great deal of 51
- 52 heterogeneity in the responses of Mongolian taiga forests to recent climate changes, but declines in larch growth and
- 53 regeneration are more widespread than the opposite trend, suggesting a net loss of forest will occur in future
- 54 (Dulamsuren et al., 2010b).

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4 2010; Zhao et al., 2010). Russia contains more permafrost than any other country: more than half of the Russian part 5 of Northern Asia lies in permafrost zones, which constitutes a significant portion of the Northern Hemisphere 6 permafrost area (FNCRF, 2010). Monitoring in most of the permafrost observatories in Asian Russia shows 7 substantial warming of permafrost during the last 20 to 30 years (Romanovsky et al., 2008; 2010). Typical 8 magnitude of warming varied from 0.5 to 2°C for different locations at the depth of zero annual amplitude. The 9 main warming occurred between the 1970s and 1990s, with no significant warming after 2000. However, since 10 2007-2008 warming has resumed at many locations predominantly near the Arctic coasts. In Northwest Siberia, new 11 closed talks (areas of unfrozen ground) and an increase in the depth of preexisting talks have been observed during 12 last 20 to 30 years. Permafrost formed during the Little Ice Age is thawing at many locations and Late Holocene 13 permafrost has begun to thaw at some undisturbed locations in northwest Siberia. Permafrost thawing is most 14 noticeable within the discontinuous permafrost domain in Northern Asia, while in the continuous permafrost zone it 15 is starting to thaw at some limited locations. As a consequence, the boundary between continuous and discontinuous 16 permafrost zones is moving northward (Romanovsky et al., 2008; 2010). Over many thousands of years, the soil

Permafrost. Degradation of permafrost, including reductions in area and increased thickness of the active layer, has

been reported from parts of Siberia, Central Asia, and the Tibetan Plateau (Romanovsky et al., 2010; Wu and Zhang,

- 17 layer and bogs in the permafrost zone of Northern Asia have been accumulating huge amounts of organic matter. As
- permafrost thaws, reinforcement of the greenhouse effect is possible due to growing emissions of greenhouse gases
- 19 (see Section 4.3.4.4., Ch 4 and Section 19.3.5., Ch 19, WG2 AR5).
- 20

The Qinghai-Tibet Plateau (QTP) and Central Asian region, including parts of Southern Siberia, Mongolia, Western China, Kazakhstan, and adjacent countries/regions, represent the largest area underlain by mountain permafrost in the world. Ongoing monitoring at numerous sites across the QTP regions over the past several decades has revealed significant permafrost degradation caused by climate warming and human activities such as deforestation, forest fire, road construction and grazing: areas of permafrost are shrinking, the depth of the active layer is increasing, the

- lower limit of permafrost is rising, and the seasonal frost depth is thinning (Zhao *et al.*, 2010; Li *et al.*, 2008). The
- 27 lower altitudinal limit of permafrost has moved up by 25 m in the north during the last 30 years and between 50 and
- 28 80 m in the south over the last 20 years in accord with long-term temperature measurements. Ground temperature at
- 29 a depth of 6 m in 2001 has been higher by about 0.1 0.3°C than in 1996 according to data taken from seven natural
- 30 sites on the Plateau (Cheng and Wu, 2007; Li *et al.*, 2008). Over the period from 1995 to 2007, the mean rate of 31 increase of the active layer thickness (ALT) uses 7.5 cm/uses (Wu and Zhang 2010). Council temperature
- 31 increase of the active layer thickness (ALT) was 7.5 cm/year (Wu and Zhang, 2010). Ground temperatures at the 32 bottom of the active layer warmed on average by 0.06°C/year over the past decade (Zhao *et al.*, 2010). In the alpine
- headwater regions of the Yangtze and Yellow Rivers, rising temperatures and permafrost degradation have resulted
- in lower lake levels, drying swamps and shrinking grasslands (Cheng and Wu, 2007; Wang *et al.*, 2011a).
- 35

In the Kazakh part of Tien Shan Mountains, the increase in permafrost temperature during 1974–2009 at depths of 14–25 m varied from 0.3°C to 0.6°C. The average active layer thickness (ALT) increased by 23% in comparison to

- the early 1970s. In the eastern Tien Shan Mountains, in the headwaters of the Urumqi River, China, significant
- 39 permafrost warming took place as the air temperature increased (Marchenko *et al.*, 2007; Zhao *et al.*, 2010). In
- 40 Mongolia, mean annual ground temperature (MAGT) at 10–15 m depth over the past 10–40 years increased on
- 41 average by 0.02–0.03°C/year in the Hovsgol Mountain region, and by 0.01– 0.02°C/year in the Hangai and Hentei
- 42 Mountain regions. During the past 15–20 years permafrost warming was greater than during the previous 15–20
- 43 years (1970s–1980s). The average rate of increase in MAGT in Mongolia was about 0.15°C/decade (Sharkhuu *et al.*,
- 44 2008; Zhao *et al.*, 2010).
- 45 46
- 47 24.4.2.3. Projected Impacts48

The projected impacts in the literature assessed here include extrapolations from the observed trends and inferences from a variety of modeling approaches, based on projected climate change and projections for other factors, such as rising carbon dioxide levels and land-use changes.

52

53 *Distributions of species and vegetation.* The current distribution of vegetation across the region is controlled 54 primarily by climate (particularly temperature and rainfall, and their seasonality; Tang *et al.*, 2009), modified over 1 large areas by soils, permafrost, topography, and a variety of human impacts. In the longer term, therefore, climate 2 change is expected to change this distribution (e.g. Wang et al., 2011b). However, the rate at which this change in vegetation is realized will be constrained by many factors, including seed dispersal, competition from established

- 3 4 plants, rates of soil development, and habitat fragmentation. As explained in section 24.3.4, climate simulations for
- 5 Asia strongly suggest that the warming trend will continue, but projections for precipitation are still uncertain. In
- 6 general, the changes in both temperature and precipitation are expected to be greater in the north and west of the
- 7 region, which are also the areas with the least fragmented vegetation. These changes in climate will lead to large and
- 8 relatively predictable changes in the distribution of potential natural ecosystems (Ni, 2011; Wang et al., 2011b;
- 9 Tchebakova *et al.*, 2011; Insarov *et al.*, in press), although the transitional stages will be less predictable.
- 10
- 11 In Northern Asia, if current climate projections are correct, it is *likely* that the boreal forest will expand northward
- 12 and eastward, and the tundra area will decrease, during the 21st century (Golubyatnikov and Denisenko, 2007;
- 13 Korzukhin and Tcelniker, 2010; Lucht et al., 2006; Sitch et al., 2008; Tchebakova et al., 2010; Woodward and 14
- Lomas, 2004). However, for a shorter time horizon, some forest retreat and tundra advance by 2020 in Central 15 Siberia have been projected (Tchebakova et al., 2011). Because models vary in accordance with their structure as
- 16 well as biome classifications, climatic projections, CO_2 level and other characteristics used as inputs, the magnitude
- 17 of the forest expansion varies greatly across models: Tchebakova et al. (2009) and Lucht et al. (2006) project that
- 93-100% of tundra area will be covered by boreal forest at the end of 21st century, Kaplan and New (2006) predict a 18
- 19 42% reduction in tundra area between 2026 and 2060, whereas Golubyatnikov and Denisenko (2007) estimate that
- 20 97% of tundra will remain unaltered by the mid-21st century.
- 21
- 22 The combination of boreal forest expansion and the continued invasion of the existing larch-dominated forest by
- 23 dark-needle conifers could lead to a situation where larch reaches the Arctic shore, a phenomenon that has happened
- 24 previously in the Holocene, whereas the traditional area of larch dominance will turn into mixed taiga forest
- 25 (Kharuk, 2006, 2010d). Both replacement of summer-green conifers (larch) with evergreen conifers (DNC) and
- 26 expansion of boreal forest and shrubs into regions now occupied by tundra decrease albedo. This change would
- 27 cause heating of the atmosphere, a response that, in its turn could possibly accelerate the replacement of larch by
- 28 DNC and of tundra by boreal forest (McGuire et al., 2007; Kharuk et al., 2006, 2010d). Energy budget feedback to
- the regional summer climate from the tundra to forest transition is estimated at 5.0 Wm⁻² (McGuire *et al.*, 2007). 29
- 30

31 The direction and rate of change in the extent of steppe vegetation is less clear, in part because of uncertainty in

- 32 precipitation trends. One projection is that steppe area will increase by 27% for the decade beginning in 2090
- 33 (Tchebakova et al., 2010) while another is that it will decrease by up to 65% for late 2030s-early 2050s
- 34 (Golubyatnikov and Denisenko, 2007). Reasons for the differences between these estimates include different
- 35 projection horizons and vegetation classifications used. Increasing aridity may expand the deserts of northern China, 36 and push the steppe to the northeast (Zhang G.G. et al., 2011), while a retreat of the southern limit of the taiga would 37 expand the steppe area in the north (Dulamsuren et al., 2010b).
- 38

39 The forest regions of East Asia are expected to remain forested, but climates suitable for subtropical evergreen forest

- 40 will expand north into the deciduous forest zone (Wang et al., 2011b). As observed elsewhere in the world, however,
- 41 vegetation changes within lowland forest regions are expected to lag behind climate change by decades or even
- 42 centuries, as fragmentation limits seed dispersal and long-lived forest dominants persist (e.g., Bertrand et al., 2011;
- 43 Zhu et al., 2012). For example, climate models predict a large increase in the potential habitat for the evergreen
- 44 broad-leaved tree species Quercus acuta in Japan, but short-distance seed dispersal by rodents will limit the ability
- 45 of this species to occupy new areas (Nakao et al., 2011).
- 46
- 47 Impacts in Central and West Asia will depend critically on the changes in precipitation, which are still highly
- 48 uncertain. Projections for China from an atmospheric-vegetation interaction model under the SRES B2 scenario
- 49 show that the arid northwest of the country is the most vulnerable ecoregion, with severe damage to desert
- 50 ecosystems possible (Wu et al., 2007, 2010). Forest is expected to expand on the more mesic parts of the Tibetan
- 51 plateau and there is expected to be a general northwestern shift of all vegetation zones (Wang et al., 2011a). In the
- drier areas of the plateau, the loss of permafrost may contribute to desertification (Cheng and Wu, 2007). In the 52
- 53 tropics, although the expected rates of warming are less, the relatively small annual temperature range means that by
- 54 the end of the century the tropical lowlands are *likely* to experience temperatures daily that are outside the current

range of extremes (Beaumont *et al.*, 2010). The potential impacts of these novel climatic conditions are largely unknown (Corlett, 2011). If the frequency and severity of droughts increases, as some projections suggest, this is

- unknown (Corlett, 2011). If the frequency and severity of droughts increases, as some projections sugg
 likely to interact with forest fragmentation and logging to increase fire risk (van der Werf *et al.*, 2008).
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Fewer studies have projected impacts on animals. Hughes *et al.* (2012) projected the effects of both climatic (A2 and B1 scenarios) and vegetation changes on the distribution and diversity of bats in SE Asia. All projections predicted widespread losses in bat diversity and large reductions in the distribution of most bat species. Projections for the

- 8 potential ranges of 63 species of galliform birds (pheasants, partridges and their relatives) in China (A2 scenario,
- 9 2071-2100) showed large (>50%), mostly northward, range shifts for 29 species (Li *et al.*, 2010), while projections
- for the 13 species of nuthatches (Sittidae) in Asia (A2 and B2 scenarios, 2040-2069) found that most ranges would
- retract along their southern fringes and at lower elevations, with the largest range contractions in SE Asia and peninsular India (Menon *et al.*, 2009). Projections for the distributions of 161 butterfly species in Thailand (A2 and
- B2 scenarios, 2070-2099) suggested that species richness within currently protected areas will decline c. 30%, but
- 14 that these areas will continue to include a similar proportion of the highest priority sites for conservation 15 (Klorvuttimontara *et al.*, 2011).
- 16

17 Permafrost. In the Northern Hemisphere as a whole, a 20-90% decrease in permafrost area and a 50-300 cm increase in active layer thickness (ALT) is projected for 2100 by different models under SRES A1B, A2, B1 18 19 scenarios (Schaefer et al., 2011). The wide range of permafrost degradation projections may be result of different 20 scenarios used, intensity of land atmosphere feedbacks and of difference in model internal structures. In Asia, it is 21 *likely* that permafrost degradation during the 21st century will spread from the southern and low-altitude margins, 22 advancing northwards and upwards as numerous models predict, but rates of change vary greatly between different 23 model projections (Cheng and Wu, 2007; Riseborough et al., 2008; Romanovsky et al., 2008, with supplement; 24 Anisimov, 2009; Eliseev et al., 2009; Nadyozhina et al., 2010; Schaefer et al., 2011; Wei et al., 2011). The spatially 25 distributed permafrost model (Sazonova and Romanovsky, 2003) has been applied to the entire permafrost domain of Northern Eurasia, Central Asia and the QTP (Romanovsky et al., 2008, with supplement). If air temperatures 26 27 continues to increase in accordance with the MIT 2D climate model output for the 21st century (Sokolov & Stone 1998), that is 2.2°C warming by 2031-50 and 4.7°C by 2080-2099 compared with 1981-2000 (Romanovsky et al., 28 2008, with supplement), permafrost models show that permafrost that is presently discontinuous with temperatures 29 between 0 and -2.5° C will cross the threshold by the end of 21^{st} century and will be thawing actively. The most 30 31 intense permafrost degradation in Russia is projected for Northwest Siberia. According to this model, the Late 32 Holocene permafrost will be actively thawing everywhere except for the south of East Siberia and the Far East of 33 Russia by the middle of 21st century. Almost all Late Holocene permafrost will be thawing, and some Late Pleistocene permafrost will begin to thaw in Siberia by the end of 21^{st} century (Romanovsky *et al.*, 2008, with 34 35 supplement). Near-surface permafrost is expected to remain only in Central and Eastern Siberia and in Tibet in the 36 late 21st century. Depths of seasonal thaw are projected to exceed 1 m (2 m) in the late 21th century under the SRES 37 B1 (A1B or A2) scenario in these regions (Eliseev et al., 2009). 38

39 On the Qinghai-Tibet Plateau (QTP) and in northeastern China, substantial retreat of permafrost is expected during 40 the 21st century due to the combined influence of climatic warming and increasing anthropogenic activities. No 41 significant change will take place in permafrost conditions on the QTP over the next 20 to 50 years, but more than 42 half of the permafrost in the southern and eastern parts of the plateau may become relict and/or even disappear by 43 2100 according to modeling results (Cheng and Wu, 2007). The result of permafrost degradation can be ground 44 surface drying, and land desertification may become an important environmental issue for the QTP (Cheng and Wu, 45 2007). In northeastern China, the southern limit of permafrost is expected to shift northwards, the total permafrost 46 area to shrink, and the area of unstable permafrost to expand, with adverse consequences for associated wetlands and 47 forests (Sun et al., 2011; Wei et al., 2011). 48

49 *Inland Waters.* Climate change impacts on inland waters will continue to interact with a wide range of other human 50 impacts, including dam construction, pollution, and catchment land-use changes (see also Chapter 3, this volume).

50 Impacts, mendening dam construction, pointuon, and caterimient rand-use enarges (see also enapter 5, this volume). 51 Increases in water temperature will be the most pervasive impact on both living organisms and a wide range of

- temperature-dependent ecological, chemical, and physical processes. The dominance of ectotherms in aquatic
- ecosystems may make them particularly vulnerable to changing temperature, although direct evidence for this is
- 54 currently lacking for Asia (Dudgeon, 2011). The other major impact of climate change is likely to be on flow

1 regimes in running waters and consequently on riverine habitats and species that are sensitive to flow extremes

2 (droughts and floods). Regionally threatened natural habitats that depend on seasonal inundation, including

3 floodplain grasslands and freshwater swamp forests, will be particularly vulnerable (Maxwell, 2009; Bezuijen,

2011). Changes in river flow, in turn, have a direct impact on the freshwater to saltwater gradient where the river
 meets the sea, with reduced dry season flows combining with sea-level rise to increase saltwater intrusion in deltas

- 6 (Hamilton, 2010), although non-climatic human impacts will probably continue to dominate in most Asian estuaries
- 7 (Syvitski *et al.*, 2009). The unique ecosystem of Lake Baikal is expected to be impacted most by changes in ice
- 8 duration and transparency, followed by water temperature and wind mixing (Moore *et al.*, 2009).
- 9

10 Thresholds and irreversible changes. Specific thresholds for terrestrial and inland water systems have not yet been 11 identified. Studies of future climate change impacts on terrestrial ecosystems in China under the SRES B2 scenario 12 suggest that moderate to severe impacts will increase significantly when temperatures increase by more than 2°C, 13 but do not suggest a sharp threshold (Wu *et al.*, 2010). Species extinctions are the most likely irreversible change, 14 with species that are unable to track climate change as a resulted of limited dispersal ability, habitat fragmentation, 15 or non-climatic constraints, such as specialized soil requirements, most vulnerable (Heller and Zavaleta, 2009).

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18 24.4.2.4. Vulnerabilities to Key Drivers

20 Changes in temperature are the most robust predictions and the most pervasive climate impact, but the biological 21 consequences of the predicted changes are still poorly understood. Adverse impacts from rising temperature are 22 likely in the wetter areas of north Asia and at high altitudes, with permafrost melting impacting ecosystems across 23 large areas (Cheng and Wu, 2007; Tchebakova et al., 2011), but the impacts of higher temperatures in the tropical 24 and subtropical lowlands are still unclear (Corlett, 2011). The biodiversity of isolated tropical, subtropical, and 25 warm-temperate mountains may be most vulnerable to warming, because many species already have small 26 geographical ranges that will shrink further in a warming climate (Liu et al., 2010; Chou et al., 2011; La Sorte and 27 Jetz, 2011; Noroozi et al., 2011; Peh et al., 2011; Jump et al., 2012).

28

29 For much of Asia, increases in aridity, as a result of declining rainfall and/or rising temperatures, are the key 30 concern. Because aridity (decreased precipitation and soil moisture and increased frequency of severe droughts) is 31 projected to increase in the northern Mongolian forest belt during the 21st century (Sato *et al.*, 2007), the larch 32 covered area will likely be reduced (Dulamsuren et al., 2010a). This will have far-reaching consequences for 33 Mongolia's biodiversity and capacity to store water and carbon. It is likely it will also have significant 34 socioeconomic consequences because the economy depends on the sustainable exploitation of natural resources. 35 Even where mean rainfall remains adequate, any increase in drought frequency and/or severity will increase 36 vulnerability to human-caused fires. The frequency and scale of both natural and manmade fires have recently 37 increased in the tundra and taiga-tundra zones, as a result of warming, especially summer droughts (Kumpula et al., 38 2011; Nuttall 2005; Walker et al., 2011). Freshwater systems are also potentially vulnerable to increases in the 39 frequency and intensity of extreme events (droughts or floods), even if average conditions are unchanged (Hamilton, 40 2010).

- 41
- 42

43 24.4.2.5. Adaptation Options

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In view of the large uncertainties in the prediction of impacts and vulnerabilities, the focus so far has been largely on
 building resilience and enhancing the capacity of natural ecosystems for autonomous adaptation. Suggested
 adaptation strategies have often been generic (e.g. reducing non-climate impacts, monitoring climate impacts,

48 maximizing landscape connectivity, making protected area networks robust to future climate scenarios; Hannah,

- 49 2010; Shoo *et al.*, 2011; Klorvuttimontara *et al.*, 2011) rather than specific to local conditions, and, in most cases,
- 50 the adaptation measures adopted so far have been continuations of programs initiated for other reasons (e.g. China's
- 51 "Grain for Green Program" and "Green Wall policy"; Piao *et al.*, 2010). In northeastern China, where climate
- 52 change is expected to increase the risk of damaging forest fires, strengthening early warning and monitoring systems,
- paying attention to post-fire recovery, and the use of prescribed burning to reduce fuel loads are among the
- 54 suggested strategies for adaptation (Tian *et al.*, 2011). For Papua New Guinea, three general strategies have been

1 suggested for adapting biodiversity conservation to climate change: conserving the 'geophysical stage' (i.e., habitats

2 across the full range of physical settings, including combinations of elevation and geology); protecting 'climatic

3 refugia' (i.e., areas where climate change is expected to be relatively attenuated); and increasing landscape

4 connectivity (Game et al., 2011). More generally, there is increasing recognition of the need to incorporate climate 5 change adaptation into all forest conservation and development programs (e.g. in India; Chaturvedi et al., 2011).

6

Species distribution models are increasingly used to forecast future species distributions in the face of climate

7 8 change, identifying areas where the species is most likely to persist and where it is most threatened, as well as

9 potential new habitats (e.g., Higa et al., 2012). Restoration of ecological habitats within and between protected areas

10 may help facilitate the movement of species across climatic gradients in response to climate change

11 (Klorvuttimontara et al., 2011; Hughes et al., 2012). Key seed dispersal agents may need to be protected because of

12 their potential role in long-distance plant movements in fragmented landscapes (Corlett, 2009). Assisted migration

13 (or 'managed translocation') of genotypes and species is an increasingly common suggestion for plants and animals

14 where adjustments to climate change are constrained by natural rates of movement, although the risks and benefits

15 in each case need to be considered carefully (e.g. Liu et al., 2010; Olden et al., 2010; Tchebakova et al., 2011; Ogawa-Onishi et al., 2011; Ishizuka & Goto, 2012). Ex situ conservation can provide back-up for populations and

16 17 species that are most at risk from climate change (Chen et al., 2009).

18

19 There is a lack of scientifically well-founded recommendations and programs aimed at development of adaptation 20 plans for the forest-tundra ecotone in North Asia at a state level (Anisimov et al., 2010). Comprehensive monitoring, 21 assessments and projections that can anticipate numerous development scenarios are needed to elaborate a plan for 22 adaptation to the cumulative effects of resource development, climate change, and demographic changes that are 23 occurring (Walker et al., 2011). Similar problems are widespread in other parts of Asia, although awareness of the 24 need for adaptation plans is increasing.

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24.4.3. Coastal Systems and Low-Lying Areas

29 24.4.3.1. Sub-Regional Diversity

31 Asia's long coastline includes the full global range of muddy, sandy, and rocky shore types, as well as extensive 32 estuarine systems. Asia's tropical and subtropical coasts support an estimated 45% of the world's total mangrove 33 forest and include the most mangrove-rich country (Indonesia) and the largest single tract of mangrove forest (the 34 Sundarbans of Bangladesh and India) (Giri et al., 2011). Low-lying areas near the coast of equatorial SE Asia 35 support most of world's peat swamp forests (Posa et al., 2011), which are a massive store of carbon, as well as 36 extensive areas of other forested swamp types. Intertidal salt marshes are widespread along temperate and arctic 37 coasts, while a variety of non-forested wetlands occur inland, including freshwater marshes and peat bogs. Asia also supports around 40% of the world's coral reef area (Spalding et al., 2001; Burke et al., 2011), mostly in SE Asia, 38 39 with the most extensive reefs and the world's most diverse reef communities in the 'coral triangle' (in Indonesia, 40 Malaysia, the Philippines, and Papua New Guinea; see also Chapter 30, this volume, Box 30-3). Seagrass beds are 41 also widespread, although less well studied, and Asia supports the majority of the world's seagrass species (Green 42 and Short, 2003). Six of the seven living species of sea turtle are found in the region and five species nest on Asian 43 beaches (Spotila, 2004). Kelp forests and other seaweed beds are important on temperate coasts (Bolton, 2010; 44 Nagai et al., 2011). Permafrost and sea-ice influence coastal processes in the far north (Are et al., 2008). The sea-ice 45 itself supports a specialized community of mammals, including the polar bear, walrus, several species of seals, and 46 the beluga and bowhead whales, as well as birds, fish and other species (Forbes, 2011; Chapter 28, Sections 47 28.2.3.3. and 28.2.3.4.).

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49 50 24.4.3.2. Observed Impacts

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52 Most of Asia's non-Arctic coastal ecosystems are under such severe pressure from non-climate human impacts, that

53 climate impacts are hard to detect. For example, observations of impacts from rising sea levels in Asia have

54 reflected coastal subsidence rather than the impact of climate change, since most of the major deltas in Asia are now 1 sinking (as a result of groundwater withdrawal, floodplain engineering, and trapping of sediments by upstream

2 dams) at rates many times faster than the global sea-level is rising (Syvitski et al., 2009). Widespread impacts can be

3 attributed with *high confidence* to climate change, however, for coral reefs, where the temporal and spatial patterns

4 of large-scale bleaching events generally correlate well with higher than normal sea surface temperatures (Hoegh-

5 Guldberg, 2011; Krishnan et al., 2011). Increases in coastal water temperatures are also one of the most plausible

6 explanations for widespread declines in beds of large seaweeds in temperate Japan: the Isoyake phenomenon (Nagai 7 et al., 2011). Longer periods of annual herbivore activity are one suggested mechanism. Warming coastal waters

- 8 have also been implicated in the northwards expansion in Japanese waters of tropical and subtropical macroalgae
- 9 and toxic phytoplankton (Nagai et al., 2011), fish (Tian et al., 2012), and tropical corals, including key reef-forming
- 10 species (Yamano et al., 2011), over recent decades.
- 11

12 The impact of warming is also evident on sparsely populated Arctic coastlines, where erosion appears to be

13 accelerating. Permafrost and sea ice are additional factors for coastal erosion in Arctic Asia and the overall influence

14 of cryogenic processes increases coastal retreat, in spite of the fact that most of the year coasts are protected by

15 continuous ice cover (Are et al., 2008; Razumov, 2010). Average erosion rates of Asian Arctic coastlines range from

16 0.27 m/year (Chukchi Sea) to 0.87 m/year (East Siberian Sea). A number of segments in the Laptev Sea and in the

17 East Siberian Sea are characterized by rates greater than 3 m/year (Lantuit et al., 2012). The decline in the extent of

18 arctic sea-ice documented in AR4 has continued, but the impacts on ice-dependent species and ecosystems in Arctic

- 19 Asia are so far unclear (WGI, Ch. 4, ZOD; WGII, Ch. 28, ZOD).
- 20 21

22 24.4.3.3. Projected Impacts

23 24 It is *likely* that there will be an overall increase in marine biodiversity at temperate latitudes as temperature 25 constraints on the distributions of warm-water taxa are relaxed, but biodiversity in tropical regions may fall if, as

26 some evidence suggests, tropical marine species are already near their thermal maxima (Cheung et al., 2009, 2010; 27 Neuheimer et al., 2011). A combination of projected shifts in species distributions and expected changes in total

28 primary production may lead to a regional redistribution of fisheries potential, with large declines in the tropics and

29 large increases in high-latitude regions (Cheung et al., 2010). Overall, however, the connectivity of marine habitats

30 and the relatively high dispersal abilities of many marine organisms are expected to keep the extinction rate below

31 that expected for terrestrial habitats (Cheung et al., 2009). Projected impacts are greatest for coral reefs, where a

32 continuation of current trends in sea-surface temperatures and ocean acidification suggests that existing coral-

33 dominated reefs will largely disappear by mid-century (Vivekanandan et al., 2009; Hoegh-Guldberg, 2011; Burke et

34 al., 2011), although the capacity of coral communities to adjust by changes in species composition, or by the

35 acclimation and/or adaptation of coral species, is not well understood (Ateweberhan and McClanahan, 2010;

36 Fabricius et al., 2011; Guest et al., 2012; Howells et al., 2012). The impacts of ocean acidification on other organisms are also poorly understood (Hendriks et al., 2010). Warm-temperate kelp beds may be more vulnerable to

37 catastrophic phase shifts with rising temperatures (Ling et al., 2009; Graham, 2010).

38

39

40 The uncertainties in future sea-level rises are still large (WG1, Ch. 13, AR5 ZOD). The major projected impacts 41 include coastal flooding, increased erosion, and saltwater intrusion into surface and groundwater. In the absence of

42 other impacts, coral reefs could grow fast enough to keep up with rising sea-levels, but mangroves, salt marshes, and

43 seagrass beds will decline unless they can move landwards or they receive sufficient sediment to keep pace, and

44 beaches may erode (Gilman et al., 2008; Bezuijen, 2011; Forbes, 2011). Loucks et al. (2010) predict a 96% decline

45 in tiger habitat in Bangladesh's Sunderbans mangroves with a 28 cm sea-level rise if sedimentation does not

46 increase surface elevations. Coastal freshwater swamps and marshes will be vulnerable to saltwater intrusion with

47 rising sea-levels. However, in most river deltas, the global sea-level rise will continue to be outpaced by local

- 48 subsidence for non-climatic reasons (Syvitski et al., 2009).
- 49

50 Cyclones affect most of the Asian coastline, except in the far north, west, and 10° either side of the equator. Natural

51 coastlines are resilient, but large cyclones can have a devastating impact on isolated ecosystem fragments. However,

52 current trends in cyclone frequency and intensity are unclear (Seneviratne et al., 2012). A combination of cyclone

- 53 intensification and sea-level rise could potentially result in a large increase in coastal flooding (Knutson et al., 2010).
- 54 Cyclones can also have a large impact on the productivity of coastal waters through increased nutrient run-off and

water circulation (Qiu *et al.*, 2010). In addition to any changes in cyclone activity, sea turtles nesting beaches may be impacted by increased temperature and sea-level rise, but the capacity of turtle populations to adapt is not well

- understood (Hawkes *et al.*, 2009; Poloczanska *et al.*, 2009; Fuentes *et al.*, 2011).
- 4

5 In the Asian Arctic it is *likely* that rates of coastal erosion will increase as a result of interactions between rising sea-6 levels and projected changes in permafrost and the length of the ice-free season (Pavlidis et al., 2007; Lantuit et al., 7 2012). The most sensitive region to potential increases in permafrost and sea surface temperatures on the Asian 8 Arctic coast is the Kara Sea region (Lantuit et al., 2012). Sea level rise may have different influences on coastal 9 processes depending on the sediment budget equilibrium, playing a minor role if there is a strong imbalance in the 10 sediment budget, but appearing to be the main factor if the sediment budget is balanced (Leont'yev, 2008). The most 11 prominent changes in the dynamics and morphology of the coastal zone are expected where the coasts are composed 12 of loose permafrost rocks and are therefore subject to intensive thermal abrasion. Assuming that sea level will rise 13 by 0.5 m over the next century, modeling studies predict that the rate of recession due to thermal erosion will 14 increase 1.5- to 2.6-fold for the coasts of Laptev Sea, East Siberian sea and of West Yamal in the Kara Sea 15 compared to the rate observed in first years of the XXI century. This rate will vary across the Asian Arctic coast 16 from 3 to 9 m/year (Pavlidis et al., 2007).

17

18 It has been suggested that the warming and acidification associated with an atmospheric CO_2 concentration of 450 19 ppm will lead to the loss of coral-dominated reef systems (Hoegh-Guldberg, 2011). Investigations of coral reefs 20 around natural volcanic seeps of CO_2 in Papua New Guinea suggest a much higher threshold (750 ppm) for

persistence of coral cover at current water temperatures, but with severe losses in biodiversity and structural
 complexity (Fabricius *et al.*, 2011).

22 23 24

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24.4.3.4. Vulnerabilities to Key Drivers

Offshore marine systems appear to be most vulnerable to rising water temperatures, plus the impacts of ocean acidification, particularly for calcifying organisms such as corals. Sea-level rise will be the key issue for many coastal areas, particularly if it is combined with changes in cyclone frequency or intensity, or in Arctic Asia, with a lengthening open-water season. Polar bears, walruses, ice-associated seals, and beluga and bowhead whales may be threatened by the expected continuing decline in the extent of sea-ice in the arctic (Forbes, 2011; Kovacs *et al.*, 2011).

33 34

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35 24.4.3.5. Adaptation Options

37 The connectivity of marine habitats and the relatively high dispersal abilities of many marine and coastal organisms 38 should maximize the capacity for autonomous (spontaneous) adaptation in natural and semi-natural coastal systems 39 (e.g., Cheung et al., 2009). 'Hard' coastal defenses, such as dykes, levees and sea walls, may protect settlements, but 40 at the cost of preventing adjustments by mangroves, salt marshes and seagrass beds to rising sea-levels. The 41 acquisition of landward buffer zones that provide an opportunity for future inland migration could mitigate this 42 problem (Erwin, 2009), but is rarely practical. Large sections of Asia's coastline are already highly degraded and 43 there are many opportunities for restoration of coastal systems (Crooks et al., 2011). The high carbon sequestration 44 potential of the organic-rich soils in mangroves and peatswamp forests provides opportunities for combining 45 adaptation with mitigation.

46 47

48 24.4.4. Food Production Systems and Food Security 49

50 24.4.4.1. Sub-Regional Diversity

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52 AR4 Section 10.4.1.1 pointed out that there will be regional differences in the impacts of climate change on food

53 production. Research since then has validated this generalization and new data are available especially for West and

54 Central Asia (see Table 24-7). These differences will be apparent in the discussion below. In addition, there are now

1 more detailed researches on the impacts to crop production. In AR4 Section 10.4.1, climate change was projected to

- mainly lead to reduction in yield. New research shows there will be gains as well. Depending on the regions and the
 crops grown, effects will vary substantially.
 - [INSERT TABLE 24-7 HERE

6 Table 24-7: Summary of observed and projected impacts in the food sector.]

24.4.4.2. Observed Impacts

While there is consensus that climate change will affect food production systems and food security, the precise nature and timing of these impacts, as well as their implications for human livelihoods are still uncertain (Hertel *et al.*, 2010). There are limited data in Asia on observed impacts of climate change on food production systems. In Jordan, it was reported that in 1999, the total production and average yield for wheat and barley were the lowest among the years 1996 to 2006. This could be explained by the low rainfall during that year, which was 30% of the average. These results suggest that both crops are vulnerable to climatic variations (Al-Bakri *et al.*, 2010).

In China, rice yield responses to recent climate change at experimental stations, was assessed for the period of
1981–2005 (Zhang *et al.* 2010).. The study concluded that there is a variable climate to yield relationships at a
regional scale. In some places, yields were positively correlated with temperature when they were also positively
related with radiation. However, in other places, lower yield with higher temperature was accompanied by positive
correlation between yield and rainfall.

One possible approach to generating new knowledge on observed impacts of climate change is to combine local knowledge with scientific assessments. For example, the nomadic herders of Mongolia demonstrated a detailed understanding of weather and climate including an account of climatic change that integrates multiple indicators (Marin, 2010. However, their evidence of change is only partly supported (or even contradicted) by meteorological records, larger scale predictions and general circulation models.

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24.4.4.3. Projected Impacts 32

Production. AR4 Section 10.4.1.1 mainly dealt with cereal crops (rice, wheat corn). Since then, impacts of climate change have been modeled for additional cereal crops and sub-regions. In semi-arid and arid regions of Western Asia, a review paper has shown that rainfed agriculture is sensitive to climate change both positively and negatively. A rise in CO_2 concentration may benefit the semi-arid crops by increasing the crop water use efficiency and net photosynthesis leading to greater biomass, yield and harvest index (Ratnakumar *et al.*, 2011). C_3 plants responded with a higher average increment in biomass production than C_4 plants. For example, wheat yield increased by 10-

39 20% with elevated CO_2 (350ppm to 700ppm). It was hypothesized that elevated CO_2 would produce more biomass 40 and seed yield through an increased water use efficiency. In Yarmouk basin, Jordan, simulation with DSSAT

40 and seed yield through an increased water use efficiency. In Yarmouk basin, Jordan, simulation with DSSAT 41 showed that wheat and barley yields will decline by 10-20% and 4-8% respectively with 10-20% reduction in

rainfall (Al-Bakri *et al.*, 2010). Increase in rainfall by 10–20% increased the expected yield by 3–5% for barley and

43 9–18% for wheat, respectively. However increase of air temperature had mixed results. Increasing temperature by 1,

2, 3 and 4°C resulted in deviation from expected yield by -14%, -28%, -38% and -46% for barley and -17%, +4%,
+43% and +113% for wheat, respectively. These results indicate that barley would be more negatively affected by

46 climate change and therefore adaptation plans should prioritize the arid areas cultivated with this crop.

47

48 In Swat and Chitral districts of Pakistan, (mountainous areas with average altitudes of 960 and 1500 m above sea

49 level, respectively), there were mixed results as well (Hussain and Mudasser, 2007). Projected temperature increase 50 of 1.5 and 3 °C would lead to wheat yields decline (by 7% and 24% respectively) in Swat district but would lead to

- an increase (by 14% and 23% respectively) in Chitral district. If precipitation increases by 5-15% during the
- growing season, the study showed a negligible impact on wheat yield.

53

1 In India, climate change impacts on sorghum were analyzed using Info Crop-SORGHUM simulation model

2 (Srivastava *et al.*, 2010). A changing climate was projected to reduce monsoon sorghum grain yield by 2 to 14% by

3 2020 with worsening yields by 2050 and 2080. In addition, climate change was projected to reduce winter crop

- 4 yields up to 7% by 2020, up to 11% by 2050 and up to 32% by 2080. In the Indo-Gangetic Plains (IGPs), a similar
- reduction in wheat yields is projected, unless appropriate cultivars and crop management practices were adopted by
 South Asian farmers (Ortiz *et al.*, 2008).
- 6 7

8 In China, a number of studies on the impacts of climate change to crop productivity had mixed results. Rice is the

- 9 most important staple food in Asia. Studies show that climate change will alter productivity in China but not always
- negatively. With rising temperatures, the process of rice development accelerates and reduces the duration for
- growth. In one study using IPCC SRES B2 without CO_2 fertilization effect, the yield of irrigated rice along the Yangtze River decreased by 14.8%, and the yield of rain-fed rice decreased by 15.2% on average (Shen *et al.*, 2011).
- With CO₂ fertilization effect factored in, the yield of irrigated rice decreased by 3.3% and the yield of rain-fed rice
- decreased by 4.1% on average. Tao *et al.* (2008) reported similar findings using all 20 combinations of four SRESs
- 15 (A1F1, A2, B2, B1) and five GCMs (HadCM3, PCM, CGCM2, CSIRO2, ECHAM4). Without CO₂ fertilization
- 16 effects, the growing period would be shorter and yield would decrease. The median values of yield decrease ranged
- 17 from 6.1% to 18.6%, 13.5% to 31.9%, and 23.6% to 40.2% for air temperature increases of 1, 2, and 3 °C,
- 18 respectively. However, if CO₂-fertilization effects were included, the median values of yield changes ranged from -
- 19 10.1% to 3.3%, -16.1% to 2.5%, and -19.3% to 0.18% for air temperature increases of 1, 2, and 3 °C, respectively,
- 20 across the stations. Other studies show similar results that higher temperature would seriously lower rice yields due
- 21 to shorter crop duration (Xiong *et al.*, 2010; Yao *et al.*, 2007).
- 22

In contrast, Zhang *et al.* (2010) reported that rice yield responses to temperature were broadly positive, which means that yields were not limited by an increase in T_{min} , T_{max} , or T_{mean} . The authors hypothesize that radiation level is the major climatic driver for yield fluctuations at these Chinese experimental stations, and the positive yield correlation to temperature can be explained by the correlations between radiation and temperature, which were positive at most studied stations. Thus, the positive effect of radiation on yield overwhelmed temperature's negative effect on rice yield.

29

Wassman *et al.* (2009a, 2009b) provide the most comprehensive review of climate change impacts and adaptation

for rice in the region. A key conclusion of the report is that in terms of risks of increasing heat stress, there are parts

- 32 of Asia where current temperatures are already approaching critical levels during the susceptible stages of the rice
- 33 plant. These include: Pakistan/north India (October), south India (April, August), east India/Bangladesh (March-
- June), Myanmar/Thailand/Laos/Cambodia (March-June), Vietnam (April/August), Philippines (April/June),
 Indonesia (August) and China (July/August).
- 36

There was also simulation research for other crops in China. In the Huang-Hai Plain, China's most productive wheat growing region, modeling work indicated that winter wheat yields would increase on average by 0.2 Mg ha-1 in 2015-2045 and by 0.8 Mg ha-1 in 20700-2099 due to warmer nighttime temperatures and higher precipitation, under A2 and B2 scenarios using HadCM3 model (Thomson *et al.*, 2006). Yields were positively influenced by increasing precipitation projected under the climate change scenarios, with the highest average yields in the 2085 time period when the precipitation increase was greatest.

43

Liu et al. (2010) worked on a wheat-maize cropping system in Huang-Huai-Hai (3H) Plain, China. Generally,

- 45 climate change (2 and 5 °C increase in temperature; precipitation increasing and deceasing by 15 and 30%;
- 46 atmospheric CO2 enrichment to 500 and 700 ppmv) would result in a mean relative yield change (RYC in %) of
- 47 –10.33% with standard deviation of 20.27%, and the lowest and highest RYC values of –46% and 49%, respectively.
- 48 However with CO₂ fertilization a positive change in RYC was obtained. In addition, increasing precipitation
- 49 mitigates the negative impact of increasing temperatures on yield. On average, without CO_2 enrichment, the mean of
- 50 RYC for irrigated land is less negative $(-18.5\pm12.6\%)$ than that for rain-fed land $(-21.5\pm14.2\%)$. With CO2
- 51 enrichment there was no significant differences between irrigated and rainfed yield. These results show that CO_2
- 52 enrichment blurs the role of irrigation.
- 53

- 1 The potential climate change impacts on the productivity of five major crops (canola, corn, potato, rice, and winter
- 2 wheat) in eastern China have also been investigated using RegCM3 regional climate model under A2 scenario

3 (Chavas *et al.*, 2009). Their results indicate that aggregate potential productivity (i.e. if the crop is grown

4 everywhere) with CO_2 fertilization increased 6.5% for rice, 8.3% for canola, 18.6% for corn, 22.9% for potato, and

- 5 24.9% for winter wheat, although with significant spatial variability for each crop. However, without the enhanced 6 CO₂- fertilization effect, potential productivity declined in all cases ranging from 2.5 to 12%.
- 6 CO₂- fertilization effect, potential productivity declined in all cases ranging from 2.5 to 12%.
- 8 Extreme weather events are expected to further negatively affect agricultural crop production (IPCC, 2012; Handner 9 et al. 2012). For example, extreme temperatures could lower yields of rice (Tian et al., 2010; Mohammed and
- 10 Tarpley, 2009). With higher precipitation, flooding could also lead to lower crop production (SREX chap 4). For
- example, cyclone Sidr which hit Bangladesh in 2007 caused more than 3,000 deaths and the damage to agriculture
 was estimated to be in excess of US\$ 3 billion (Hasegawa, 2008). Another example is from the Philippines which
- lies in the typhoon belt with an average of 20 tropical cyclones per year in addition to other extreme weather events it experiences (Y<u>umul</u> et al., 2011; Yu<u>mul et al.</u>, 2010), One study showed that relative loss per crop as part of the annual farm household income due to one tropical cyclone event for yellow corn, banana, and rice were 64%, 24%, and 27%, respectively (Huigen and Jens, 2006).
- 17
- Farming systems and crop areas. Since AR4 (Section 10.4.1.2), more information is available on the impacts of climate change on farming systems and cropping areas in more countries in Asia and especially in Central Asia. In general, recent studies validate the northward shifts of crop production with current crop lands under threat from the impacts of climate change as mentioned in AR4.
- 22

Climate change threatens the food security of West Asia where most of drylands are comprised of rangelands
 (Thomas, 2008). The region has the world's lowest rate of renewable water resources per capita and is already the
 major grain importing region of the world. Climate change will exacerbate existing threats to food production and
 security such as high population growth rates, water scarcity, and land degradation.

27

28 In Central Asia, changes in temperature and precipitation regimes could to lead to: changes in area suitable for 29 growing rain-fed production of cereals and other food crops, changing sustainable stocking rates, and modifying crop irrigation requirements (Lioubimtseva and Henebry, 2009). The region is expected to become warmer during 30 31 the coming decades and increasingly arid across the entire region, especially in the western parts of Turkmenistan, 32 Uzbekistan, and Kazakhstan. The impacts to food production will vary by country. Some parts of the region could 33 be gainers (cereal production in northern and eastern Kazakhstan could benefit from the longer growing season, 34 warmer winters and slight increase in winter precipitation), while others could be losers (particularly western 35 Turkmenistan and Uzbekistan, where frequent droughts could negatively affect cotton production, increase already 36 extremely high water demands for irrigation, and exacerbate the already existing water crisis and human-induced 37 desertification). In addition Central Asia and the Caucasus is the second most vulnerable region of the world to crop 38 loss by pollinator loss (Christmann and Aw-Hassanb, 2011). Agricultural production in general depends on honey 39 bees (Apis mellifera), but honey bees are highly sensitive to change of temperatures and can provide service only on 40 sunny, warm, dry and not too windy days. The tolerance of local honey bees to climate change needs further

- 41 elucidation.
- 42

43 In India, the Indo-Gangetic Plains (IGPs) are under threat of significant reduction in wheat yields (Ortiz *et al.*, 2008).

This area produces 90 million tons of wheat grain annually (about 14–15% of global production). Climate

45 projections based on a doubling of CO_2 using a CCM3 model downscaled to a 30 arc-second resolution as part of the

- Worldclim data set showed that there will be a 51% decrease in the most favorable and high yielding area due to heat stress. About 200 million people (using current population), who's food intake relies on crop harvests would
- 48 experience adverse impacts.
- 49

50 In Sri Lanka, a number of studies reviewed by Eriyagama *et al.* (2010) showed varying results. Tea cultivation at

- 51 low and mid-elevations are more vulnerable to the adverse impacts of climate change than those at high elevations.
- 52 Projected coconut production after 2040 in all climate scenarios will not be sufficient to meet local consumption.
- 53 The total impact on agriculture (rice, tea, rubber and coconut) production ranges from a decrease of US\$96.4 million
- 54 (-20%) to an increase of US\$342 million (+72%) depending on the climate scenarios.

1

In eastern China, a study showed corn and winter wheat production would benefit significantly from climate change
in the North China Plain (Chavas *et al.*, 2009). Rice would remain dominant in the southeast but emerges in the
northeast, potato and corn yields would become viable in the northwest, and potato yields suffer in the southwest.
The study defined vulnerable and emergent regions under future climate conditions as those having a greater than

6 10% decrease or increase in productivity, respectively.

7

8 Rice growing areas are also expected to shift with climate change throughout the region. In Japan, increasing water 9 temperature (1.6–2.0 °C) could lead to a northward shift of the isochrones of safe transplanting dates for rice 10 seedlings (Ohta and Kimura, 2007). As a result, rice cultivation period will be prolonged by approximately 25-30 11 days. This will allow greater flexibility of variation in the cropping season as compared with that at present; thus, 12 resulting in a reduction in the frequency of cool summer damage in the northern districts. In Indonesia, a marked 13 increase in the probability of a 30-day delay in monsoon onset in 2050 is projected, as a result of changes in the 14 mean climate, from 9–18% today (depending on the region) to 30–40% at the upper tail of the distribution (Naylor 15 et al. 2007). In addition, there would be an increase in precipitation later in the crop year (April–June) of $\approx 10\%$ but 16 a substantial decrease (up to 75% at the tail) in precipitation later in the dry season (July–September). However, the 17 increase in April-June rainfall would not compensate for reduced rainfall later in the crop year, particularly if water 18 storage for agriculture was inadequate. Second, the extraordinarily dry conditions in JAS could preclude the planting 19 of rice and all other crops without irrigation during these months by 2050. In Sri Lanka, studies on rice production 20 have mixed results (Eriyagama et al., 2010). An earlier study showed that a 0.1-0.5°C increase in temperature could 21 depress rice yield by approximately 1-5%. However, another experiment suggested that rice yields respond 22 positively (increases of 24 and 39% in the two seasons) to elevated CO₂ even at higher growing temperatures 23 (>30°C) in subhumid tropical environments. The real threat to rice cultivation might be changes in the amount of 24 precipitation and temporal distribution. Climate change is expected to affect water supply for rice cultivation in Sri 25 Lanka (De Silva et al., 2007). During the wet season, irrigated rice production is projected to be positive in the 26 extreme south of the country, confirming results of a previous study. However, the impacts are negative across most 27 of Sri Lanka. During the wet season, average rainfall would decline by 17% (A2) and 9% (B2), with rains ending

earlier. Consequently, the average paddy irrigation water requirement would increase by 23% (A2) and 13% (B2).

30 Similarly in China, Xiong et al. (2010) reported there would be insufficient water for agriculture in the 2020s and 31 2040s, due to increases in water demand for non-agricultural uses, using HadAM3H GCM and PRECIS regional 32 model, especially under the A2 scenario. The proportion of water demanded by rice (which consumes 79% of total 33 baseline potential water demand of three grain crops) is projected to increase, because of significant increases in the 34 projected water demand by rice under A2 (+62% for the 2020s above the baseline, and +58% for the 2040s), and 35 moderate increases under B2 (5% and 2% for the 2020s, and the 2040s, respectively). However, due to increases in 36 demand in other sectors (domestic, environmental and industrial) captured in the socio-economic scenarios (SES), 37 the water available for agriculture decreases dramatically under A2 by 5% (2020s) and 21% (2040s), and by 3% and 38 16%, respectively under B2.

39

40 Livestock, fishery, aquaculture. Since AR4, very limited information has been added on the impacts of climate 41 change on livestock, fishery, and aquaculture. In Mongolia, Marin (2010) showed that both local knowledge of 42 herders and meteorological data and projections are important in assessing the impacts of climate change as well as 43 potential adaption strategies. While regional models and local analyses agree that Mongolia has become warmer, 44 predictions either ignore or are contradictory about the changes in precipitations and sand storms. The nomadic 45 herders of Mongolia demonstrated a detailed understanding of weather and climate. According to the herders, the 46 dust storms and droughts were more frequent and severe, rains were patchier, less effective ('harder') and delayed. 47 All of these could affect livestock production in the country.

48

49 Future food supply and demand. AR4 Section 10.4.1.4 was largely based on global models including Asia. Since 50 then there are now a few quantitative studies on the whole continent and countries. In general, these studies show 51 that the risk of hunger, food insecurity and loss of livelihood due to climate change will be high as discussed below. 52

Rice is a key staple crop in Asia and 90% or more of the world's production is from Asia. An Asia-wide study revealed that climate change scenarios (using 18 GCMs for A1B; 14 GCMs for A2 and 17 GCMs for B1would 1 reduce rice yield over a large portion of the continent (Masutomi *et al.*, 2009). The most vulnerable regions were

2 western Japan, eastern China, the southern part of the Indochina peninsula, and the northern part of South Asia. In

these areas, rise in temperature during the growing periods would be the main cause of the decreases in yield. The

4 negative impacts of climate change were diminished but not totally eliminated by the positive effect of CO_2

fertilization. In a global study, Hertel *et al.* (2010) showed that under the low-productivity scenario (due to climate
 change), prices for major staples would rise 10–60% by 2030 in Asia. Poverty rates in some non-agricultural

household could rise by 20–50% in parts of Asia and fall by significant proportions for agriculture households.

8

In Russia, climate change may also lead to "food production shortfall" which was defined as an event in which the annual potential (i.e. climate-related) production of the most important crops in an administrative region in a specific year falls below 50% of its climate-normal (1961–1990) average (Alcamo *et al.*, 2007). The frequency of shortfalls in the main crop growing regions in the same year is around 2 years/decade under climate baseline conditions but could climb to 5–6 years/decade in the 2070s using the ECHAM and HadCM3 models and the A2 and B2 SRES. The increasing shortfalls was attributed to severe droughts. The study estimated that the number of people living in these regions may grow to 82–139 million in the 2070s. Increasing frequency of extreme climate events will pose an

- 16 increasing threat to the security of Russia's food system.
- 17

18 Likewise, most of the studies reviewed in the previous sections show negative impacts of climate change to crop

19 yield and therefore presumably on food supply. In contrast, climate change may also lead to increase food supply of

20 some countries. For example, climate change may provide a windfall for wheat farmers in parts of Pakistan.

21 Warming temperatures would make it possible to grow at least two crops (wheat and maize)/year in the mountain

areas (Hussain and Mudasser, 2007). It will also allow more time for land preparation of the subsequent maize crop,
 with beneficial effects on yield. The increased productivity of the wheat–maize cropping system is expected to

improve food security, increase farm income and reduce overall poverty of the farm households in the area.

25

26 Pests and diseases. AR4 contained a generalization about the possibility of increasing pests and diseases due to 27 climate change. Since then, there have been very few studies on climate change and pests and diseases which 28 support the aforementioned conclusion. For example in South Asia, warming temperatures could lead to higher 29 incidence of spot blotch (caused by Cochliobolus sativus), already a serious constraint to wheat production at 30 present. An increasing mean minimum temperature in March showed a positive relationship with spot blotch 31 severity (Sharma et al., 2007). In the future, Sharma et al. (2010) recommended the need to regularly monitor pest 32 populations to determine if a threshold has been exceeded and if control measures are required. This information 33 will also be valuable for forecasting pest populations, severity of damage, and pest outbreaks. Climate change may 34 also modify the effectiveness of biological control (e.g. natural enemies), biopesticides, and synthetic insecticides. 35

36

37 24.4.4.4. Vulnerabilities to Key Drivers

38 39 Vulnerability of rainfed agriculture is expected to increase with decreasing precipitation. However, decreasing 40 availability of water due to economic and population growth will negatively influence the irrigated agriculture as 41 well. Rapid population growth will raise food demand, and further industrialization of developing countries could 42 lead to massive migration from rural areas into urban ones. One cannot ignore the impact of governmental decision, 43 such as land policies, or improvements in agricultural technologies, and market oriented land-management, which 44 can affect the efficiency and scale of cultivated land. Due to this plurality of factors in determining vulnerability of 45 the food production systems it is becoming more and more difficult to ascertain a clear picture of future climate 46 change impacts.

- 47 48
- 49 24.4.4.5. Adaptation Options50

51 Since AR4, there have been additional studies on recommended and potential adaptation strategies and practices in 52 Asia (see Table 24-8). There is new information on West and Central Asia. There are also much more crop specific

and country specific adaptation options available.

1 **[INSERT TABLE 24-8 HERE**

- 2 Table 24-8: Summary of adaptation options for agriculture in Asia.]
- 3

4 It is noteworthy that farmers have been adapting to climate risks for generations. Indigenous and local adaptation

5 strategies have been documented Southeast Asia (Peras et al., 2008; Lasco et al., 2011; Lasco et al., 2010). These 6 strategies could be used as a basis for future climate change adaption. In addition, social and institutional aspects of

7 climate change adaptation have also been investigated in the Philippines. Agent-based modeling showed that small

8 holder farmers face a number of constraints in adapting new technologies to cope climate risks (Acosta-Michlik and 9 Espaldon, 2008). In general, lack of knowledge and money were the most important reasons for not adopting

10 drought-related technical measures. It is interesting that in the above studies there are many non-farm related

11 adaptation strategies. Local government units (LGUs) can also play a catalytic role in climate change adaptation as

- 12 shown by the experience of Albay province in the Philippines (Lasco et al. 2008).
- 13 14

15

16 17

18

24.4.5. Human Settlements, Industry, and Infrastructure

24.4.5.1. Sub-Regional Diversity

19 Asia, being the largest continent of the world in terms of area and population, is both diverse and complex.

20 Sustainable development will be challenged as climate change compounds the pressures that rapid urbanization,

21 industrialization and economic development have placed on natural resources (IPCC, 2007b). Settlement patterns,

22 urbanization and changes in socioeconomic conditions greatly influence trends in exposure and vulnerability to 23 climate extremes (IPCC, 2012).

24

25 Population distribution is uneven within Asia. For example, two sub-regions i.e. East Asia and South-Central Asia, 26 account for 80% of the continents population (UNFPA, 2010). At present 69% of the world's rural population is 27 highly concentrated in a few Asian countries. India and China has the largest rural population amounting to 45% of 28 the world's rural population, followed by Bangladesh, Indonesia and Pakistan each with over 107 million rural

29 inhabitants. Much of the increase projected in the world population is expected to come from 39 high-fertility

30 countries of which nine are located in Asia. Notwithstanding this, population growth rates have been decreasing in 31 almost all sub regions of Asia since 2000 (UN ESCAP, 2011).

32

33 Presently, about one in every five urban dwellers in Asia lives in large urban agglomerations and little less than 50% 34 of urban dwellers live in small cities (UN, 2012). However, there is wide sub-regional level variation. For example,

35 North and Central Asia are the most urbanized areas where over 63% of the population live in urban areas with the

36 exception of Kyrgyzstan and Tajikistan, followed by East and North-East Asia where rapid urbanization in last two

37 decades led to 50 % population living in cities by 2010 (UN ESCAP, 2011; UN Habitat, 2010). South and South-

West Asia are the least urbanized sub-regions with only 33% of the population living in urban areas. However, the 38

39 sub-region has the highest urban population growth rate within Asia at an average of 2.4% per year during 2005-

40 2010 (UN-ESCAP, 2011). By the middle of this century, Asia's urban population will increase by 1.4 billion, and

41 that alone will account for over 50% of the global population, with China and India projected to account for about a

- 42 third of the increase in the coming decades (UN, 2012).
- 43

44 Most Asian countries are witnessing significant development opportunities as well as a myriad of challenges. In 45 2010, seven Asian economies (China, India, Indonesia, Japan, Korea, Malaysia and Thailand) shared 78% of Asia's 46 population and 87 % of Asia's GDP (ADB, 2011). Across all the sub-regions of Asia, poor people and urban slum 47 dwellers tend to live in high-risk areas such as unstable slopes and flood plains, and often cannot afford well-built 48 houses. The poorest people are expected to suffer the most from climate change.

49 50

52

51 24.4.5.2. Observed Impacts

53 Asia experienced the highest number of weather- and climate-related disasters during the period of 2000 to 2008 and 54 suffered huge economic loss, accounting for the second highest proportion (27.5%) of total global economic losses.

1 Loss of human lives, cultural heritage, and ecosystem services, are difficult to value and monetize, and thus are

- 2 poorly reflected in estimates of losses. Impacts on the informal or undocumented economy, as well as indirect
- effects, can be very important in some areas and sectors, but are generally not counted in reported estimates of losses
 (IPCC, 2012).
- 5

Flood mortality risk is heavily concentrated in Asia [see Figure 24-2]. The top ten countries at risk of floods (based
on number of lives lost) are India, Bangladesh, China, Viet Nam, Cambodia, Myanmar, Sudan, Korea, Afghanistan
and Pakistan (UNISDR, 2009). Severe floods of July 26, 2005 in Mumbai, which happened after receiving 944 mm
rainfall within 24 hours is attributed to both climate as well as non-climate factors such as lack of early warning,
preparedness and response capacities at the local level, lack of modern rain gauges, poor urban drainage systems,
blockages in the natural drainage channels, poor waste management, poor urban planning, lack of civic sense among
citizens, among others (IPCC, 2012; Surjan, et al., 2010).

12 13

14 [INSERT FIGURE 24-2 HERE

- 15 Figure 24-2: Hazard mortality risk.]
- 16

On the contrary, in many parts of Asia, there exist seasonal shortfalls in the availability of water, which is also a growing crisis (ADB, et al., 2012). Despite the increasing number of people living in floodplains, strengthening of capacities to address the mortality risk associated with major weather-related hazards (e.g. floods), mortality risk relative to population size is showing downward trend, such as in East Asia and the Pacific, where mortality risk is now at a third of its 1980 level (UNISDR, 2011).

22 23

25

24 24.4.5.3. Projected Impacts

About half to two-third of Asian cities with 1 million or more population are located in regions exposed to natural hazards (UN, 2012). The possibility is high for underestimating the impact of rare or more severe natural disasters on urban areas. Asian mega-deltas are susceptible to extreme impacts due to a combination of high-hazard rivers, coastal flooding and increased population exposure from expanding urban areas with large proportions of high vulnerability groups (IPCC, 2012).

31

32 *Floodplains*. Three of the world's five most populated cities in 2011 are located in areas with high risk of floods. 33 They are Tokyo, Delhi and Shanghai (UN, 2012). Flood risk and associated human and material losses are heavily 34 concentrated in India, Bangladesh, and China. East Asian region in particular experienced increasing dryness, 35 affecting its socioeconomic, agricultural, and environmental conditions negatively, which is attributed to lack of 36 rains, high evapotranspiration as well as over-exploitation of water resources. Increase in climatic and weather 37 extremes is expected to aggravate the problem of pollution and flooding. While most urban centers in Asia have no 38 sewers, aging infrastructure may hinder the presently operational sewer systems, particularly in Central Asia (IPCC, 39 2012).

40

41 Coastal Areas. By the year 2025, 70% of Asia's urban population will live in the coastal areas, with the majority 42 located in low-elevation coastal zones (Balk *et al.*, 2009). Climate change is expected to increase the risk of 43 cyclones, flooding, landslides and drought, the adverse events which have direct influence on urban and rural 44 settlements, infrastructure and industries alike. Large parts of South, East and South-east Asia is exposed to higher 45 degree of cumulative climate related risk (UN-Habitat, 2011).

- 46
- 47 In absolute terms, Asia has more than 90% of the global population exposed to tropical cyclones (IPCC, 2012).
- 48 Damage due to storm surge is sensitive to any change in the magnitude of tropical cyclones. For example,
- 49 projections for the inner parts of three major bays (Tokyo, Ise, and Osaka) in Japan indicated that a typhoon that is
- 50 1.3 times as strong as the design standard with a sea level rise of 60 cm would cause damage costs of about US\$ 3,
- 51 40, and 27 billion, respectively, in the investigated bays.
- 52
- 53 Exposure of the world's large port cities (population exceeding 1 million inhabitants in 2005) to coastal flooding due

economic and climate changes (Hanson *et al.* 2011). About 40 million people (0.6% of the global population or roughly 1 in 10 of the total port city population in the cities considered) are currently exposed to a 1 in 100 year coastal flood event (Hanson *et al.* 2011). The bulk of exposed assets in Asia are currently concentrated in Japan where 46% of the population, 47% of industrial production and 77% of commercial sales are concentrated in oceanfront cities, towns and villages (Yasuhara, et al., 2011). Mumbai, Kolkata, Dhaka, Guangzhou, Ho Chi Minh City, Shanghai, Bangkok, Rangoon, and Hai Phòng will be the cities with the greatest population exposure to coastal flooding in 2070 (IPCC, 2012).

8

9 Port authorities from around the world perceive sea-level rise as an issue of great concern especially in the next 10 century (Becker *et al.*, 2011). There is consensus that planned rapid expansion of ports should take into account 11 adaptation measures as ports construct new infrastructure that may still be in use at the end of the century.

12

13 *Population and Assets.* By the 2070s, the top Asian cities in terms of population exposure (including all

14 environmental and socioeconomic factors), are expected to be Kolkata, Mumbai, Dhaka, Guangzhou, Ho Chi Minh

15 City, Shanghai, Bangkok, Rangoon, and Hai Phòng (Nicholls *et al.* 2008). The top Asian cities in terms of assets

16 exposed included Guangdong, Kolkata, Shanghai, Mumbai, Tianjin, Tokyo, Hong Kong, and Bangkok. Hence,

17 cities in Asia, particularly those in China, India and Thailand, become even more dominant in terms of population

and asset exposure, as a result of the rapid urbanization and economic growth expected in these countries". This

19 study also estimates that by 2070, population and asset exposure within Asia's large port cities will be

- disproportionately concentrated in China, India, Japan, Thailand, Vietnam, Bangladesh, Myanmar and Indonesia
 (Nicholls, 2008).
- 21 22

Cities susceptible to human-induced subsidence (mainly, developing county cities in deltaic regions with rapidly
 growing populations) could see significant increases in exposure due to human-induced subsidence as shown
 historically in several Asian cities (Nicholls, 2008).

26

Settlements on unstable slopes or landslide prone areas face increased prospect of rainfall induced landslides.
 Disturbance in water-cycle due to changing climate is already affecting agriculture output but also resulting into
 serious socio-economic problems forcing people to either fall into vicious circle of poverty or migrate.

30

Water-scarcity, especially in summer, is now beyond the control of local governments in urban areas in a number of cities and towns in Asia. Groundwater sources, which are affordable means of usually high-quality water supply in cities of developing countries, are threatened due to over withdrawals. Aquifer levels have fallen by 20 to 50 meters

in cities such as Bangkok, Manila and Tianjin and between 10 and 20 meters in many other cities (UNESCO, 2012).
 The drop in groundwater levels often results in land subsidence, which can enhance hazard exposure due to coastal

inundation and sea-level rise especially in settlements near the coast, and deterioration of groundwater quality.

37

38 The impacts on human settlements and living facilities can be summarized as: (i) increasing shortage of water

39 resources, climate change has been shaping the Yangtze River Delta and its socioeconomic development

40 (Immerzeel, 2010; Vineis, 2011; Shrestha, 2011; Gu et al., 2011; Kang, et al., 2009); (ii) growth in health care

41 expenditure (Ebi et al., 2007; McBean, 2009); (iii) impact on seasonal tourism, the simulation results from seven

42 Japanese ski grounds show that the temperature increase of 3 degree Celsius will cause 30 percent decrease of skiers

43 (Chunyan, et al, 2010; Jianming, et al., 2010; Jian-chao et al., 2011); (iv) impact on livelihoods, the combination of

44 social impacts (e.g. loss of livelihood, displacement) and economic impacts (e.g. damage to industry) could have

45 cumulative or multiplicative effects that eventually interfere with the function and activity of communities within

urban areas (Lioubimtseva et al., 2009; Binyi, et al., 2010; Gasper, et al., 2011); (v) physical and mental health of

47 residents that is closely related to climate change, where cold climate easily causes depression (Jin-qi, et al., 2010;

- 48 Yingjun, et al., 2010; Yonghing, et al. 2008).
- 49

50 *Industry and Infrastructure.* The impacts of climate change on industry include the direct impacts on industry

- 51 production and the indirect impacts on industrial enterprises due to the implementation of the mitigate activities (Li,
- 52 2008). The impact of climate change on infrastructure deterioration cannot be ignored, but can be addressed by
- 53 changes to design procedures including increases in cover thickness, improved quality of concrete, and coatings and

barriers (Stewart, et al., 2012). Climate change and extreme events may have the greater impact on large and
 medium-sized construction projects (Kim, 2007).

3

Climate change has little influence on general travel decisions for tourism, even though weather extremes such as tropical storms are relevant, as revealed by a case study from Israel (Gossling and Hall, 2006). Tourist perceptions of weather and climate vary widely. Many Asian countries are major tourist destinations and more studies are needed to understand the impact of climate change on tourism. With respect to beach tourism, large developing countries and small islands states may be among the most vulnerable due to high exposure and low adaptive

- 9 capacity (Perch-Nielsen, 2010). A number of Asian countries were found vulnerable in this regard.
- 10 11

12 24.4.5.4. Vulnerabilities to Key Drivers13

The impacts of climate change on human settlements, industry and infrastructure will not only be due to sea-level rise and extreme weather events. Disruption of basic services such as water supply, sanitation, energy provision, and transportation system have implications on local economies "and strip populations of their assets and livelihoods", in some cases leading to mass migration. Such impacts are not expected to be evenly spread among regions and cities, across sectors of the economy or among socioeconomic groups. They tend to reinforce existing inequalities

and disrupt the social fabric of cities and exacerbate poverty" (UN-Habitat, 2011).

20

21 A study of Chittagong, Bangladesh concludes that urban adaptation and strengthening of local government capacity

to reduce vulnerability of the urban poor is not considered a priority in national climate change adaptation policy

23 (Ahammad, 2011). As a result, those most at risk from climate extremes are not given adequate attention. In addition,

unequal access to education, health and other public services not only contribute to increase in income disparities, but can also weaken resilience to climate extremes. ADB reported that in the last two decades, 11 economies of Asia,

- which account for more than four-fifths of the region's population have also experienced widening gap between rich
- and poor (ADB, 2012). These development challenges can negatively affect impacts of climate extremes and
- 28 undermine opportunities arising from adaptation.
- 29

Rapid economic growth in Asia is translating into land use related changes, faster construction of buildings and infrastructure, and corresponding industrial development. While such development is improving the quality of life, it

is also creating more impervious surfaces creating both localized heat-island effect as well as flooding in dense
 urban built environments. UN-Habitat (2011) informs that "Climate change has direct effects on the physical
 infrastructure of a city – its network of buildings, roads, drainage, and energy systems – which in turn impact the
 welfare and livelihoods of its residents". The increasing frequency and intensity of extreme climatic events and

siow-onset changes will increase the vulnerability of urban economic assets and subsequently the cost of doing business.

- 38
- 39

41

40 24.4.5.5. Adaptation Options

An ADB and UN report estimates that "about two thirds of the \$8 trillion needed for infrastructure investment in Asia and the Pacific between 2010 and 2020 will be in the form of new infrastructure, which creates tremendous opportunities to design, finance and manage more sustainable infrastructure" (ADB, 2012). A recent study estimated that direct and indirect losses for a 1-in-100 year flooding in Mumbai could triple by the 2080s compared with the present (increasing from US\$ 700 to 2,305 million), and suggests adaptation measures to reduce future damages

47 (Ranger, et al., 2011).48

49 The massive investment may not be affordable for most of the developing countries of Asia (Zevenbergen and

50 Herath, 2008). Hallegatte *et al.* (2011) suggests that adaptation measures, especially in developing countries, offer a

51 'no regret' solution "where basic urban infrastructure is often absent (e.g. appropriate drainage infrastructure),

- 52 leaving room for actions that both increase immediate well-being and reduce vulnerability to future climate change".
- 53 A comprehensive approach featuring non-structural flood control measures is essential for effectively addressing
- 54 future flood risks in complex urban systems (see Table 24-9). Adaptation measures such as improvement of city's

1 drainage system and extending insurance to 100% penetration, can reduce losses associated with a 1-in-100 year 2 flood event by 50% - 70% (Ranger, et al., 2011). 3 4 **[INSERT TABLE 24-9 HERE** 5 Table 24-9: Summary of adaptation options.] 6 7 The role of urban planning and urban planners is emphasized towards adaptation to climate change impacts (IPCC, 8 2012; Fuchs, 2011). City planners with greater understanding of climate change related hazards and capable to 9 communicating associated risk can effectively utilize spatial planning and social infrastructure as tools for 10 adaptation in cities (Fuchs, et.al. 2011). Climate sensitive urban planning is effective even as long-term adaptation 11 measure if takes into account climate variability including uncertainty, and systems vulnerability and capacity (IPCC, 12 2012). 13 14 Awareness, improved governance, development and local partnerships are essential for promoting resilience and 15 adaptation. This is reflected by the significantly different number of fatalities experienced from the impacts of 16 cyclones Sidr and Nargis in developing Asian countries (IPCC, 2012) and in reducing flood risks in Mumbai. 17 18 Green infrastructure is an important new role in protecting urban areas from the consequences of inevitable climate 19 change (Barber, et al., 2009). Climate change brings a significant effect on the building's cooling and heating load, 20 electricity consumption and the outdoor design conditions for air conditioning systems (Yau, et al., 2011). Climate 21 change is expected to influence the demand for space cooling and heating (Vuuren, et al., 2011). 22 23 24 24.4.6. Human Health, Security, Livelihoods, and Poverty 25 26 24.4.6.1. Sub-Regional Diversity 27 28 Asia is predominantly an agrarian society as is evident from 58% of its total population living in rural areas out of

Asia is predominantly an agrarian society as is evident from 58% of its total population living in rural areas out of which 81.8% are dependent on agriculture for their livelihoods (FAOSTAT, 2011). In addition, agriculture employs 24.7% of total population in these countries and contributes to 15.3% of total value added GDP (FAOSTAT, 2011; World Bank, 2011a). Asia also has high levels of rural poverty compared to the urban poverty, with relatively higher poverty incidence in the 8 least developing countries in the region (FAOSTAT, 2011). The high incidence of rural poverty and hunger is closely related to heavy dependence on natural resources that are directly influenced by changes in weather and climate, indicating a close connection between rural livelihoods and poverty (IFAD, 2010; Haggblade et al., 2010).

36

Though Asia has emerged as an economic power during recent decades, there is still a considerable gap in progress in developmental indicators when compared to rest of the world (World Bank, 2011b). In terms of developmental indicators, Southeast Asia is the third poorest region in the world after Sub-Saharan Africa and Southern Asia, and ranks poorly in terms of labor productivity, access to food, maternal health, and forestation (United Nations, 2009). Consequently, as a large proportion of rural populations depend on agriculture, agriculture has been identified as a key driver of economic growth in the region (World Bank, 2007).

43

Impacts on human security in Asia will primarily manifest due to direct and indirect impacts on water resources, agriculture, coastal areas, resource-dependent livelihoods and on urban settlements and infrastructure, with implications for human health and well-being. To a large extent, regional disparities on account of socio-economic context and geographical characteristics among others, define the differential vulnerabilities and impacts within countries in Asia (Sivakumar and Stefanski 2011; Thomas, 2008).

- 49
- 50

51 24.4.6.2. Observed Impacts52

Floods and health. Epidemics have been reported in the aftermath of floods and storms (Bagchi, 2007) due to
 decreased drinking water quality (Harris *et al.*, 2008; Solberg, 2010), invasions of mosquitos (Pawar *et al.*, 2008),

1 and exposure to rodent-borne pathogens like hantavirus and Leptospira (Kawaguchi et al., 2008; Zhou et al., 2011).

2 Contaminated flood waters in urban environments have caused exposure to pathogens and toxic compounds, as

3 noted in e.g. India and Pakistan (Sohan et al., 2008; Warraich et al., 2011). Mental disorders and posttraumatic

4 stress syndrome are observed in disaster prone areas (Li et al., 2010; Udomratn, 2008), and have in India been 5 linked to age and educational level (Telles et al., 2009).

6 7

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Heat and health. The effects of heat on mortality and morbidity, mainly in terms of hospital admission, have been studied in many countries throughout Asia, with specific focus on effects among the elderly and persons with cardiovascular and respiratory disorders (Guo et al., 2009; Huang et al., 2010; Kan et al., 2007). Linear correlations between temperature rise and mortality have been shown for India (McMichael et al., 2008) and several cities in East Asia (Chung et al., 2009; Kim et al., 2006). Several studies have analyzed health effects of air pollution in combination with increased temperatures (Lee et al., 2007; Qian et al., 2010; Wong et al., 2010; Yi et al., 2010).

12 13 Intense heatwaves have also been shown to affect outdoor workers in South and East Asia (Hyatt et al., 2010; Nag et 14 al., 2007).

15

16 Drought and health. Prolonged drought in combination with windy conditions increase the exposure to sand and 17 dust, often mixed with toxic compounds (Wang et al., 2011). There are indications that dust storms in South West,

18 Central and East Asia increase hospital admissions and worsen asthmatic conditions, as well as cause skin and eve

19 irritations (Griffin et al., 2007; Hashizume et al., 2010; Kan et al., 2011; Tam et al., 2012). Prolonged drought may

20 also lead to wildfires and haze exposure with increased morbidity and mortality, as observed in Southeast Asia

- 21 (Johnston et al., 2012).
- 22

23 Water-borne diseases. Many pathogens and parasites multiply faster at higher temperatures. Increases in

24 temperatures have been correlated with outbreaks of water-borne diseases in for example East Asia (Huang et al.,

25 2008; Onozuka et al., 2010, Zhang et al., 2007). Other studies from South and East Asia have shown a correlation

26 between diarrheal outbreaks and a combination of higher temperatures and heavy rainfall (Chou et al., 2010;

27 Hashizume et al., 2007; Majra and Gur, 2009). Increasing coastal water temperatures have been correlated with 28 outbreaks of systemic Vibrio vulnificus infection in Israel (Paz et al., 2007) and Taiwan of China (Kim and Jang,

29 2010). Cholera outbreaks in coastal populations in South Asia have been associated with increasing water

30 temperatures and algal blooms (Huq et al., 2005).

31

32 Vector-borne diseases. Increasing temperatures affect vector-borne pathogens during the extrinsic incubation period 33 and shortens the life-cycles of arthropod vectors, thereby facilitating for larger vector populations and enhanced

34 disease transmission. Several Asian studies have focused on the emergence of dengue fever. Outbreaks have been

35 correlated with temperatures and rainfall (Sriprom et al., 2010; Hsieh and Chen 2009, Nitatpattana et al., 2008;

36 Shang et al., 2010; Su, 2008), in one study with linear correlations with a time lag of 1-3 weeks (Hii et al., 2009).

37 Outbreaks of the vaccine-preventable Japanese encephalitis have been linked to rainfall in studies from the

Himalayan region (Bhattachan et al., 2009; Patridge et al., 2007), and to a combination of rainfall and temperatures 38

39 in South and East Asia (Bi et al., 2007; Murty et al., 2010). Malaria prevalence is often influenced by other factors

40 than climate variability, but studies from India and Nepal have found correlations with rainfall (Dahal, 2008; Dev

41 and Dash, 2007; Devi et al., 2006; Laneri et al., 2010), whereas temperatures were linked to malaria distribution and 42

seasonality in Saudi Arabia (Kheir et al., 2010). Re-emergence of malaria in central China has been suggested to be 43

explained by rainfall and increases in temperature close to water bodies (Zhou et al., 2010). Temperature,

44 precipitation, and virus-carrying index among rodents have been found to be correlated to the prevalence of

45 hemorrhagic fever with renal syndrome in China (Guan et al., 2009; Yan et al., 2008).

46

47 *Livelihood and Poverty.* There have been significant changes in terms of livelihood diversification in Asia over the 48 decades due to rapid economic development (see Table 24-10). Estimates suggest that currently about 51% of total

income in rural Asia come from non-farm sources (Haggblade et al., 2010; Haggblade et al., 2009), out of which 49

50 major proportion comes from local non-farm business and employment. There has also been steady growth in the

proportion of remittances contributing to rural income (Estudillo and Otsuka, 2010). Asia has made significant 51

- 52
- improvement in poverty eradication over the past decade (World Bank, 2008). At the sub-regional level, the East 53 Asia has recorded much rapid reduction in poverty followed by South Asia (IFAD, 2010). Significant part of this

reduction has come from population shift, rapid growth in agriculture, and urban contribution (Janvry and Sadoulet,
 2010).

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[INSERT TABLE 24-10 HERE

5 Table 24-10: Summary of observed changes and projected impacts for livelihoods and poverty.]

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24.4.6.3. Projected Impacts

10 Health effects. An emerging interregional public health concern in Asia is increasing mortality and morbidity due to 11 heat waves. An ageing population in Asia will increase the number of people at risk, i.e. the elderly, and especially 12 those with cardio-vascular and respiratory disorders. The rapid urbanization and growth of megacities in Asia add to 13 the magnitude of the problem with the urban heat island effect that may increase downtown temperatures 14 considerably compared to surrounding rural areas (Tan et al., 2010), even though local adaptation of the built 15 environment and urban planning will define the magnitude of the impacts on public health. The relationship between temperature and mortality often show a U-shaped curve (Guo et al., 2009). Studies from both tropical and temperate 16 17 environments in Asia show increased mortality in particular in rural environments during cold events, even if 18 temperatures do not fall below 0°C (Hashizume et al., 2009; Wu et al., 2011). However, some studies on cold-19 related deaths in developing areas suggest that other factors than climate are important contributors here, and that 20 climate change will not decrease cold-related deaths to any larger extent in such environments (Honda and Ono 21 2009).

22

23 Climate change will affect the local transmission of many climate-sensitive diseases. Increases in heavy rain and 24 temperature are projected to increase the risk of diarrhoeal diseases in for example China (Zhang et al., 2008). The 25 impact of climate change on malaria risk will differ between areas, as projected for e.g. West and South Asia 26 (Husain and Chaudhary, 2008; Garg et al., 2009; Majra and Gur, 2009). Some studies have developed climate 27 change-disease prevalence models, for example for schistosomiasis in China that shows an increased northern 28 distribution range of the disease with climate change (Kan et al., 2011, Zhou et al., 2008). Impacts of climate change on fish production (Qiu et al., 2010) is being studied, along with impacts on chemical pathways in the marine 29 30 environment and consequent impacts on food safety (Tirado et al., 2010b), including seafood safety (Marques et al.,

2010).
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2008). Rural poverty in parts of Asia could be exacerbated (Skoufias et al., 2011) due to negative climate change
 impacts on the rice crop and increase in food price and cost of living (Hertel et al., 2010; Rosegrant, 2011). Poverty

impacts of the fice crop and increase in food price and cost of fiving (Herter et al., 2010, Rosegrant, 2011). Fover impacts of climate change would be heterogeneous among countries and social groups [see Table 24-10]. In low

crop productivity scenario, food exporters such as Indonesia, Philippines and Thailand would benefit from climate

- 39 change related global food price rise and be able to reduce poverty while countries such as Bangladesh would
- 40 experience a net increase in poverty to the tune of 15% by 2030 (Hertel et al., 2010). Regression studies conducted
- by Skoufias et al. (2011) indicate significant negative impacts of shortfall in rainfall on the welfare of rice farmers in
- 42 Indonesia, compared to the delay in onset of rainfall. These impacts may lead to global mass migration and related
- 43 conflicts (Laczko and Aghazarm, 2009; Barnett and Webber, 2010; Warner, 2010; World Bank, 2010).
- 44 45
- 46 24.4.6.4. Vulnerabilities to Key Drivers
- 47

48 Key vulnerabilities vary widely within the region. Climate change can exacerbate current socio-economic and

49 political disparities and add to the vulnerability of Southeast Asia and Central Asia to security threats that may be

transnational in nature (Jasparro and Taylor, 2008; Lioubimtseva and Henebry, 2009). Apart from detrimental

51 impacts of extreme events the vulnerability of livelihoods in agrarian communities also arise from geographic

52 settings, demographic trends, socio-economic factors, access to resources and markets, unsustainable water

- 53 consumption, farming practices and lack of capacity to adapt (Mulligan et al., 2011; Acosta-Michlik and Espaldon,
- 54 2008; Allison et al., 2009; Knox et al., 2011; Lioubimtseva and Henebry, 2009; Byg and Salick, 2009; Salick, 2009;

1 Salick et al., 2009; Xu et al., 2009; Winters et al., 2009; United Nations, 2009). Urban wage labourers were found to

2 be most vulnerable to cost of living related poverty impacts of climate change than those who directly depend on

3 agriculture for their livelihoods (Hertel et al., 2010). In Southeast Asia, an important topic of focus is forest and

4 landfires; for example vulnerability of agriculture, forestry and human settlements on peat land areas in Indonesia

5 (Murdiyarso and Lebel, 2007). Human health is also a major area of focus for Asia (Munslowa and O'Dempseya,

2010), where the magnitude and type of health effects from climate change will differ within Asia depending on
 differences in socio-economic and demographic factors, health systems, the natural and built environment, land use

8 changes, and migration in relation to local resilience and adaptive capacity.

9 10

24.4.6.5. Adaptation Options

13 Cross-sectorial collaborations will be needed for the development of sustainable adaptive measures with interactions 14 between the health sector and disaster preparedness programs, water management, sanitation, urban planning, food 15 industry and the animal health sector. Disaster preparedness on a local community level could include a combination 16 of indigenous coping strategies, early-warning systems, and adaptive measures (Paul and Routray, 2010). Heat 17 warning systems have shown to be successful in preventing deaths among risk groups, like in Shanghai (Tan et al 18 2007). Also proven successful are the implementation of new work practices to avoid heat stress among outdoor 19 workers, as shown in studies from Japan and UAE (Joubert et al., 2011; Morioka et al., 2006). As described in 20 section 24.7 there are many win-win solutions for public health from the interaction of adaptation and mitigation 21 measures that involve urban environments and air pollution. Early warning models have been developed for haze 22 exposure from wildfires, in e.g. Thailand (Kim Oanh and Leelasakultum, 2011). Early warning models are also 23 being tested in infectious disease prevention and vector control programs, like for malaria in Bhutan (Wangdi et al.,

24 2010) and Iran (Haghdoost *et al.*, 2008), or are being developed, like for dengue fever region-wide (Wilder-Smith *et al.*, 2012).

25 26 27

Available literature suggests the need for identifying and promoting technologies and policy options that will provide both mitigation potential as well as sustained income generation potential in a changed climate (Bhan

provide both mitigation potential as well as sustained income generation potential in a changed climate (Bhandari *et al.*, 2007; Rosenzweig and Tubiello, 2007; Paul *et al.*, 2009). Interesting examples seem to emerge on how some practices provide completely unexpected livelihood benefits which otherwise may not be captured in standard

31 evaluation frameworks, as in the case of introduction of traditional flood mitigation measures in China could

- 32 positively impact the local livelihoods leading to both reductions of physical and economic vulnerabilities of
- 33 communities (Xu *et al.*, 2009). Significant amount of literature has stressed for the greater role of local communities
- 34 in decision making (Alauddin and Quiggin, 2008) and in prioritization and adoption of adaptation options
- 35 (Prabhakar *et al.*, 2010; Prabhakar and Srinivasan, 2011). Defining adequate community property rights, including
- 36 solving the issues such as land tenure, reducing income disparity, exploring market based and diversified off-farm
- 37 livelihood options, moving from production based approaches to productivity and efficiency decision making based
- approaches, and promoting integrated decision making approaches were suggested (Merrey *et al.*, 2005; Brouwer *et al.*, 2007; Paul *et al.*, 2009; Niino, 2011; Stucki and Smith, 2011).
- 39 *al* 40

41 Climate resilient livelihoods can be fostered through the creation of a bundle of capitals (natural, physical, human,

42 financial and social capital) and poverty eradication (see Table 24-11). In general, greater emphasis on agriculture

43 growth has been suggested as an effective means of reducing rural poverty (Janvry and Sadoulet, 2010; Rosegrant,

44 2011). Bundled approaches are known to facilitate better adaptation than individual adaptation options (Acosta-

- 45 Michlik and Espaldon, 2008; Fleischer et al., 2011). Community based approaches, as against top-down
- 46 interventions, have been suggested to identify adaptation options that address poverty and livelihoods, as these
- 47 techniques help capture information at the grassroots (Aalst et al., 2008), help integration of disaster risk reduction,
- 48 development, and climate change adaptation (Heltberg et al., 2010), connect local communities and outsiders (Aalst
- 49 et al., 2008), and addresses the location specific nature of adaptation (Iwasaki, et al., 2009; Rosegrant, 2011). Some 50 groups can become more vulnerable to changes after being 'locked into' specialized livelihood patterns as shown in
- 50 groups can become more vulnerable to changes after 51 the case of fish farmers in India (Coulthard, 2008).
- 52
- 53 [INSERT TABLE 24-11 HERE
- Table 24-11: Summary of adaptation options for securing livelihoods in Asia.]

1 2 3

24.4.7. Valuation of Impacts and Adaptation

4 5 Research on the valuation of climate change impacts and adaptation in Asia has been highly limited. However, 6 recently there is growing attention to the research efforts of assessing aggregate costs of climate change impacts and 7 adaptation. There are a few studies focusing on disperse sectors though without comprehensive economic valuation 8 of the costs and benefits of adaptation. Examples of such studies include exploring low-cost adaptation strategies to 9 reduce the net vulnerability of sorghum production system in India (Srivastava et al., 2010); assessing vulnerability 10 and adaptation of agriculture and food security, water resources and human health in Central Asia (Lioubimtseva 11 and Henebry, 2009); socio-economic impacts of drought and flood in South Asia (Muhammed, et al., 2007); 12 investigation of vulnerability and adaptive capacity to climate variability and water stress in the Lakhwar watershed 13 in Uttarakhand State, India (Kelkar et al., 2008), assessing socio-economic vulnerability and adaptation measures in 14 West Coast of Peninsular Malaysia (Drainage and Irrigation Department, 2007); and simulation impacts on rice 15 yields in a number of Asian countries (Matthews et al. 1997). In addition to changes in temperature and rainfall, 16 changes in the frequency of extreme climatic events could be damaging and costly to agriculture (Aydinalp and 17 Cresser, 2008; Muhammed et al., 2007; Su et al., 2009).

18

19 A study of the economics of climate change in Southeast Asia (ADB, 2009) with focus on Indonesia, Philippines,

20 Thailand, and Vietnam reported that many of the impacts from climate change are not in traditional economic

sectors such as agriculture including fisheries and aquaculture, forestry and mining, with the result that their valuations are difficult with uncertainly and incomplete information. Furthermore, some of the economic and social

valuations are difficult with uncertainly and incomplete information. Furthermore, some of the economic and valuations, such as loss of life or damage to ecosystems, can be contentious. Without further mitigation or

adaptation (under the A2 scenario of IPCC, 2000), the PAGE2002 integrated assessment model projects for the four

countries to suffer a mean loss of 2.2% of gross domestic product (GDP) by 2100 on an annual basis, if only the

26 market impact (mainly related to agriculture and coastal zones) is considered. This is well above the world's mean

GDP loss of 0.6% each year by 2100 due to market impact alone. In addition, the mean cost for the four ASEAN

countries by 2100, could reach 5.7% of the GDP if non-market impacts related to health and ecosystems are

included and 6.7% of the GDP if catastrophic risks are also taken into account.

31 The PAGE2002 model also found that the cost of adaptation for the agriculture and coastal zones (mainly the 32 construction of sea walls and development of drought- and heat resistant crops) would be about \$5 billion/year by 33 2020 on average, and that this investment would be paid in the future. For instance, the annual benefit of avoided 34 damage from climate change is likely to exceed the annual cost by 2060 and by 2100, benefits could reach 1.9% of 35 GDP, compared to the cost at 0.2% of GDP with the results at mean and 5% probability level under the A2 scenario. 36 This shows that the benefits from adaptation are projected to outweigh the costs of implementing adaptation 37 measures in the long term. It was also stressed that there are currently great uncertainties associated with the 38 economic aspects of climate change (ADB, 2009). Adaptation cannot entirely remove the projected damage of 39 climate change, and thus must be complemented with global mitigation of CO2 in order to avoid the greater impact

40 of future climate change (Begum *et al.*, 2011; ADB, 2009; MNRE, 2010).

41 42

43 24.5. Adaptation and Managing Risks

44 45

24.5.1. Conservation of Natural Resources

46 47 Even without climate change, natural resources are already under severe pressure in most of East, Southeast, and 48 South Asia, as well as in much of Central and West Asia, and parts of North Asia and the Tibetan Plateau. The 49 extraordinarily high rates of deforestation and forest degradation in Southeast Asia have received most attention 50 (Sodhi et al., 2010; Miettinen et al., 2011), but ecosystem degradation, with the resulting loss of natural goods and services, is also a major problem in other forest types and in non-forest ecosystems. These pressures result from 51 52 rising populations and rapid economic development, exacerbated by poor governance and the low priority of natural 53 resource conservation. The impacts of projected climate change are expected to intensify these pressures in most 54 areas, but the relative importance of climate and non-climate stressors is difficult to predict in most cases. Coral

reefs are an exception, with climate change and ocean acidification a clear threat to all reefs in the region and thus
the millions of people who depend on them (Hoegh-Guldberg, 2011; Burke *et al.*, 2011; see also Chapter 30, this
volume).

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5 With natural resource conservation already under stress, the focus has been on actions that would be beneficial even 6 without climate change, including minimizing non-climate pressures on natural resources and restoring connectivity 7 to allow movements of genes and species between fragmented populations (Lindenmayer et al., 2010). Authors have 8 also suggested a need to identify and prioritize for protection areas that will be subject to the least damaging climate 9 change ('climate refugia') and to identify additions to the protected area network that will allow for expected range 10 shifts, for example by extending existing protected areas to higher altitudes or latitudes (Hannah, 2010; Hole et al., 11 2011; Shoo et al., 2011). Assisted migration may be useful for some species in fragmented landscapes (Thomas, 12 2011). More generally, conservationists may need to consider abandoning the current focus on the preservation and 13 restoration of 20th century reference conditions, which may no longer be relevant in a changing world (Thomas, 14 2011).

15 16

17 24.5.2. Flood Risks and Coastal Inundation

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Many coastal areas in Asia are anticipated to face threats of flood risk and coastal inundation exacerbated by climate change. Responding to a large number of climate change impact studies for each country over the past decade (e.g. Huang et al., 2004; Karim and Mimura, 2008; Pal and Al-Tabbaa, 2009), various downscaled tools to support formulate and implement climate change adaptation policy for local governments are under development. One of the major tools is vulnerability assessment and identifying policy options with Geographical Information System (GIS). As a matter of course, these have been developed for flood risk management so far, most of the tools have begun to give consideration in varying degrees to climate change impact such as sea level rise in long term.

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27 In India, for example, coastal vulnerability index for mainly sea level rise were calculated and mapped to inform the 28 vulnerability in each area of the west coast (Dwarakish et al., 2009), and physical and social vulnerability to storm 29 surge considering climate change ware mapped in each area of the east coast (Rao et al., 2010). In Bangladesh, a positive relationship between flood risk, poverty and socioeconomic vulnerability was identified and an importance 30 31 of preparedness of the poor household and support from community were indicated from a case study on the 32 southeast region (Brouwer, et al. 2007). In Indonesia, involvement of stakeholder and community were proposed to 33 improve the existing flood risk management from a case study in outlying city of Dhaka (Marfai and King, 2008), 34 and actually community based vulnerability assessment was implemented to identify various adaptation measures 35 (Taylor, 2011). The similar approach was conducted in the central province in Vietnam. Integrating technology like 36 GIS and indigenous local knowledge through the participatory technique was stressed in this case (Tran et al., 2009). 37 Also in Ho Chi Minh City in Vietnam, intensive approach to integrate climate change adaptation policy and urban 38 planning is in progress providing a toolkit which aims at the empowerment of local decision makers and other 39 relevant actors providing a broad range of potential options for climate change adaptation policy (Schwartze et al., 40 2011; Storch et al, 2011). 41

42 All these tools and systems tend to have a direction of community-based approaches. These approaches have been 43 salient over the past two or three decades in environmental policy, disaster risk management and so on. Behind these 44 backgrounds, there is a growing recognition that such approaches are indispensable to reduce vulnerability and to 45 build adaptive capacity effectively. However, a key challenge for a successful implementation of community-based 46 approaches is to keep it easy enough for wider application (Van Aalst, et al. 2008). Also it requires an understanding 47 of the community structure and other factors while the approaches have primary weakness of lack of resource and 48 decision-making, legislative and regulatory powers available to local-level actors (Allen, 2006). One of the key 49 components to overcome the weakness is social capital. In Vietnam, climate change adaptation strategies were facilitated by social capital that emerged in the absence of governmental support or frameworks (Adger, 2003). It 50 implies that community-based approaches have possibilities to vary depending on works of social capital in the 51 52 context of the community.

53 54 1 2

24.5.3. Economic Growth and Equitable Development

3 Economic, social, and environmental equity is an enduring challenge in many parts of Asia. Attempts have been 4 made to use the level of wealth (typically GDP) as a measure of human vulnerability of a country or region, but this 5 approach has serious limitations. In many cases, social capital, an indicator of equity in income distribution within 6 countries, is a more important factor of vulnerability and resilience than GDP per capita. Furthermore, political and 7 institutional instabilities can undermine the influence of economic development (Lioubimtseva and Henebry, 2009). 8 Poor and vulnerable countries are at greater risk from the impacts of climate extremes as their options for coping 9 which such events are limited. This is particularly true for developing countries in Asia with a high level of natural-10 resource dependency. Provision of adequate resources based on the burden sharing and the equity principle will 11 serve to strengthen appropriate adaptation policies and measures in such countries (Su et al., 2009). Mainstreaming 12 adaptation into government's sustainable development policy portrays a potential opportunity for good practice to 13 build resilience and reduce vulnerability depending on effective, equitable and legitimate actions to overcome barriers and limits to adaptation (Lioubimtseva and Henebry, 2009; Agrawala and van Aalst, 2005; Lim et al., 2005; 14 15 ADB, 2005). It requires growth with economic stability, development with social equity and poverty eradication, 16 and the continued functioning of ecosystems as life support systems to sustain development.

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19 24.5.4. Mainstreaming and Institutional Barriers

21 The level of climate change adaptation mainstreaming is most advanced in the context of official development 22 assistance where donor agencies and international financial institutions have taken significant steps in taking into 23 account climate change adaptation in their loan and grant making process (Gigli and Agrawala, 2007; Klein et al., 24 2007b). In contrast, in developing countries, actual projects on the ground to mainstream adaptation to climate 25 change remains limited and significant institutional and cognitive barriers remain (Yohe et al., 2007; Gigli and 26 Agrawala, 2007). For example, in the Philippines, the reasons that hindered climate change mainstreaming are the 27 following: national priorities are geared towards what are perceived to be more pressing concerns such as 28 employment generation and education and a pervasive lack of awareness on the impacts of climate change to 29 sustainable development (Lasco *et al.*, 2009). However, there are massive investments on infrastructure projects 30 designed to adapt to weather-related hazards. Local government units could play a crucial role as shown by the 31 experience of Albay province in the Philippines which pioneered climate action at the grassroots level (Lasco et al., 32 2012)

33

34 While some practical experiences of adaptation in Asia at the regional, national and local level are emerging, there 35 can be barriers that impede or limit adaptation. This can include lack of financial resources for adaptation 36 implementation, institutional barriers, biophysical limits to ecosystem adaptation etc. (Moser and Ekstrom, 2010). 37 Regional adaptation strategies are necessary to tackle issues such as food security. There are already some groups such as the Association of South East Asian Nations (ASEAN) but there is need for global and regional strategic 38 39 partnerships in this regard (Singleton et al., 2010). The success of deployment, implementation and sustainability of 40 adaptation options can be influenced by the political economy of the region. Issues with resource availability might 41 not only be as a result of climate change but also weak governance mechanisms and breakdown of policy and 42 regulatory structures, especially in the context of common-pool resources (Moser and Ekstrom, 2010). Furthermore, 43 this impact depends on the inherent vulnerability of the socio-ecological systems in a region, as much as on the 44 magnitude of climate impact (Evans, 2010). Recent studies linking climate-related resource scarcities and conflict, 45 call for enhanced regional cooperation (Gautam, 2012).

46 47

48 24.5.5. Role of Higher Education in Adaptation and Risk Management 49

50 To enhance the young professional development in the field of climate change adaptation, it is of utmost importance

51 to include the topic in the higher education, especially in the formal education programs. Shaw et al. (2011) 52

emphasized the need of higher education in adaptation and disaster risk reduction in the Asia-Pacific region through: 53 environment disaster linkage, focus on hydro-meteorological disasters, and emphasizing synergy issues adaptation

54

Nomura and Abe (2010), Chhokar (2010) and Niu et al. (2010). Higher education should be done through lectures and course work, field studies, internship, and establishing education-research linkages by exposing the students to field realities. In this regard, a few guiding principles should be: inclusive curriculum, theoretical focus, field orientation, multi-disciplinary courses and practical skill enhancement. Bi-lateral or multi-lateral practical research programs on adaptation and risk management by the graduate students and young faculty members would expose them to the real field problems.

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24.6. Intra-regional and Inter-regional Issues

24.6.1. Trade and Economy

13 A well-functioning international trading system can support the adaptation to climate change-related challenges. 14 Hence welfare gains from reforms to trade policies may be greater than normally measured if they also reduce GHG 15 emissions globally (Huang et al., 2011). In recent years, there has been a growing interest in the environmental 16 impacts of regional trade liberalization. A study by Gumilang, et al. (2011) suggests that overall AFTA (ASEAN 17 Free Trade Agreement) has a greater impact on the Indonesian economy compared to IJEPA (Indonesia-Japan 18 Economic Partnership Agreement) while the adoption of both agreements contributes to increasing CO2 emission by 19 0.47% compared to the BAU case. This is mainly due to a high emission coefficient by the transportation sector. On 20 the other hand, the agreements did have a positive impact on water pollution indicators. 21

China's high economic growth flourishing trade activities on both domestic and international levels have resulted in significant amounts of water withdrawal and water pollution. For instance, Guan and Hubacek (2007) found that North China as a water scarce region virtually exports about 5% of its total available freshwater resources while accepting large amounts of wastewater for other regions' consumption. By contrast, South China a region with abundant water resources is virtually importing water from other regions while their imports are creating waste water polluting other regions' hydro-ecosystems. Thus, the effective trade liberalization and regional trade policy might be useful to mitigate some of major climate change challenges affecting the environment and health such as

29 air pollution, water scarcity and security as well as waste management.

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32 24.6.2. Migration and Population Displacement

Migration has received attention in the literature as an adaptation option (Reuveny, 2007; Warner, 2010). Studying environment and other natural resources-induced migration can help to effectively manage climate change induced migration (Reuveny, 2007). While some form of environmentally induced migration may be adaptive, other forms of environmental migration may indicate a failure of social-ecological systems to adapt (Warner, 2010), suggesting need for differentiating the root cause of migration and treating them through new forms of governance that connects the migrants with those who returned and remained.

40

Migration has become one of the strategies to sustain livelihoods in the wake of climate and environmental change (Barnett and Webber, 2010). The shift towards non-farm income activities, including migration, appears to be more prominent in countries and communities with least access to land (Winters et al., 2009) and in those communities with better access to education (Estudillo and Otsuka, 2010). Rapid-onset environmental change such as floods, as in the case of Mekong Delta, are increasingly playing a role in migration (Warner, 2010). These migration induced remittances have significantly contributed to Asian economies and decreased the poverty gap but had negligible effect on poverty rate (Vargas-Silva et al., 2009).

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50 24.7. Adaptation and Mitigation Interactions

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52 Climate change mitigation benefits climate change adaptation in Asia by increasing the prospects that adaptation can

address many unavoidable impacts, and adaptation benefits mitigation by somewhat moderating impacts of

54 particular GHG concentration levels due to reduced sensitivities or increased coping capacities. One of the most

prominent examples is increasing the efficiency and affordability of air conditioning, which would extend space conditioning benefits to a larger share of populations with rising standards of living while at the same time reducing carbon emissions associated with electricity generation. Other examples include the development of sustainable cities in Asia with less fossil fuel driven vehicles (mitigation) and with more trees and greenery (carbon storage as well as adaptation to urban heat island effect), which would have a number of co-benefits including public health – a promising strategy for "triple win" interventions (Romero-Lankao et al, 2011). A further example is China's leadership in promoting solar energy technologies, where reduced requirements for carbon-based electricity generation are combined with technological change, job creation, and skill development that enhance adaptive capacities.

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11 Other possible synergies (and/or conflicts) are likely to be more subtle. In general, integrated mitigation and

12 adaptation responses tend to focus on either land use changes, often involving ecosystem functions, or on

technology development and use. For instance, changes in land use, such as agroforestry, may provide both

14 mitigation and adaptation benefits (Verchot et al., 2007). Agroforestry practices will provide carbon storage and

15 may at the same time decrease soil erosion, increase the resilience against floods, landslides and drought, increase 16 soil organic matter, reduce the financial impact of crop failure, as well as have biodiversity benefits over other forms

of agriculture as shown in e.g. Indonesia (Clough et al., 2011). Integrated approaches are often needed when

developing mitigation-adaptation synergies, as seen in waste-to-compost projects in Bangladesh (Avers and Hug.

2009). Linking adaptation to mitigation makes mitigation action more relevant for many low-income regions.

20

21 Ecological adaptation measures that increase plant biomass, such as ecosystem protection and reforestation, will

22 contribute to climate mitigation by carbon sequestration. However, exotic monocultures may fix more carbon than

23 native species mixtures while at the same time they decrease biodiversity and contribute less to ecological services.

Biodiversity-rich carbon storage that is resilient to future climate change would be a more sustainable choice (Díaz et al., 2009). The potential for both adaptation and mitigation through forest restoration appears to be greatest in the

tropics (Sasaki et al., 2011). In boreal and high latitudes temperate regions it will also be necessary to consider

albedo effects, with the possibility that adaptation-driven reforestation could have negative consequences for

mitigation by reducing surface albedo (Thompson et al., 2009). On rivers and coasts, the use of hard defenses (e.g.

29 sea-walls, channelization, bunds, dams) to protect agriculture and human settlements from flooding will often have

30 negative consequences for both natural ecosystems and carbon sequestration by preventing natural adjustments to

31 changing conditions. Conversely, setting aside landward buffer zones along coasts and rivers would be positive for

- both (Erwin, 2009), although this will often be difficult in practice.
- 33

34 Several mitigation technologies and measures will have public health benefits, such as controlled composting, state-35 of-the-art incineration, expanded sanitation coverage, and waste water management (Bogner et al., 2008). There are 36 potentially large benefits for both public health and other sectors through climate change mitigation policies that 37 reduce exposure to outdoor and indoor air-pollution (Haines et al., 2009). Decarbonizing electricity production 38 efforts in India and China (coal) are projected to decrease mortality due to reduced PM5 and PM2.5 particulate 39 matters (Markandya et al., 2009). Mitigation policies to reduce fossil fuel vehicles will increase air quality and 40 decrease the health burden in particular in urban environments as projected in India (Woodcock et al., 2009). The 41 use of more public transportation as well as active (bicycling, walking, etc) transports and less private cars could 42 also improve public health (Woodcock et al., 2007). Abandoning the use of biomass fuel or coal for in-door cooking 43 and domestic heating would substantially increase indoor air quality and respiratory and cardiac health among, in 44 particular, women and children in India and China (Wilkinson et al., 2009). In reverse, actions to reduce current 45 environmental-public health issues may often as an additional bonus have beneficial mitigation effects, like traffic 46 emission reduction programs in China (Wu et al., 2011) and in India (Reynolds and Kandlikar, 2008). At the same 47 time, climate change adaptation technologies such as improved stormwater and wastewater management can reduce 48 electricity requirements for water pumping and water treatment; and advances in information, communication, and 49 control technologies can contribute to both adaptation and mitigation efforts. In a number of cases, from Dubai and 50 Abu Dhabi in the western part of the region to Singapore in the eastern part of the region, Asia is becoming a test 51 bed for innovative applications of technologies that are at the frontier of new energy pathways for sustainable 52 development.

53
1 There has also been emphasis on forests and their management for providing resilient livelihoods and reduce 2 poverty (Persha et al., 2010; Larson, 2011; Noordwijk, 2010; Chhatre and Agrawal, 2009). Securing rights to 3 resources was found essential for greater livelihood benefits to the poor indigenous and traditional people (Macchi et 4 al., 2008) for which REDD+ schemes have been urged to respect and promote community forest tenure rights 5 (Angelsen, 2009). It was suggested that indigenous people can provide a bridge between biodiversity protection and 6 climate change adaptation (Salick, 2009) which appears to be missing in the current discourse on ecosystems based 7 adaptation. However, there are arguments against REDD supporting poverty reduction due to its inability to promote 8 productive use of forests that may keep communities in perpetual poverty (Campbell, 2009). Among financial means, 9 low-risk liquidity options such as microfinance programs and risk transfer products can help lift rural poor from the 10 poverty and accumulate assets (Barret et al., 2007; Jarvis et al., 2011). 11 12 START BOX 24-1 HERE 13 14 Box 24-1. Rice-Wheat Systems in India 15 16 Autonomous adaptation may have undesirable impacts. This case shows how adaptation actions today may 17 negatively impact the possibility of future adaptation. In the rice-wheat systems on the Indo-Gangetic plains rice is 18 planted in July, harvested in October-November and then wheat is planted in November and harvested in April. If 19 there are any delays in the system, or if as a result of changing weather patterns temperatures are higher, wheat 20 yields are reduced due to increased temperature during grain filling in March and April. To avoid this, farmers need 21 to plant wheat immediately after rice. Some farmers therefore burn rice residues to vacate fields and to plant wheat 22 in time. This unfortunately increases GHG emissions. Minimum tillage approaches may be appropriate in these 23 circumstances, though incentives to farmers to adopt such practices will need to be put in place.

25 ___ END BOX 24-1 HERE _____

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24.8. **Research and Data Gaps**

30 There are still regions within Asia that are not sufficiently represented in observed climate change studies, in 31 particular Central and West Asia. Also, numerical data on trends in precipitation is hard to find compared to trends 32 in temperature. Furthermore, research data on changes in extreme climate events does not cover most Asian regions. 33 For freshwater resources studies, research priorities are as follows: (1) to increase the knowledge of future rainfall 34 changes in regions by model ensembles to provide a better idea of future water supply, (2) to develop water 35 management strategies across scales to adapt future changes in water demand and supply associated with climate change, (3) to elaborate more study on successful water saving technologies and other adaptation options.

36 37 38 Scientific understanding of the impacts of climate change on ecosystems and biodiversity in Asia is currently limited 39 by the poor quality and low accessibility of biodiversity information (GEO-5 Assessment Report, 2012). National 40 biodiversity inventories are incomplete and very few sites have the accurate baseline information needed to identify 41 changes brought about by climatic trends and other stressors. Quantitative information for sites in protected areas 42 where non-climate impacts are minimized will be particularly valuable in the future. New and old data need to be 43 digitized and made available on-line. Current warming projections suggest that large areas in the Asian tropical 44 lowlands will experience climates in 2100 that have not existed anywhere on Earth for several million years (Wright 45 et al., 2009). This novelty makes reliance on extrapolation from our current, limited, understanding of climatic 46 controls on biological processes dangerous, and underlines the need for new research. Key priorities include the 47 temperature dependence of carbon fixation by tropical trees and the thermal tolerance and acclimation capacity of 48 both plants and animals (Corlett, 2011). 49 50 Boreal forest dynamics will be influenced by complex interactions between rising temperatures and CO_2

- 51 concentrations, permafrost thawing, forest fires, and insect outbreaks (Osawa et al., 2009; Zhang et al., 2011b).
- 52 Understanding this complexity will require enhanced monitoring of biodiversity and especially of species ranges,
- 53 improved modeling, and a greater knowledge of species biology (Anisimov et al., 2008). Long-term monitoring of
- 54 biome boundary shifts and vegetation change is also needed because of slow rate of these changes. In remote and

1 inaccessible areas such monitoring has been provided since 1978 by broad-swath satellite remote sensing data,

- 2 however lack of coincidence in estimates of vegetation vigor provided by remote sensing techniques and by
- vegetation models requires further research and methods intercalibration (Xu *et al.*, 2012).
- 5 There are still many gaps in our understanding of climate change impacts and vulnerabilities in the agricultural
- 6 sector as well as appropriate adaptation options. The most studied crop is rice but there are still significant
- uncertainties in terms of accuracy of models, effect of CO2 fertilization, regional effects (Shuang-1 He *et al.*, 2011;
 Zhang *et al.*, 2010; Masutomi *et al.*, 2009). For other crops, there is even greater uncertainty in terms of magnitude
- 8 Zhang *et al.*, 2010; Masutomi *et al.*, 2009). For other crops, there is even greater uncertainty in terms of 9 and direction of impacts of rising temperatures, precipitation changes, and CO2 fertilization.
- 10
- 11 There is a need to increase the knowledge on heat and air pollution interactions and health effects in different risk
- 12 groups, in both urban and rural environments. There are research gaps on climate change impacts on children's
- health in different socioeconomic and regional context to fill in. More trans-disciplinary research is needed on direct
- and indirect health effects from climate change impacts on water quality and quantity in different parts of Asia.
 Studies on social-economic and institutional dimension should also be given priority. For example, the impacts of
- 16 climate change to women and their role in climate change adaptation need to be investigated (Mula *et al.*, 2010).
- There is also a need to identify low cost options and a need for scaling up of the same, considering the vast majority
- of population living below the poverty line in some of the least developed countries. Greater understanding is
- required on linkages between local livelihoods, ecosystem functions, and land resources for creating positive impact
- on local livelihoods and poverty reduction in areas with greater dependency on natural resources (Paul *et al.*, 2009).
- Research on adaptation and mitigation interactions that promotes sustainable development should be increased, as
- 22 well as research on possible economic gains from different adaptation-mitigation strategies and measures.
- 23 24

More focused research is needed on climate change impacts, vulnerability and adaptation on urban settlements,

- especially cities with population under 500,000, sharing about half of region's urban population. While urban areas
- account for over 80% of region's GDP, detailed estimates on impact of climate change on various sectors of urban
- economy, including tourism industry needs priority. Research priority for promoting adaptation polices at municipal
- 28 level should be given emphasis. It is assumed that the existing policies should be expanded into adaptation; however
- the implementation of adaptation measures is still in its infancy. In order to promote adaptation policies at municipal level, two types of research should be highly prioritized. The first is on research regarding quantitative assessment of
- impacts and adaptation of climate change, which would also include different target years, different stabilized
- 32 purposes, multiple GCM results, and social economic scenarios. This would be useful in determining specific target
- periods and quantitative countermeasure levels, while taking account of the progress of future global warming. In
- this process, uncertainty should be noted in correspondence to climate change scenarios and assessment techniques.
- 35 The second type of research should be action oriented, focusing on implementing adaptation policy, taking into
- 36 account necessary cost and socio-economic innovation. In assessing the quantitative effects of an adaptation policy,
- 37 especially in Asia, researches utilizing various social-economic scenarios are significant to more accurately reflect
- 38 on diversities in a social system, life style, culture, and climate.
- 39
- Climate change will not have uniform impact on a population within a country but rather depends on location, socio economic conditions and level of preparedness (Begum et al, 2011). Negative impacts on agriculture productivity
- 42 would have significant impact on the aggregated household welfare, livelihoods and poverty in the region (Zhai and
- 43 Zhuang, 2009) and this needs to be adequately studied. Low cost options are limted considering the vast majority of
- 44 population living below poverty line in some of the least developed countries such as Bangladesh (Iwasaki *et al.*,
- 45 2009; Rawlani and Sovacool, 2011). Greater understanding is required on linkages between local livelihoods,
- 46 ecosystem functions, and land resources for creating positive impact on local livelihoods and poverty reduction in
- 47 areas with greater dependency on natural resources (Paul *et al.*, 2009). Keeping in view the interconnected nature of
- 48 the problems across geographical, social and political scales, an emphasis on increased regional collaboration in
- 49 scientific research and policy making was suggested for reducing climate change impacts on water, biodiversity and
- 50 livelihoods in Himalayan region (Xu *et al.*, 2009).
- 5152 While mitigation efforts are essential, literature suggests that work must begin on building understanding of the
- 52 while initigation errors are essential, interactive suggests that work must begin on bunding understanding of the 53 impacts of climate change and moving forward with the most cost-effective adaptation measures (Stage, 2010;
- 54 Mathy and Guivarch, 2010; Cai *et al.*, 2008; ADB, 2007). Consequently, for mitigation policies, most cost-effective

1 mitigation measures within sector and across sectors would be the key information needed to devise these policies

2 (Mathy and Guivarch, 2010; Cai et al., 2008; Nguyen, et al., 2007). The costs and benefits of climate change

3 adaptation cannot be analyzed using economic aspects only; other aspects such as climate science, behavioral 4 science, legal and moral aspects also have crucial implications for the outcome of the analysis (Stage, 2010;

Agrawala and Fankhauser, 2008; Lecocq and Shalizi, 2007; Begum et al., 2006; Metroeconomica, 2004).

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24.9. **Case Studies**

10 24.9.1. Transboundary Issues – Mekong River Basin Case Study

12 The lower Mekong River Basin (LMB) covers an area of approximately 606,000 sq km across the countries of Thailand, Laos, Cambodia and Vietnam (Hinkel and Menniken, 2007) [see Figure 24-3]. More than 60 million people in the densely populated LMB are heavily reliant on natural resources, in particular agriculture and fisheries for their well-being (MRC, 2009; UNEP, 2010). As two of the five top rice exporting countries globally, Thailand 16 and Vietnam produced 51% of the world's rice exports in 2008. The majority of rice production in these countries is located in the LMB (Mainuddin et al., 2011a). About two-thirds of the Mekong Basin's population are involved in fishing to sustain their livelihoods; fishing is particularly important for rural households in the LMB (Hortle, 2009;

19 Mainuddin et al., 2011b). Although there is no precise data on fishery exports originating in the LMB, the exports of

- 20 fishery products from the four riparian countries in total reached US\$5.6 billion in 2008 (FAOSTAT, 2008;
- 21 Mainuddin et al., 2011b).

22 23 **[INSERT FIGURE 24-3 HERE**

24 Figure 24-3: Map of Lower Mekong Basin from Mekong River Commission Technical Paper No. 24, 2009 (MRC, 25 2009).]

26

27 Across the LMB countries observations of climate change over the past 30-50 years include (MRC, 2010): increase 28 in temperature (for all riparian countries), changes in rainfall patterns (e.g. Thailand and Vietnam), intensification of 29 flooding and droughts (e.g. Laos) and sea level rise (e.g. Vietnam's Mekong Delta). Agricultural output has been 30 noticably impacted by these climate related events, for example resulting in rice production loss in Cambodia and

31 Laos (1995 – 2001). Negative impacts on capture fisheries in the LMB as a result of climate change as well as dam

32 construction are observed (MRC, 2010; Hortle, 2009; Wyatt and Baird, 2007).

33

34 National level climate change adaptation plans have been formulated in all four riparian countries. A commonly

35 shared scientific forecast on possible future climate impacts as well as an integrated and co-ordinated adaptation

36 program across the LMB does not exist to date. A range of individual studies that assess future LMB climate differ

37 in the use of underlying climate models and emission scenarios. The existing studies however broadly share a set of

expected future climate changes in the Mekong Basin (MRC, 2009): increase in temperature, wet season rainfall, 38

39 flooding frequency and duration along the Mekong River; decrease in dry season rainfall; sea level rise and salinity

40 intrusion in the Mekong delta.

- 42 While significant uncertainties about both magnitude and location-specific impacts of climate change in the LMB remain, it is expected that *vulnerabilities* will be exacerbated in three areas: 43
 - 1. Reduced agricultural output and yields, particularly for rice (MRC, 2009)
 - 2. Loss of fertile land and population displacement in the Mekong river delta (MRC, 2009; MRC, 2010)
 - Reduced fish survival, growth and reproductive success (UNEP, 2010) 3.
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48 To address these vulnerabilities, *adaptation needs* are primarily in areas of improved water management, farming

- and fishing practices as well as coastal protection (Johnson et al., 2010; Hoanh et al., 2003). Transboundary 49
- initiatives to address climate change are driven by multiple actors including the Mekong River Commission (MRC), 50
- the United Nations Development Program (UNDP) and the Asia Development Bank's Greater Mekong Sub-region 51
- programme (ADB GMS) among others (MRC, 2009; Lian and Bhullar, 2011). Despite these initiatives, strong inter-52
- 53 governmental policy development and planning co-ordination between ministries and different levels of government

1 are largely absent, which has adversely affected the development and implementation of appropriate large scale 2 adaptation strategies (Lian and Bhullar, 2011).

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4 Key challenges and barriers for an effective future transboundary adaptation planning and management include: 5

- Lack of a commonly shared scientific forecast on possible future climate impacts across LMB countries as the basis for transboundary adaptation planning (MRC, 2009)
- Sub-optimal co-ordination among adaptation stakeholders and sharing of best-practices across countries ٠ (MRC, 2009)
- Insufficient mainstreaming of climate change adaptation into the broader policy frameworks of the National Governments in all the four LMB countries (MRC, 2009; Lian and Bhullar, 2011)
 - Insufficient integration of transboundary policy recommendations into national climate change plans and policies (Kranz et al., 2010).

14 Currently sub-optimal resource allocation and adaptation gaps for some sectors or geographies in the LMB most 15 likely exist. A common *framework* of what constitutes 'successful' adaptation initiatives and a holistic

16 transboundary climate change adaptation management framework in the LMB context does not exist to date and is 17 currently subject of an ongoing study.

18 19

20 24.9.2. Tropical Peatlands in Southeast Asia 21

22 Tropical peatlands develop only in flat lowland regions with year-round rainfall and are most extensive in SE Asia, 23 particularly on the islands of Sumatra, Borneo, and New Guinea (Posa et al., 2011). The largest areas are on coastal 24 plains and river deltas, but peatlands can also develop inland on flat or gently convex areas between rivers. They 25 eventually form dome-shaped structures less than 20 m deep that are above the local water table and fed only by 26 rainwater. The modern peatlands of SE Asia are relatively young ecosystems, having started growth between the 27 Late Glacial and Mid-Holocene, and peat accumulation appears to have ceased during the late Holocene in Central 28 Kalimantan, possibly as a result of enhanced El Niño activity (Dommain et al., 2011). In recent times these peatlands covered around 250,000 km² and contained more than 65 Gt of carbon, with two-thirds of this in Indonesia 29 30 (Page et al., 2011). Although traditionally viewed as species-poor, peat swamp forests provide an important habitat

31 for much of the region's fauna, including orangutans and a high diversity of specialized freshwater fish (Posa *et al.*, 32 2011).

33

34 SE Asian peatland ecosystems were largely intact in 1970 but have been massively impacted over the last 20 years, 35 as a result of logging and conversion to oil palm and pulpwood (Acacia spp.) plantations (Murdiyarso et al., 2010).

36 Between 1990 and 2010, forest cover on the peatlands of Peninsular Malaysia, Sumatra and Borneo fell from 77% to

37 36%, to be replaced by industrial plantations of unknown sustainability and degraded areas covered in ferns, grasses

38 and shrubs (Miettinen et al., 2011a). Draining the peat leads to shrinkage and microbial decomposition, and makes

39 the peat itself highly flammable, so the degraded peatlands have become globally significant carbon sources,

40 particular during ENSO-associated droughts (Miettinen et al., 2011b; Page et al., 2011). Pressures for peatland

41 conversion continue despite these concerns. Climate change projections suggest that many peatland areas in SE Asia 42 will experience reduced rainfall and increased seasonality over the coming decades (IPCC, 2007), leading to lower

43 water tables, enhanced peat decomposition, and greater susceptibility to fire (Page et al., 2011). On the other hand,

44 the exceptionally high carbon content makes tropical peatlands a very attractive target for greenhouse gas mitigation 45 projects involving the restoration of groundwater levels (Jaenicke, 2011).

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48 24.9.3. Glaciers of Central Asia and Siberia

50 The Altai, Pamir, and Tien Shan glaciers represent significant part of the Asian alpine cryosphere supplying up to

51 40% of water to the Aral, Balkhash and Issik Kul Lakes, and Ob and Tarim rivers (Aizen et al., 1995; 1998). All

rivers, except the Ob discharge water to central Asian arid endorheic basins populated with over 150 million people 52

53 from Turkmenistan, Afghanistan, Uzbekistan, Tajikistan, Kyrgyzstan, Kazakhstan, Mongolia and Xinjiang and other 1 north-western provinces of China, and Russia. In the last 50 years (1960-2009), central Asian glaciers lost on 2 average 10% of their area and 15% of their ice volume.

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4 The rate of glacier area change varies. Accelerated glacier ice melt increases total river runoff in heavy glacierized 5 basins by 8% (Aizen and Aizen, 2012a). The glaciers of the Altai-Sayan mountains are located in the most northern 6 periphery of the Central Asia mountain system at a south edge of the Arctic basin in Siberia (see Table 24-12 and 7 Figure 24-4). Altai-Sayan glaciers lost 14% area on average. The accelerated glacier melt and glacier area reduction in the Altai-Sayan was caused mainly by an increase of summer air temperatures by 1.03°C for the last 50 years 8 9 (Surazakov et al., 2007; Shahgedanova et al., 2010; Aizen et al., 2012b). The elevation of glaciers in the Pamir 10 mountains reaches 7,700 m a.s.l. (Muztagata-Kongur glacierized massifs). Pamir glaciers nourish the Amu Dariya 11 River, the major Aral Sea water stream. During the last 50 years (1960-2009), the largest glacier area losses (up to 12 15%) have been observed in the western and south-western Pamir and the smallest in central and eastern Pamir (3-13 5%) (Khromova et al., 2006; Aizen et al., 2012c). The Fedchenko Glacier in central Pamir, which is the world's largest alpine glacier outside of the Polar regions (72 km long, 714 km² area, and 900 m max ice thickness), 14 15 retreated 755 m between 1958 and 2009, losing only 2 km². The Tien Shan glaciers are located in the largest 16 mountain system in central Asia, stretching 2000 km from west to east. The Tien Shan glaciers are the major sources 17 of water for Balkhash and Issik Kul lakes, and the Sir Darýa and Tarim rivers. Summer precipitation decreased by 18 10% and the Tien Shan glaciers lost 8.5% of their total area on average during the last 50 years. The largest glacier 19 area lost is observed in the northern and western Tien Shan (14.3%) due to a decrease in annual precipitation (20 20 mm) at elevations above 3,000 m a.s.l. and increased air temperatures by 0.44°C. Smaller glacier recessions have 21 been observed in the inner and central Tien Shan (10% and 5% respectively). In central Tien Shan glacier recession 22 is minimal due to high-elevated accumulation areas (up to 7,000 m a.s.l.). Thus, the central Tien Shan and Pamir 23 glaciers have been revealed as more stable glaciers to climate changes in central Asia (Aizen and Aizen, 2012a; 24 Bamber, 2012; Jacob, et al., 2012). The eastern Tien Shan lost 12% of the total glacier area. On average, air

- 25 temperatures increased by 0.8°C and precipitation decrease by 7% at the equilibrium line altitude (ELA) between 26 the 1960s and 2009 in Tien Shan (Aizen and Aizen, 2012d).
- 27

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28 [INSERT TABLE 24-12 HERE

29 Table 24-12: Location and major characteristics of central Asia glaciations.]

31 **[INSERT FIGURE 24-4 HERE**

32 Figure 24-4: The difference in losses of glacier area in Altai-Sayan, Pamir and Tien Shan determined by location of

33 the mountain ridges in relation to major atmospheric moisture flow and by elevation above sea-level. Remote

- 34 sensing data analysis from 1960s (Corona) through 2009 (Landsat, ASTER and Alos Prism).]
- 35
 - Simulation models forecast that significant glacier degradation will begin when ELA has increased by 600 m
- 36 37 compared to the end of the 20th century (Aizen et al., 2007; Mitchel et al., 2004). Then, the area covered by central
- 38 Asian glaciers may shrink by 40% and the glacier volume by 60% of the current state. The IPCC scenarios predict,
- 39 on average, an increase in summer air temperature of 2°C to 8°C (about 4°C) and an increase in magnitude of
- 40 precipitation of 0.84-1.24 (about 1.1 times) (Mitchel et al., 2004). If the air temperature increases to the greatest
- 41 predicted value, i.e. by 8°C, and precipitation increases to its maximum predicted value, i.e. by 1.24 times the
- 42 current rate, then the model predicts a 970 m increase in ELA and the number of Tien Shan glaciers, glacier covered
- 43 areas, and glacier volume are predicted to shrink correspondingly by 94%, 69%, and 75% of the current state.
- 44 However, under the threshold predicted conditions, if air temperature increases by 8°C and precipitation decreases to
- 45 the minimum predicted value, i.e. by 0.84 times the current rate, then current glaciations will disappear (Aizen et al.,
- 46 2007). During the last 12,000 years, the warmest period was in the Holocene Climatic Optimum (Thermal
- 47 Maximum, circa 7,500-7,600BP), when mean air temperature was 4.2°C higher than modern, i.e. the annual average
- 48 temperature in the last three decades. Nevertheless, central Asian glaciers were able to survive during the Thermal
- 49 Maximum. Thus, for complete glacier disappearance mean air temperature should be at least 5°C higher than
- 50 modern (Aizen et al., 2012e).
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- 52 53

24.9.4. Is the Aral Sea Dying?

The Aral Sea (see Figure 24-5) was a very large sea (lake) in Central Asia that was number four (in area) in the list of the world's lakes before the 1960s (Letolle, 2008; Kostianoy and Kosarev, 2010). It is located in the Karakum and Kyzylkum deserts. Navigation and the fishery (yearly catches of 44,000 tons) were developed there. The deltas of two major rivers of Central Asia, the Amudarya and the Syrdarya, that bring waters to the Aral Sea, were known for their fisheries, biodiversity, reed production, and muskrat rearing. The local population used to work in water infrastructure related spheres (Nihoul *et al.*, 2002; Zonn *et al.*, 2009).

9 10 [INSERT FIGURE 24-5 HERE

11 Figure 24-5: The satellite view of the Aral Sea acquired on 18 August 2008 from MODIS-Terra. Image courtesy by

- 12 A.G. Kostianoy (P.P. Shirshov Institute of Oceanology, Moscow, Russia) and D.M. Solovyov (Marine
- 13 Hydrophysical Institute, Sevastopol, the Ukraine), based on the LAADS Web, NASA-Goddard Space Flight Center
- data (http://ladsweb.nascom.nasa.gov/). The red line indicates the Aral Sea coastline back in 1960. The yellow line
 indicates the border between Kazakhstan and Uzbekistan.]
- 16

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17 Since 1960, the water resources of the Amudarya and Syrdarya rivers have been irrationally used in order to increase

- 18 irrigation of agricultural lands as well as to create artificial water reservoirs (Glantz, 1999; Kostianoy and Kosarev,
- 19 2010). Hence the water balance of the Aral Sea was disrupted, and irreversible changes in the regime of the sea
- 20 occurred which later led to one of the "largest ecological disasters of the twentieth century" (Letolle and Mainguet,
- 21 1993; Glantz, 1999; Micklin and Williams, 1996). For the last fifty years we have been observing a progressive
- desiccation of the Aral Sea and deterioration of its environment. During those years the sea surface shrunk from
- 23 $66,100 \text{ km}^2$ (1961) to 10,400 km² (2008); the sea volume decreased to 110 km³ from 1,066 km³ (1961); the sea level
- fell by 24 m (in 1961 the maximum depth was 69 m); and its salinity (mineralization) increased from 10 to 116 p.p.t.
- in the western part and to 210 p.p.t. in the eastern part of the Large Aral Sea (Kostianoy and Wiseman, 2004;
 Zavialov, 2005; Kostianoy and Kosarev, 2010).
- 20

The ongoing Aral Sea desiccation and salinization have resulted in critical changes in its shape, physical and chemical state, and biodiversity. The Aral Sea related economic spheres lost their importance. The consequences of the sea degradation represent a big threat to the quickly growing population in the Priaralie (from 14 million people

- in 1960 to 45 million people in 2006) due to such factors as water quality loss, lack of fresh water, dust and salt
- 32 storms, salinization of soils, various diseases, and regional climate change (Kostianoy and Kosarev, 2010).
- 33

34 Irrational use of waters of Amudarya and Syrdarya is not the only reason for the Aral Sea desiccation. Regional

- 35 climate change (decrease in atmospheric precipitation and increase in air temperature) also seems to play a
- 36 significant role in this process. Assessments of the water amount precipitated over the Amudarya catchment area for
- the period between 1979 and 2001 showed critical decrease from about 7,5 to 4,5 km³ per month on average (Nezlin
- *et al.*, 2004). According to estimates of the IPCC AR4, the rise of the mean annual air temperature in the Aral region
- in 1960–2000 was 1°C (Lioubimtseva and Henebry, 2009). Thus, regional climate change significantly influenced
- 40 the water balance of the Aral Sea in the past 30 years leading to its "supplementary" desiccation in addition to
- 41 irrational water use.
- 42

By 2012, the main progress in saving the Aral Sea was achieved only in the Kazakh part of the sea with the Kokaral dam construction between the eastern part of the Large Aral Sea and the Small Aral Sea in August 2005 (Kostianoy and Kosarev, 2010). Today, the Small Aral Sea is slowly reviving and small fishery production is growing, while the Large Aral Sea keeps on disappearing. Since 2010 the former eastern part of the Large Aral Sea has been a wetland which is periodically filled with snowmelt and rain water and partly desiccated in dry seasons. The western part of the Large Aral Sea, being a relatively narrow and deep lake, may slowly die in the absence of external water supply (Kostianoy and Kosarev, 2010; Micklin, 2010; Breckle *et al.*, 2012; Kostianoy, 2012).

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1 Frequently Asked Questions

3 FAQ 24.1: Since AR4, what is new in our knowledge about the changing climate in Asia?

The observed increasing trend of annual mean temperature of between less than 1°C to 3°C per century and warming in daily temperature extremes has been confirmed in many countries of Asia. The warming trend is projected to continue during the 21st century across the region irrespective of stabilization scenarios. Observed trends of annual and heavy precipitation are varied throughout Asia. Variability of trends in average and extreme precipitation is projected to be wider within the region.

9

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FAQ 24.2: How will the projected impact of climate change on freshwater resources by the 2050s affect natural ecosystems and society?

Shrinking of glaciers in Central Asia and the Himalayas is projected to affect water resources positively in the near future but negatively in the long term perspective. Changes in river flow will impact natural habitats and species that are sensitive to flow extremes. Freshwater availability in Central, South, East and South-East Asia, particularly in large river basins, is projected to decrease due to climate change. Water scarcity is expected to be a big challenge in these regions. Population growth and increasing demand arising from higher standards of living, could adversely affect more than 1 billion people. Better water management strategies are needed to ease water scarcity. Water

- 18 saving technologies and changing of crops into drought tolerant crops are found to be successful adaptation options
- 19 in the region.

20

21 FAQ 24.3: How will climate change affect food production and food security in Asia?

Climate change impacts on crop production would be generally negative in many regions. For rice, most models show that higher temperatures will lead to lower rice yields as a result of shorter growing period. However, with CO₂ fertilization effect, rice yield could increase with climate change. This is also generally true for other crops. The impacts of climate change on food production and food security will vary within regions and countries- increasing yields for some areas (eg. cereal production in north and east Kazakhstan) and declining yields in others (eg. wheat in the Indo-Gangetic Plain of South Asia). There are many potential adaptation strategies such as crop breeding but research on their effectiveness is limited.

29

30 FAQ 24.4: Who are the people most at risk in Asia from climate change?

People living in low lying coastal zones and flood plains are most at risk from climate extremes and disasters in Asia. Such areas are home to 50% of Asia's urban population. Asia has more than 90% of the global population exposed to tropical cyclones. Settlements on unstable slopes or landslide prone areas face increased likelihood of rainfall induced landslides. Rural poverty in parts of Asia could be exacerbated due to negative climate change impacts on the rice crop and increase in food price and cost of living. More frequent and intense heat-waves in Asia will increase mortality and morbidity in vulnerable groups, particularly in urban environments. Urban population growth will lead to urban land-use and land-cover changes and in turn will have considerable impacts on climate.

38

39 FAQ 24.5: How will climate change affect human health in different parts of Asia?

40 More frequent and intense heatwaves will increase mortality and morbidity in vulnerable groups in urban areas. The 41 transmission of infectious disease will be affected due to changes in air and water temperatures (such as cholera 42 epidemics in coastal Bangladesh, and schistosomiasis in inland lakes in China) and altered rain patterns and water 43 flows (e.g., affecting diarrheal outbreaks in rural children). Changes in the geographical distribution of vector-borne 44 diseases will be most noted close to their distribution limits. Outbreaks of the vaccine-preventable Japanese 45 encephalitis in the Himalayan region and malaria in India and Nepal have been linked to rainfall. Cross-sector 46 collaborations are required to develop adaptive measures, involving the health sector and disaster preparedness 47 programs, water management, sanitation, urban planning, food industry and the animal health sector.

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49 FAQ 24.6: What are the challenges in climate impacts, vulnerabilities and adaptation research in Asia?

50 Gaps in data are a major challenge for Asia. For example, trends in precipitation are less available than data on

- 51 trends in temperature, data on observed climate change and changes in extreme climate events does not cover most
- 52 Asian regions. For freshwater resources, new models of future rainfall changes, developing of water managing
- 53 strategies and study on water saving technologies are needed. Biodiversity data and data on biome boundaries shift
- are incomplete, and long-term monitoring, especially in protected areas is needed to fill these gaps. Studies on

agricultural sector and appropriate adaptation options, on social-economic and institutional dimension, on urban settlements and industry should also be given priority.

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- 7

Sub-region	Count	ries/regions
Central Asia (5)	KazakhstanKyrgyzstanTajikistan	TurkmenistanUzbekistan
East Asia (7)	 China, Hong Kong Special Administrative Region China, Macao Special Administrative Region Japan 	 North Korea People's Republic of China South Korea Taiwan Province of China
North Asia (2)	Mongolia	Russia (East of Ural)
South Asia (8)	 Afghanistan Bangladesh Bhutan India 	 Maldives Nepal Pakistan Sri Lanka
South East Asia (12)	 Brunei Indonesia Lao People's Democratic Malaysia Myanmar 	 The Philippines Republic Cambodia Singapore Thailand Timor-Leste
West Asia (17)	 Papua New Guinea Armenia Azerbaijan Bahrain Georgia Iran Iraq Israel Jordan 	 Vietnam Kuwait Lebanon Occupied Palestinian Territory Oman Qatar Saudi Arabia Syria United Arab Emirates Yemen

Table 24-1: The 51		.1 . 1 .	C A '
$13 \text{ me} / 4_{-1} \cdot 1 \text{ me} 1$	countries/regions if	n the six sub-regior	IS OT AS19
1000 27 1.10001	countries/regions in	1 110 517 540 102101	15 OI 7 1510.

Table 24-2: Summary of key observed past and present climate trends and variability.

Region	Countries	Parameter	Unit	Change	Base year	Period	Reference
-	Mongolia	AMT	°C	+2.14	N/A	1940-2005	Dagvadorj et al., 2009
	C	AMP	mm/y	-0.1 ~ -2.0	N/A	1940-2005	Dagvadorj et al., 2009
	*NW Khentey	AMT	°C	+2.5	N/A	1961-2007	Dulamsuren et al., 2010
	*SW Khentey	AMT	°C	+4.4	N/A	1950-2007	Dulamsuren et al., 2010
	*SW Khentey (winter)	AMT	°C	+7.1	N/A	1950-2007	Dulamsuren et al., 2010
	*E Khengay	AMT	°C	+1.8	N/A	1937-2007	Dulamsuren et al., 2010
North Asia		AMT	°C	+1.2	N/A	1942-2007	Dulamsuren et al., 2010
(2)	*NW Khentey	AMP	mm	-100	N/A	1961-2007	Dulamsuren et al., 2010
. ,	*SW Khentey	AMP	mm	+15	N/A	1950-2007	Dulamsuren et al., 2010
	*E Khengay	AMP	mm	-20	N/A	1937-2007	Dulamsuren et al., 2010
	*SE Khentey	AMP	mm	+50	N/A	1942-2007	Dulamsuren et al., 2010
	Russia	AMT	°C	+1.29	1961-1990	1907-2006	Anisimov et al., 2008
		AMT	°C	+1.33	1961-1990	1976-2006	Anisimov et al., 2008
		AMP	mm/10y	+7.2	1961-1990	1976-2006	Anisimov et al., 2008
	China, Hong Kong Special	AMT	°C/10y	+0.12	N/A	1885-2008	Ginn et al., 2010
	Administrative Region	AMT	°C/10y	+0.16	N/A	1947-2008	Ginn et al., 2010
	A commission of the second	AMT	°C/10y	+0.27	N/A	1979-2008	Ginn et al., 2010
		Tmin.	°C/10y		N/A	1947-2008	Ginn et al., 2010
		AMP	mm/10y		N/A	1885-2008	Ginn et al., 2010
		AMP	mm/10y	not signific:		1947-2008	Ginn et al., 2010
	Japan	AMT	°C/100y		1971-2000	1898-2010	JMA, 2011
	Japan	AMP	C/1003	No clear		1070 2010	MEXT et al., 2009
	*Tokyo	AMT	°C	+2.93 (+0.24/10y)	N/A	1876-2000	Schaefer and Domroes, 2009
	*Tokyo	AMT	°C	+2.95 (+0.30/10y)	N/A	1901-2000	Schaefer and Domroes, 2009
	*Hakodate	AMT	°C	+0.35 (+0.04/10y)	N/A	1901-2000	Schaefer and Domroes, 2009
		AMT	°C	+0.33 (+0.04/10y) +2.14 (+0.44°C/10y)	N/A N/A	1901-2000	Schaefer and Domroes, 2009
	*Okayama *Hiroshima	AMT	°C	+2.35 (+0.98/10y)	N/A	1976-2000	Schaefer and Domroes, 2009
	*North Japan (March/April)	AMT	°C/y	$+0.047 \sim +0.0771$	N/A	1977-2004	Fujisawa and Kobayashi, 2010
	People's Republic of China	AMT		0.09±0.017	1971-2000	1977-2004	Li et al., 2010
	reopie's Republic of Chillia	AMT		0.26±0.032	1971-2000	1954-2006	Li et al., 2010
East Asia		AMT		0.45±0.13	1971-2000	1979-2006	Li et al., 2010
(7)		AMT	°C/10y	+0.22	N/A	1979-2000	Ren et al., 2005
		AMT	°C/10y	+0.22	N/A N/A	1951-2001	Ren et al., 2003
			°C		N/A N/A		
	*North China	AMT	-	+1.16 (+0.29/10y)		1961-2000	Ren et al., 2008
	*North China	AMT	°C/10y	+0.18 (adjusted UHI)	N/A N/A	1961-2000	Ren et al., 2008
	*North China (winter)	AMT	°C	+2.48 (+0.62/10y)	N/A	1961-2000	Ren et al., 2008
	*North China (winter)	AMT	°C/10y	+0.5 (adjusted UHI)	1051 0000	1961-2000	Ren et al., 2008
	South Korea	AMT	°C	+1.87	1971-2000	1908-2008	Kim et al., 2010
		AMT	°C	+1.37	1971-2000	1954-2008	Kim et al., 2010
		AMT	°C	+1.44	1971-2000	1969-2008	Kim et al., 2010
		AMP	%	+5.6	1973-1980	2001-2008	Kim et al., 2010
	Taiwan Province of China	AMT	°C/10y	+0.14	1980-1999	1911-2009	Hsu et al., 2011
		AMT	°C/10y	+0.19	1980-1999	1959-2009	Hsu et al., 2011
		AMT	°C/10y		1980-1999	1979-2009	Hsu et al., 2011
		DP≧0.1mm			1980-1999	1911-2009	Hsu et al., 2011
		$DP \ge 0.1 mm$			1980-1999	1959-2009	Hsu et al., 2011
		DP≧0.1mm	days/10y		1980-1999	1979-2009	Hsu et al., 2011
	Indonesia *Brontas Catchment	MMP	mm/m	-0.12 ~ -2.12	N/A	1955-2005	Aldrian and Djamil, 2008
	*Brontas Catchment	AMP	mm/y	-1.23 ~ -24.25	N/A	1955-2005	Aldrian and Djamil, 2008
outh East	The Philippines	AMT	°C	+0.648	1971-2000	1951-2010	PAGASA, 2011
Asia (12)		AMT	°C/y	+0.0108	1971-2000	1951-2010	PAGASA, 2011
		Tmax.	°C	+0.36	1971-2000	1951-2010	PAGASA, 2011
	1	Tmin.	°C	+1	1971-2000	1951-2010	PAGASA, 2011

Region	Countries	Parameter	Unit	Change	Base year	Period	Reference
	Afghanistan	AMT	°C	+0.6	N/A	1960-2008	Savage et al., 2009
	-	AMT	°C/10y	+0.13	N/A	1960-2008	Savage et al., 2009
		AMP	mm/m	-0.5	N/A	1960-2008	Savage et al., 2009
		AMP	%/10y	-2	N/A	1960-2008	Savage et al., 2009
	Bangladesh	AMT	°C/10y	+0.097	N/A	1958-2007	Shahid, 2010
	5	AMP	mm/y	+5.53	N/A	1958-2007	Shahid, 2010
	*Western Bangladesh	AMP	mm/y	+6.97 ~ +7.79	N/A	1958-2007	Shahid, 2010
	*Northern Bangladesh	AMP	mm/y	+14.39 ~ +16.45	N/A	1958-2007	Shahid, 2010
	India	AMT	°C	+0.56	1961-1990	1901-2009	Attri and Tyagi, 2010
		Tmax.	°C	+1.02	1961-1990	1901-2009	Attri and Tyagi, 2010
		T _{min.}	°C	+0.12	1961-1990	1901-2009	Attri and Tyagi, 2010
		AMP	C	No significant national t		1901-2009	Attri and Tyagi, 2010
		AMI	°C/100y	<u> </u>	N/A	1880-2000	Lal, 2003
		AMT		+0.0056°	N/A N/A	1948-2008	Ganguly, 2011
	Namal	AMT	C/y °C/y				Shrestha et al., 1999
	Nepal	_	°C	+0.06 +0.57	N/A 1961-1990	1977-1994 1901-2000	Chaudhry et al., 2009
Carrella A alla	Pakistan	AMT		+0.099			
South Asia		AMT	°C/10y		1961-2000	1960-2007	Chaudhry et al., 2009
(8)		AMT	°C	+0.47±0.21	1961-2000	1960-2007	Chaudhry et al., 2009
		Tmax.		+0.18	N/A	1960-2007	Chaudhry et al., 2009
		Tmax.	°C	+0.87±0.26	N/A	1960-2007	Chaudhry et al., 2009
		Tmin.	°C/10y	+0.1	N/A	1960-2007	Chaudhry et al., 2009
		Tmin.	°C	+0.48±0.2	N/A	1960-2007	Chaudhry et al., 2009
		AMP	mm	+61	N/A	1901-2007	Chaudhry et al., 2009
		AMP	mm	-156	N/A	1901-1954	Chaudhry et al., 2009
		AMP	mm	+35	N/A	1955-2007	Chaudhry et al., 2009
	*Karachi	AMT	°C	+2.25 (+0.38/10y)	N/A	1947-2005	Sajjad et al., 2009
	*Upper Indus River basin	AMT	°C	+1.79	N/A	1967-2005	Khattak et al., 2011
	*Middle Indus River basin	AMT	°C	+1.66	N/A	1967-2005	Khattak et al., 2011
	*Lower Indus River basin	AMT	°C	+1.20	N/A	1967-2005	Khattak et al., 2011
	Sri Lanka	AMT	°C/y	+0.005 ~ +0.035	N/A	1961-2000	Iqbal, 2010
		AMP	mm/y	-1.55 ~ -19.06	N/A	1961-2000	Iqbal, 2010
		AMT	°C/10y	+0.3 ~ +0.93	N/A	1869-2007	De Costa, 2008
		AMT	°C/10y	+0.75 ~ +0.94	N/A	1910-2007	De Costa, 2008
	*Four of 7 study areas	AMP	mm/y	-0.28 ~ -0.84	N/A	1869-2007	De Costa, 2008
West Asia		AMT	°C	+0.85	1961-1990	1935-2007	Gabrielyan et al., 2010
(17)		AMP	%	-6	1961-1990	1935-2007	Gabrielyan et al., 2010
()	*General	AMT	°C	+1 ~ +2	N/A	1880-2000	Lioubimtseva et al., 2005
	Kazakhstan	AMT	°C/10y	+0.31	N/A	1936-2005	Kryukova et al., 2009
		AMP	C, 10y	No definite national tre		1936-2005	Kryukova et al., 2009
	Kyrgyzstan	AMT	°C	+1.6	N/A	1901-2000	Iliasov et al., 2003
	ixyigyzstan	AMP	mm	+1.0 +23mm (+6%)	N/A N/A	1901-2000	Iliasov et al., 2003
Central	Tajikistan *plain region	AMP	°C/10y	+0.1 ~ +0.2	N/A N/A	1901-2000	Karimov et al., 2008
Asia		AMT	°C	$+0.1 \sim +0.2$ $+0.3 \sim +0.5$	N/A	1940-2005	Karimov et al., 2008
(5)	*mountainos region	AMP	%	$+0.3 \approx +0.3$ +8 (insignificant)	N/A N/A	1940-2005	Karimov et al., 2008
	*up to 2500 masl	AMP		-3 (insignificant)	N/A N/A	1940-2005 1940-2005	Karimov et al., 2008 Karimov et al., 2008
	*mountainous areas						
	Turkmenistan	AMT	°C/10y		N/A	1931-1995	MNPT, 2000
	TT 1.1 '	AMP	mm/10y		N/A	1931-1995	MNPT, 2000
	Uzbekistan	Tmax.	°C/10y		N/A	1951-2008	Uzhydromet, 2008
		Tmin.	°C/10y		N/A	1951-2008	Uzhydromet, 2008
Tibetan Pla	teau	AMT	°C	+1.8 (0.36/10y)	N/A	1961-2007	Wang et al., 2008
		AMT	,	+0.447	N/A	1962-2001	Xu et al., 2008
		AMP	mm/y	+0.614	N/A	1961-2001	Xu et al., 2008

Region	Countries	Key trend	Period	Reference
Temperat	ure Extreme	98		
North Asia	*SREX	Likely increases in warm days/nights and likely decreases in cold days/nights		SREX, Ch.3, Table 3.2
		Spatially varying trends in warm spells, overall increase in warm spell duration index		SREX, Ch.3, Table 3.2
		(WSDI);WSDI decrease in some areas		
	Mongolia	Decrease in warm day-times and nights	1948-2006	Fang et al., 2008
	Russia	Increase in warm daytimes and nights in northeastern Siberia	1948-2006	Fang et al., 2008
East Asia	*SREX	Likely increases in warm days and likely decreases in cold days		SREX, Ch.3, Table 3.2
		Decreases in cold nights and increases in warm nights		SREX, Ch.3, Table 3.2
		Increase in warm season heat waves in China		SREX, Ch.3, Table 3.2
		Increase in WSDI in North China, but decline in South China		SREX, Ch.3, Table 3.2
	People's	Increasing frequency and severity of regional wet heatwaves events with a magnitude	1960-2008	Ding and Qian, 2011
	Republic of	of 0.29 times per decade		
	China	Extreme warm-month events have strong spatial dependance, with smaller variability	1960-2007	Wan, 2009
		over the Tibetan Plateau, North China plain and coastal areas of South China, and		
		larger variability over North China		
		Significant decrease in warm daytimes and nights in North China	1948-2006	Fang et al., 2008
	South Korea	20 heatwaves with mean annual duration of 9.3 days (longest being 33 days);	1991-2005	Kysely and Kim, 2009
		Mean relative excess total mortality shows a positive trend of +5.9%;		
		Cardiovascular disease mortality shows a positive trend of +9%		
S-E Asia	*SREX	Increases in warm days, decreases in cold nights		SREX, Ch.3, Table 3.2
		Decreases in cold days, increases in warm nights in the northern part of domain		SREX, Ch.3, Table 3.2
	*General	Significant increase in warm day-times and nights in inland and on the coast	1948-2006	Fang et al., 2008
	Malaysia	Significant increase in warm nights	1948-2006	Fang et al., 2008
South Asia	*SREX	Increase in warm days/nights and decrease in cold days/nights		SREX, Ch.3, Table 3.2
	*General	Increase in warm daytimes and nights	1948-2006	Fang et al., 2008
	Afghanistan	Decrease in warm daytimes	1948-2006	Fang et al., 2008
	Pakistan	Decrease in warm daytimes	1948-2006	Fang et al., 2008
West Asia	*SREX	More likely than not decrease in cold days and a very likely increase in warm days		SREX, Ch.3, Table 3.2
		Likely decrease in cold nights and likely increase in warm nights		SREX, Ch.3, Table 3.2
		WSDI increase		SREX, Ch.3, Table 3.2
	*General	Increase in warm daytimes and nights	1948-2006	Fang et al., 2008
Central Asia	*SREX	Likely increases in warm days/nights and likely decreases in cold days/nights		SREX, Ch.3, Table 3.2

Table 24-3: Summary of observed changes in extreme events and severe climate anomalies.

Region	Countries	Key trend	Period	Reference						
leavy pre	cipitation									
North Asia	*SREX	Increase in some regions, but spatial variations		SREX, Ch.3, Table 3.2						
		Some increase western Russia, especially in winter	1950-2000	SREX, Ch.3, Table 3.2						
	Russia	-4 to +4 days in absolute terms, or -40% to +40% in relative terms	1936-2000	Bogdanova et al., 2010						
		In the western part, areas that show increase considerably exceed areas of decrease								
		In the eastern part, speed of the increase is lower, and the speed of decrease is higher								
		than in the western part								
East Asia	*SREX	Spatially varying trends in heavy precipitation								
	Japan	+2.49%/decade for \geq 100mm precipitation days	1901-2004	Fujibe et al., 2006						
		Increased heavy precipitation mainly in West Japan and in autumn, although weak								
		positive trends are found in most other regions and seasons								
		Trend in annual maximum number of heavy daily precipitation indices is								
		+0.89%/decade for whole territory of Japan								
		+4.2%/decade for \geq 200mm/day	1901-2006	Fujibe, 2008						
		+2.4%/decade for \geq 100mm/day								
		-0.9%/decade to -1.5%/decade for \geq 1mm/day								
		+63.2% ±52.2%/decade for ≥300mm/6h	1979-2007							
		$+37.6\% \pm 30.4\%$ /decade for ≥ 200 mm/6h								
		$+48.4\% \pm 45\%$ /decade for ≥ 100 mm/h								
	People's	Increases in >50mm/day, and/or heavy (25-50mm/day) precipitation in SE China	1978-2002	Yao et al., 2008;						
	Republic of	Sudden increase in severe floods in Poyang Lake	during past	Shankman et al., 2006						
	China		few decades							
		All of the severest floods since 1950 occurred during or immediately following El Niño events	since 1950							
	South Korea	A gradual increase in heavy summer precipitation days (≥30mm/day) around mid-late	1954-2001	Ho et al., 2003						
	Soull Horea	1970s	1901 2001	110 01 41, 2000						
		Significant increasing trends for indices measuring heavy precipitation frequency and	1971-2000	Im et al., 2008						
		intensity	1971 2000	1111 et uii, 2000						
		Pronounced enhancement of the number of days with precipitation above 80mm		Im et al., 2011						
		intensity, percentage of total rainfall from events above longterm 95th perceltile, and		1111 of uil, 2011						
		greatest 10-day total precipitation in southern parts								
-E Asia	*SREX	Spatially varying trends in heavy precipitation								
LITION	Malaysia	Decreasing trend in frequency of daily rainfall exceeding the 1971-2005 mean 99th	1971-2005	Zin et al., 2010						
	Willingshi	percentile (days) at 60% of stations	1971 2005	2.11 01 01., 2010						
		Increasing trends in wet day intensities greater or equal to 95th and 99th percentile are								
		+5.08mm/decade and +8.75mm/decade respectively (Petaling Java)								
		Increasing trends in wet day intensities greater or equal to 95th and 99th percentile are								
		+3.41mm/decade and +5.57mm/decade respectively (Subarg)								
outh Asia	*SREX	Mixed signal in India								
outinnista	India	+6%/decade for ≥150mm/day	1901-2004	Rajeevan et al., 2008						
	man	+14.5%/decade for \geq 150mm/day	1951-2004	Rujee van et al., 2000						
		$+10\%$ /decade for \geq 100mm/day	1951-2000	Goswami et al., 2006						
Vest Asia	*SREX	Decrease in heavy precipitation events	1991 2000	005 Wallin et al., 2000						
entral Asia		Spatially varying trends in heavy precipitation								
ropical Cy		spatially tarying denas in neury preeplation	l							
last Asia	*General	Increasing typhoon influence in subtropical East Asia and considerable decrease over	1965-2003	Wu et al., 2005						
	General	South China Sea due to changes in large-scale steering flow (tropospheric cooling in	1700 2000	11 a ot all, 2000						
		the last 20 years was suggested as cause)								
	People's	Tropical cyclone frequency shows a decreasing trend over most part of China except	1955-2007	Ying et al., 2011						
	Republic of	at some locations (low reaches of Yangtze River) where averaged number of tropical	1955 2007	1 mg ot un, 2011						
	China	cyclones over last 25 years decreased about 1-2 per year, relative to first 25 years								
S-E Asia	*General	Growing duration of the most extreme winds (tropical storms and typhoons) over	1960-2000	Rozynski et al, 2009						
L 136	General	South East Asian seas, mainly the South China Sea and the Philippine Sea	1700-2000	1002ynoki et al, 2009						
Fast China	Sea and	Significant decrease in frequency of typhoon passage in East China Sea and Philippine	1951-2001	Ho et al. 2004						
*East China Sea and Philippine Sea		Significant decrease in frequency of typnoon passage in East China Sea and Philippine Sea in the 1980-2001 period, relative to 1951-1979	1751-2001	Ho et al., 2004						
		A continuous downward trend over Philippine Sea is found at a rate of change of -								
acific		0.6%/year, which amounted to 45% decrease over the study period	1050 2004	Chap. 2000						
Pacific		Decreasing trend in tropical cyclone number in northwestern Pacific	1959-2006	Chen, 2009						
		Trend of tropical cyclone frequency in southeastern Pacific shows an increase until								
		early 1990s and then a moderate decrease		1						
		-		eeted enanges i	r		-	-		
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egion	Countries	Paramete		Projected change	Scenario	GCM	RCM	Base year		Reference
		AMT	°C	+1.7±0.6 (1.0,2.8)	RCP2.6			1986-2005	2081-2100	WGI AR5
		AMT	°C	+2.8±0.8 (1.8,4.0)	RCP4.5			1986-2005	2081-2100	WGI AR5
	Asia	AMT	°C	+3.5±0.9 (2.3,4.9)	RCP6.0			1986-2005	2081-2100	WGI AR5
	71510	AMT	°C	+5.6±1.3 (3.7,7.8)	RCP8.5			1986-2005	2081-2100	WGI AR5
		AMT	°C	large increases	SRES A2	MRI-CGCM2.2	CRIEPI-RegCM3	1981-1990	2046-2055	Takayabu et al., 2007
		AMT	°C	large increases	SRES A2	MRI-CGCM2.2	MRI-RCM	1981-1990	2046-2055	Takayabu et al., 2007
	*General	AMP	%	+15 ~ +25	SRES A1B	*1		1981-2000	2081-2100	Kim and Byun, 2009
	Mongolia	AMP	mm/y	increase	SRES A2	MRI-CGCM2.2	CRIEPI-RegCM3	1979-2005	2046-2055	Takayabu et al., 2007
		AMP	mm/y	increase	SRES A2	MRI-CGCM2.2	MRI-RCM	1979-2005	2046-2055	Takayabu et al., 2007
	*intensity of	AMT	°C	+1	SRES A2	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
	warming in summer	AMT	°C	+2.7	SRES A2	HadCM3		1980-1999	2046-2065	Dagvadorj et al., 2009
	season is higher	AMT	°C	+5	SRES A2	HadCM3		1980-1999	2080-2099	Dagvadorj et al., 2009
	than winter	AMT	°C	+0.9	SRES A1B	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
		AMT	°Č	+3	SRES A1B	HadCM3		1980-1999	2046-2065	Dagvadorj et al., 2009
	**increase in	AMT	°Č	+4.6	SRES A1B	HadCM3		1980-1999	2080-2099	Dagvadorj et al., 2009
	summer	AMT	°C	+0.8	SRES B1	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
rth Asia	precipitation will be	AMT	°C	+2.1	SRES B1	HadCM3		1980-1999	2046-2065	Dagvadorj et al., 2009
(2)	smaller than winter	AMT	°C	+3.1	SRES B1	HadCM3		1980-1999	2080-2099	Dagvadorj et al., 2009
	precipitation	AMP	%	+2	SRES A2	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
	precipitation	AMP	70 %	+2	SRES A2 SRES A2	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
		AMP	%	+15	SRES A2	HadCM3		1980-1999	2080-2099	Dagvadorj et al., 2009
		AMP	%	0	SRES A1B	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
		AMP	%	+7	SRES A1B	HadCM3		1980-1999	2046-2065	Dagvadorj et al., 2009
		AMP	%	+16	SRES A1B	HadCM3		1980-1999	2080-2099	Dagvadorj et al., 2009
		AMP	%	+3	SRES B1	HadCM3		1980-1999	2011-2030	Dagvadorj et al., 2009
		AMP	%	+6	SRES B1	HadCM3		1980-1999	2046-2065	Dagvadorj et al., 2009
		AMP	%	+11	SRES B1	HadCM3		1980-1999	2080-2099	Dagvadorj et al., 2009
	*General	AMP	%	+5 ~ +15	SRES A1B	*1		1981-2000	2081-2100	Kim and Byun, 2009
	China, Hong Kong	AMT	°C	+3	"low-end"			1980-1999	2090-2099	Ginn et al., 2010
	Special	AMT	°C	+6.8	"high-end"			1980-1999	2090-2099	Ginn et al., 2010
	Administrative	AMT	°C	+4.8	"middle-of-the-road"			1980-1999	2090-2099	Ginn et al., 2010
	Region	AMP	%	+11	above three			1980-1999	2090-2099	Ginn et al., 2010
	People's Republic	AMP	mm/y	increase in N. China	SRES A2	MRI-CGCM2.2	CRIEPI-RegCM3	1979-2005	2046-2055	Takayabu et al., 2007
	of China	AMP	mm/y	increase in N. China	SRES A2	MRI-CGCM2.2	MRI-RCM	1979-2005	2046-2055	Takayabu et al., 2007
	or crima	HWfreq.	times/y	increase (max. >5)	SRES A2	HadAM3P/HadCM3	PRECIS	1961-1990	2071-2100	Yang et al., 2010
		HW neq.	days	7 to 14 (mostly 9)	SRES A2	HadAM3P/HadCM3	PRECIS	1961-1990	2071-2100	Yang et al., 2010
st Asia				+3.5(RCM)~+3.7(GCM)	SRES A2 SRES A2		ICTP RegCM3	1961-1990	2071-2100	Gao et al., 2010
(7)		AMT	°C			FvGCM/CCM3				
	a	AMP	%	+5.5(RCM)~+11.3(GCM)	SRES A2	FvGCM/CCM3	ICTP RegCM3	1961-1990	2071-2100	Gao et al., 2011
	South Korea	Tmax.	°C	33.2	SRES A1B	ECHAM5/MPI-OM	RegCM3	1971-2000	2001-2100	Im et al., 2011
		Tmin.	°C	-6.6	SRES A1B	ECHAM5/MPI-OM	RegCM3	1971-2000	2001-2100	Im et al., 2011
		FD	days	65.3	SRES A1B	ECHAM5/MPI-OM	RegCM3	1971-2000	2001-2100	Im et al., 2011
		HD	days	47.2	SRES A1B	ECHAM5/MPI-OM	RegCM3	1971-2000	2001-2100	Im et al., 2011
		HWdur.	days	14	SRES A1B	ECHAM5/MPI-OM	RegCM3	1971-2000	2001-2100	Im et al., 2011
		DRYmax.	days	20	SRES A1B	ECHAM5/MPI-OM	RegCM3	1971-2000	2001-2100	Im et al., 2011
	Taiwan, Province	AMT	°C	+1.7 ~ +3.4	SRES A1B	XX [in Chinese]		1980-1999	2080-2099	Hsu et al., 2011
	of China	AMP	%	-3 ~ -22	SRES A1B	XX [in Chinese]		1980-1999	2080-2099	Hsu et al., 2011
	*General	AMT _{max} .	°C	+0.5 ~ +1.5	SRES A1B	ECHAM5/MPI-OM	WRF	1990-1999	2045-2054	Chotamonsak et al., 2011
		AMTmin.	°C	+0.81 ~ +1.52	SRES A1B	ECHAM5/MPI-OM	WRF	1990-1999	2045-2054	Chotamonsak et al., 2011
		RSLR	m	+0.03	SRES A2	*DIVA, IMAGE2.2,		1995	2010	McLeod et al., 2010
		RSLR	m	+0.08 ~ +0.09	SRES A2	CLIMBER-2		1995	2030	McLeod et al., 2010
		RSLR	m	+0.16 ~ +0.17	SRES A2	CLIMBLIC 2		1995	2050	McLeod et al., 2010
		RSLR	m	+0.43 ~ +0.46	SRES A2	-		1995	2100	McLeod et al., 2010
		RSLR		+0.43 ~ +0.40		-		1995	2010	
			m		SRES B1	-				McLeod et al., 2010
		RSLR	m	+0.08 ~ +0.09	SRES B1	-		1995	2030	McLeod et al., 2010
		RSLR	m	+0.14 ~ +0.15	SRES B1	_		1995	2050	McLeod et al., 2010
		RSLR	m	+0.30 ~ +0.32	SRES B1			1995	2100	McLeod et al., 2010
	The Philippines	AMT	°C	+0.7	SRES A2	ECHAM4	PRECIS	1971-2000	2006-2035	PAGASA, 2011
		AMT	°C	+1	SRES A1B	HadCM3Q0	PRECIS	1971-2000	2006-2035	PAGASA, 2011
		AMT	°C	+0.7	SRES B2	ECHAM4	PRECIS	1971-2000	2006-2035	PAGASA, 2011
th East		AMT	°C	+1.7	SRES A2	ECHAM4	PRECIS	1971-2000	2036-2065	PAGASA, 2011
		AMT	°C	+2	SRES A1B	HadCM3Q0	PRECIS	1971-2000	2036-2065	PAGASA, 2011
Asia		AMT	°C	+1.6	SRES B2	ECHAM4	PRECIS	1971-2000	2036-2065	PAGASA, 2011
(12)		AMT	°C	+3.4	SRES A2	ECHAM4	PRECIS	1971-2000	2036-2065	PAGASA, 2011
		AMT	°C	+3.1	SRES A1B	HadCM3Q0	PRECIS	1971-2000	2036-2065	PAGASA, 2011
		AMT	°C	+2.5	SRES B2	ECHAM4	PRECIS	1971-2000	2036-2065	PAGASA, 2011
	Thailand *Ping	AMT	°C	+0.4 ~ +0.5	SRES A2	ECHAM4/OPYC3		1990-1999	2020s	Sharma et al., 2007
	River Basin	AMT	°C	+1.3 ~ +1.5	SRES A2	ECHAM4/OPYC3	1	1990-1999	2020s	Sharma et al., 2007
	- and Bush	AMT	°C	+0.3 ~ +0.4	SRES B2	ECHAM4/OPYC3	1	1990-1999	2030s 2020s	Sharma et al., 2007
		AMT	°C	+0.3 ~ +0.4 +0.9 ~ +1.1	SRES B2 SRES B2	ECHAM4/OPYC3		1990-1999 1990-1999	2020s 2050s	
	Timor last-		°C		SRES B2 SRES A1B,A2, B1		1			Sharma et al., 2007
	Timor-leste	AMT		$+0.8 (+0.4 \sim +1.5)$	SRES A1B,A2, B1	*2		1961-1990	2020s	4
		AMT	°C	+1.5 (+0.7 ~ +2.8)	SRES A1B,A2, B1 SRES A1B,A2, B1	*2	1	1961-1990	2050s	a 1 1 1 1 1 1 1 a 1 a 1
		AMT	°C	+2.2 (+0.8 ~ +4)	SRES A1B,A2, B1 SRES A1B,A2, B1	*2		1961-1990	2080s	Cardno Acil and KWK Consul
		AMP	%	+2 (-12 ~ +15)		*2	1	1961-1990	2020s	2010; Kirono 2010
		AMP	%	+4 (-25 ~ +15)	SRES A1B,A2, B1	*2		1961-1990	2050s	
		AMP	%	+6 (-21 ~ +32)	SRES A1B,A2, B1	*2		1961-1990	2080s	
		HWDI	days	+2		CSIRO-CCAM		1981-2000	2041-2060	Kirono, 2010
	*General	AMP	%	+5 ~ +10	SRES A1B	*1		1981-2000	2081-2100	Kim and Byun, 2009
	Afghanistan	AMT	°C	+1.4	SRES A1, A2, B1	15 GCM ensemble	1	1970-1999	2001 2100	Savage et al., 2009
	Britting and	AMT	°C	+2.8 ~ +5	SRES A1, A2, B1 SRES A1, A2, B1	15 GCM ensemble		1970-1999	2020	Savage et al., 2009
							1			
		AMP	mm	+10 ~ +20	SRES A1, A2, B1	15 GCM ensemble		1970-1999	2030s	Savage et al., 2009
		AMP	mm	-10 ~ -40	SRES A1, A2, B1	15 GCM ensemble		1970-1999	2090s	Savage et al., 2009
h Asia			°C/10y	+1.73	SRES A2	17-model ensemble	NCC-RCM		2011-2050	Chaudhry et al., 2009
	Pakistan	AMT							10044 00 00	Chaudhry et al., 2009
	Pakistan	AMT AMT	°C/10y	+1.26	SRES A1B	17-model ensemble	NCC-RCM		2011-2050	
	Pakistan				SRES A1B SRES B1	17-model ensemble ECHAM-5	NCC-RCM NCC-RCM		2011-2050 2011-2050	Chaudhry et al., 2009 Chaudhry et al., 2009
	Pakistan	AMT	°C/10y	+1.26						
th Asia (8)	Pakistan	AMT AMT	°C/10y °C/10y	+1.26 -0.89	SRES B1	ECHAM-5	NCC-RCM		2011-2050	Chaudhry et al., 2009

Table 24-4: Summary of projected changes for a variety of climate parameters.

Region	Countries	Paramete	Unit	Projected change	Scenario	GCM	RCM	Base year	Period	Reference
rtogion	*General	AMT	°K	+1.41±0.32	SRES A2	*3		2000-2009	2045-2054	Evans, 2009
	General	AMT	°K	+3.95±0.73	SRES A2	*3		2000-2009	2090-2099	Evans, 2009
		AMP	mm	-8.42±16.08	SRES A2	*3		2000-2009	2090-2099	Evans, 2009 Evans, 2009
		AMP	mm	-25.45±28.66	SRES A2	*3		2000-2009	2090-2099	Evans, 2009
		AMP	%	0~-25	SRES A1B	*1		1981-2000	2090-2099	Kim and Byun, 2009
	Armenia	AMT	°C	+1.1 ~ +1.2	SRES A2	1		1961-1990	2031-2100	Gabrielyan et al., 2010
	Armenia	AMT	°C	+3.2 ~ +3.4	SRES A2 SRES A2	+		1961-1990	2030	Gabrielyan et al., 2010
						ł				
		AMT	°C	+5.3 ~ +5.7	SRES A2	ł		1961-1990	2100	Gabrielyan et al., 2010
		AMT	°C	+1 ~ +1.1	SRES B2	-		1961-1990	2030	Gabrielyan et al., 2010
		AMT	°C	+2.9 ~ +3	SRES B2			1961-1990	2070	Gabrielyan et al., 2010
West Asia		AMT	°C	+4.8 ~ +5.1	SRES B2	MAGICC/SCENGEN		1961-1990	2100	Gabrielyan et al., 2010
(18)		AMP	%	-2 ~ -6	SRES A2	(combination of models)		1961-1990	2030	Gabrielyan et al., 2010
. ,		AMP	%	-6 ~ -17	SRES A2	ł		1961-1990	2070	Gabrielyan et al., 2010
		AMP	%	-10 ~ -27	SRES A2			1961-1990	2100	Gabrielyan et al., 2010
		AMP	%	-2 ~ -6	SRES B2			1961-1990	2030	Gabrielyan et al., 2010
		AMP	%	-3 ~ -15	SRES B2			1961-1990	2070	Gabrielyan et al., 2010
		AMP	%	-8 ~ -24	SRES B2			1961-1990	2100	Gabrielyan et al., 2010
		AMT	°C	+1	SRES A2		PRECIS	1961-1990	2030	Gabrielyan et al., 2010
		AMT	°C	+3	SRES A2		PRECIS	1961-1990	2070	Gabrielyan et al., 2010
		AMT	°C	+4	SRES A2		PRECIS	1961-1990	2100	Gabrielyan et al., 2010
		AMP	%	-3	SRES A2		PRECIS	1961-1990	2030	Gabrielyan et al., 2010
		AMP	%	-6	SRES A2		PRECIS	1961-1990	2070	Gabrielyan et al., 2010
		AMP	%	-9	SRES A2		PRECIS	1961-1990	2100	Gabrielyan et al., 2010
	*General	AMT	°C	+2.87 ~ +5.49	SRES A1	1		1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMT	°C	+2.68 ~ +4.55	SRES A2	İ		1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMT	°C	+1.93 ~ +2.49	SRES B1	t		1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMT	°C	+1.93 ~ +3.8	SRES B2	t		1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMT	°C	+3.99 ~ +7.17	SRES A1	+		1961-1990	2080	Lioubimtseva and Henebry, 2009
		AMT	°C	+2.87 ~ +6.42	SRES A2	+		1961-1990	2080	Lioubimiseva and Henebry, 2009
		AMT	°C	+2.49 ~ +4.74	SRES B1	HadCM3,		1961-1990	2080	Lioubimtseva and Henebry, 2009
			°C	+2.68 ~ +4.18		ECHAM4,		1961-1990		
		AMT			SRES B2	ECHAM5,			2080	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.6 ~ +0.59	SRES A1	CSIRO-Mk3, CGCM3		1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.49 ~ +0.42	SRES A2			1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.43 ~ +0.08	SRES B1	-		1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-1 ~ +1	SRES B2			1961-1990	2050	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.77 ~ +0.08	SRES A1			1961-1990	2080	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.43 ~ +0.08	SRES A2	-		1961-1990	2080	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.43 ~ -0.09	SRES B1			1961-1990	2080	Lioubimtseva and Henebry, 2009
		AMP	mm/d	-0.26 ~ +0.08	SRES B2			1961-1990	2080	Lioubimtseva and Henebry, 2009
	Kazakhstan	AMT	°C	+1.4				1961-1990	2016-2045	Kryukova et al., 2009
		AMT	°C	+2.7	Median results of			1961-1990	2036-2065	Kryukova et al., 2009
		AMT	°C	+4.6				1961-1990	2071-2100	Kryukova et al., 2009
		AMP (rain)	%	+2	SRES A1F1, A2,			1961-1990	2016-2045	Kryukova et al., 2009
		AMP (rain)	%	+4	B2, B1			1961-1990	2036-2065	Kryukova et al., 2009
		AMP (rain)	%	+5	1			1961-1990	2071-2100	Kryukova et al., 2009
		AMT	°C	+1.2 ~ +1.9	SRES A1F1	GTD 700		1961-1990	2016-2045	Kryukova et al., 2009
Central		AMT	°C	+2.5 ~ +4	SRES A1F1	CERF98,		1961-1990	2036-2065	Kryukova et al., 2009
Asia		AMT	°C	+5.7 ~ +8	SRES A1F1	CSI296,		1961-1990	2071-2100	Kryukova et al., 2009
(5)		AMP (rain)	%	-2 ~ +8	SRES A1F1	ECH498,		1961-1990	2016-2045	Kryukova et al., 2009
		AMP (rain)	%	-4 ~ +15	SRES A1F1	CSM_98,		1961-1990	2036-2065	Kryukova et al., 2009
		AMP (rain)	%	+8 ~ +28	SRES A1F1	HAD300		1961-1990	2071-2100	Kryukova et al., 2009
		AMT AMT	~°C	+0 ~ +20	SRES B1	t		1961-1990	2016-2045	Kryukova et al., 2009
		AMT	°C	+1.5 ~ +2.2 +1.6 ~ +2.6	SRES B1	ł		1961-1990	2016-2045	Kryukova et al., 2009 Kryukova et al., 2009
			°C			ł				
		AMT AMB (roin)		+3.1 ~ +3.4	SRES B1	ł		1961-1990	2071-2100	Kryukova et al., 2009
		AMP (rain)	%	0~+8	SRES B1	ł		1961-1990	2016-2045	Kryukova et al., 2009
		AMP (rain)	%	-3 ~ +9	SRES B1	ł		1961-1990	2036-2065	Kryukova et al., 2009
		AMP (rain)	%	-2 ~ +13	SRES B1			1961-1990	2071-2100	Kryukova et al., 2009
	Kyrgyzstan	AMT	°C	+4.5 ~ +8.4	A2-ASF	-		1961-1990	2100	Iliasov and Yakimov, 2009
		AMT	°C	+3.5 ~ +6.1	B2-MESSAGE	MAGICC/SCENGEN		1961-1990	2100	Iliasov and Yakimov, 2009
		AMP	%	-43.4 ~ +59.9	A2-ASF	(combination of models)		1961-1990	2100	Iliasov and Yakimov, 2009
		AMP	%	-30.9 ~ +40.9	B2-MESSAGE			1961-1990	2100	Iliasov and Yakimov, 2009
	Turkmenistan	AMT	°C	+4.6	Double CO ₂	GISS		1961-1990	2050-2100	MNPT, 2000
		AMT	°C	+4.2	Double CO ₂	CCC		1961-1990	2050-2100	MNPT, 2000
		AMT	°C	+5.5	Double CO ₂	UK89		1961-1990	2050-2100	MNPT, 2000
		AMT	°C	+6.1	Double CO ₂	GFDL		1961-1990	2050-2100	MNPT, 2000
		AMT	°C	+4.8	Double CO ₂	GFDL-T		1961-1990	2050-2100	MNPT, 2000
		AMP	%	-56	Double CO ₂	GISS		1961-1990	2050-2100	MNPT, 2000
		AMP	%	0	Double CO ₂	CCC		1961-1990	2050-2100	MNPT, 2000
		AMP	%	-17	Double CO2	UK89		1961-1990	2050-2100	MNPT, 2000
		AMP	%	-15	Double CO2	GFDL	1	1961-1990	2050-2100	MNPT, 2000
		AMP	%	-4.4	Double CO ₂ Double CO ₂	GFDL-T		1961-1990	2050-2100	MNPT, 2000
Tibetan Pla	teau	AMP			SRES A2	MRI-CGCM2.2	CRIEPI-RegCM3	1961-1990	2030-2100	Takayabu et al., 2007
i iocian Pla	wati		mm/y	increase						
		AMP	mm/y	increase	SRES A2	MRI-CGCM2.2	MRI-RCM	1979-2005	2046-2055	Takayabu et al., 2007

^{*1} CCSM3, CGCM3.1 (T47), CGCM3.1 (T63), CNRM-CM3, CSIRO-MK3.0, ECHAM5/MPI-OM, FGOALS-g1.0, GFDL-CM2.0, GFDL-CM2.1, GISS_AOM, GISS-ER, INM-CM3.0, MIROC3.2 (hiers), MIROC3.2 (medres), MRI-CGCM2.3.2

^{*&}lt;sup>2</sup> BCCR-BCM2.0, CCCMA-CGCM3, CCCMA-CGCM3_T63, CNRM-CM3, CSIRO-MK3.0, GFDL-CM2.0, GFDL-CM2.1, GISS_AOM, GISS_EH, GISS-ER, IAP_FGOALS1.0G, INMCM30, IPSL_CM4, MIROC3.2_HIRES, MIROC3.2_MEDRES, MIUB-ECHOG, MPI-ECHAM5, MRI-CGCM2.3.2a, NCAR-CCSM3.0, NCAR-PCM1, UKMO-HADCM3, UKMO-HADGEM1

^{*&}lt;sup>3</sup> BCC-CM1, BCCR-BCM2.0, CCCMA-CGCM3.1 (T47), CNRM-CM3, CSIRO-Mk3.0, GFDL-CM2.0, GFDL-CM2.1, GISS-ER, INM-CM3.0, IPSL-CM4, MIROC3.2, MIUB-ECHO-G, MPI-ECHAM5, MRI-CGCM2.3.2a, NCAR-CCSM3.0, NCAR-PCM1, UKMO-HADCM3, UKMO-HADGEM1

Variable	Abbreviation	Description
Temperature	AMT	annual mean temperature
	AMT _{max.}	annual mean maximum temperature
	AMT min.	annual mean minimum temperature
	HW freq.	frequency of hot waves per year
	HW _{dur} .	maximum duration of consecutive hot days (days with Tmax above 30 Celsius)
	T max.	averaged daily maximum temperature
	Tmin.	averaged daily minimum temperature
	FD	number of frost days with Tmin below 0 Celsius
	HWDI	heat wave duration index (defined as the maximum number of consecutive
		days during the year when the daily maximum temperature was greater than 5
		degrees Celsius above the normal maximum temp.
	HD	number of hot days with Tmax above 30 Celsius
Precipitation	DRY _{max.}	maximum number of consecutive dry days
	MMP	mean monthly precipitation
	$DP \ge 0.1mm$	days when daily precipitation is 0.1mm or less
	AMP	annual mean precipitation /rainfall anomaly
Sea level rise	RSLR	relative sea level rise

Table 24-5: Descrip	ption of climate paramete	er abbreviations used i	n Tables 24-2 to 24-4
	phon of enhance parameter		11 1 1 1 1 1 2 1 2 1 0 2 1 1.

	Central Asia	East Asia	North Asia	South East Asia	South Asia	West Asia
Observed Impacts	- High impact, mountain glaciers melt (Casassa et al., 2009)	 High impact in arid area, e.g. Mongolia, and northwest China Groundwater drops in northeast Mongolia Monsoon rainfall impact on water quality in South Korea Increased carbon and nutrients from mountainous watershed during typhoons in Japan and Taiwan China 		 Precipitation relates to dissolved oxygen, PH, and productivity in Mekong river 		
Projected impacts		 Possible positive impact in Yellow River basin - 	- In most of Russia, an increase of evaporation in warmer climate and precipitation is projected to have positive impact on water availability (Alcamo et al., 2007)		 Projected heavy impact due to high dependence on irrigated agriculture and melt water in Indus and Brahmaputra (Immerzeel et al., 2010) Possible increase the risk of floods in Mahanadi (Asokan and Dutta, 2008) Projected increase of rainfall offset the water demand in Ganges (Fung et al., 2011) Projected vulnerability in Coastal ground freshwater in South India, Bangladesh and China (Ranjan et al., 	

Table 24-6: Summary of observed and projected impacts in the water sector.

	Central Asia	East Asia	North Asia	South East Asia	South Asia	West Asia
Observed Impacts		In China, assessed rice yield responses to recent climate change at experimental stations for the period of 1981–2005 (Zhang et al. 2010). There was a variable climate to yield relationships. In some places, yields were positively correlated with temperature when they were also positively related with radiation. However, in other places, lower yields with higher temperature was accompanied by positive				In Jordan, in year 1999, the total production and average yield of wheat and barley were the lowest among the years 1996 to 2006.
Projected Impacts	The impacts to food production will vary by country. Cereal production in northern and eastern Kazakhstan can could benefit from the longer growing season, warmer winters and slight increase in winter precipitation, Western Turkmenistan and Uzbekistan, where frequent droughts will could negatively affect cotton production, increase already extremely high water demands for irrigation, and exacerbate the already existing water crisis and human- induced desertification.	correlation between yield and rainfall. In China, impacts of climate change to crop productivity have mixed results. Rice yields could decline with increasing temperature if CO2 effect is not considered. With CO2 fertilization, rice yields may increase with higher temperature. In China's most productive wheat growing region, winter wheat yields would increase on average by 0.2 Mg ha-1 in the earlier period in 2015- 2045 and by 0.8 Mg ha-1 in the later period in 2070-2099 due to warmer nighttime temperatures and higher precipitation, under A2 and B2 scenarios using HadCM3 model (Thomson et al., 2006). In a wheat-maize cropping system in Huang-Huai-Hai (3H) Plain, China, a 2 and 5 oC increase in temperature, precipitation increasing and deceasing by 15 and 30%; and atmospheric CO2 enrichment to 500 and 700 ppm would result to in a mean relative yield change (%) (RYC in %) of -10.33% and the lowest and highest RYC values of -46% and 49%, respectively. However with CO2 fertilization a positive change in RYC	In Russia, climate change may also lead to "food production shortfall" which is defined as an event in which the annual potential (i.e. climate- related) production of the most important crops in an administrative region in a specific year falls below 50% of its climate-normal (1961– 1990) average (Alcamo et al., 2007). The frequency of shortfalls in the main crop growing regions in the same year is around 2 years/decade under climate baseline conditions but could climb to 5–6 years/decade in the 2070s using the ECHAM and HadCM3 models and the A2 and B2 SRES.	In Indonesia, the date of rice planting could shift with a marked increase in the probability of a 30-day delay in monsoon onset in 2050 as a result of changes in the mean climate (Naylor et al. 2007).	In Swat and Chitral districts of Pakistan, there were mixed results as well (Hussain and Mudasser, 2007). Projected temperature increase of 1.5 and 3 °C would cause to wheat yields to decline by 7% and 24% respectively in Swat district but would lead to an increase (by 14% and 23% respectively) in Chitral district.	In Western Asia, a rise in CO2 concentration may benefit semi-arid crops by increasing crop water use efficiency and net photosynthesis leading to greater biomass, yield and harvest index (Ratnakumar et al., 2011). For example, wheat and rice grain yield increased by an average of 10-20% at ample N and water with elevated CO2 (350pm to 700ppm).
		2010). In Japan, increasing water temperature (1.6– 2.0 °C) could lead to a northward shift of the isochrones of safe transplanting dates for rice seedlings (Ohta and Kimura, 2007).			In India, changing climate projected to reduce monsoon sorghum grain yield by 2 to 14% by 2020 with worsening yields by 2050 and 2080 (Srivastava et al., 2010). In the Indo-Gangetic Plains (IGPs), a similar reduction in wheat yields is projected, unless appropriate cultivars and crop management practices were adopted (Ortiz et	In Yarmouk basin, Jordan, simulation with DSSAT showed that wheat and barley yields will decline by 10-20% and 4-8% respectively with 10-20% reduction in rainfall (Al-Bakri et al., 2010).

Table 24-7: Summary of observed and projected impacts in the food sector.

		al 2008)	1
		 al., 2008).	
		In the Indo-Gangetic Plains (IGPs) climate projections based on a doubling of CO2 using a CCM3 model downscaled to a 30 arc- second resolution as part of the Worldclim data set showed that there will be a 51% decrease of in the most favorable and high yielding area due to heat stress.	
		In Sri Lanka, tea cultivation at low and mid-elevations are more vulnerable to the adverse impacts of climate change than those at high elevations. Projected coconut production after 2040 in all climate scenarios will not be sufficient to meet local consumption. The total impact on agriculture (rice, tea, rubber and coconut) production ranges from a decrease of US\$96.4 million (-20%) to an increase of US\$34,214 million (+72%) depending on the climate scenarios (Eriyagama et al. 2010).	
		In Sri Lanka, studies on rice production have mixed results (Eriyagama et al., 2010). An earlier study showed that a 0.1-0.5°C increase in temperature could depress rice yield by approximately 1-5%. However, another experiment revealed that rice yields respond positively (increases of 24 and 39% in the two seasons) to elevated CO2 even at higher growing temperatures (>30°C) in subhumid tropical environments.	

Table 24-8: Summary of adaptation options for agriculture in Asia.

Crop	Country/ Region	Recommended/ Potential Adaptation strategies	Benefits/ Co-Benefits	References
Wheat	General	Conservation agriculture (reductions in tillage, surface retention of adequate crop residues, and diversified, economically viable crop rotations)	Improve rural incomes and livelihoods by reducing production costs, managing agroecosystem productivity and diversity more sustainably, and minimizing unfavorable environmental impacts	Ortiz et al., 2008
Wheat	Pakistan	Development of short duration and high yield varieties of wheat.	Can withstand climatic anomalies expected in the future	Hussain and Mudasser (2007)
Wheat	Indo- Gangetic Plains, India	Development of heat-tolerant wheat germplasm, as well as cultivars.	Better adapted to heat and conservation agriculture	Ortiz et al., 2008
Barley; wheat	Jordan	Soil water conservation. Selection of drought tolerant genotypes with shorter growing seasons.	Increase available water to crop	Al-Bakri et al., 2010
Sorghum	India	Changing variety and sowing date	Reduce the impacts on monsoon sorghum to about 10%, 2% and 3% in 2020 scenario. Reduced impacts on winter crop to 1–2% in 2020, 3–8% in 2050 and 4–9% in 2080.	Srivastava et al., (2010
Rice	Sri Lanka	Traditional approaches for resolving water stress, such as increasing water use efficiency, water harvesting and/or reducing cropped areas. Earlier planting and shorter duration varieties to avoid the impacts of less rainfall in January and February.		De Silva et al., 2007.
Rice	China	Shifts in planting dates and automatic application of irrigation and fertilization. Selection for more temperature-tolerant cultivars and later-maturing cultivars to take advantage of longer growing seasons		Tao et al., (2008)
Corn	China	Using high-temperature sensitive varieties Early planting, fixing variety growing duration, and late planting	Using high-temperature sensitive varieties, maize yield could averagely increase by 1.0–6.0%, 9.9–15.2%, and 4.1–5.6%, by adopting adaptation options of early planting, fixing variety growing duration, and late planting, respectively	Tao and Zhang (2010)
General	India	Water harvesting		Kelkar et al (2008)
General	South Asia	Increasing livestock production relative to crops Selection of crop varieties Livelihood diversification		Morton, 2007
General	Central Asia	The replacement of the existing network of open irrigation canals by more efficient drip irrigation systems Development of early warning systems, such as drought forecast, pest and epidemic disease forecasts, and water quality monitoring systems.	Could significantly reduce evaporative water loss, while simultaneously improving crop productivity, reducing soil salinization, and decreasing risks of water contamination and transmission of vector-borne and waterborne diseases.	Lioubimtseva and Henebry (2009)
General	West Asia	Changing of cropping systems and patterns, switching from cereal-based systems to cereal-legumes and diversifying production systems into higher value and greater water use efficient options. Using supplementary irrigation systems, more efficient irrigation practices and the adaptation and adoption of existing and new water harvesting technologies. Development of more drought and heat tolerant germplasm using traditional and participatory plant breeding methodologies and better predictions of extreme climatic events.		Thomas 2008
General	Russia	Crop substitution Diversification of crops Expanding irrigated agricultural areas Strategic food reserves, Improving management, Monitoring and early warning systems, Food imports from abroad.		Alcamo et al., (2007)
General	Philippines	Crop diversification; change of crop varieties, use of water conservation pratices		Peras et al., 2008; Lasco et al., 2011
		Cultivars with multiple resistance to insects and diseases		Sharma et al., 2010

Issues	Country/	Recommended/ Potential Adaptation	Benefits/ Co-	References
	Region	strategies	Benefits	
Environmental impacts	Asia	Develop new or transform existing settlements with green infrastructure	Contribute towards low- carbon society	ADB, 2012
Water infrastructure	Bangladesh	Consider factors such as changes in development patterns, water use upstream, land use change and population and economic growth in addition to scenarios. Incorporation of climate proofing of infrastructure in the project design		ADB, 2011a
Disaster resilience (flood risk)	India Bangladesh	Integrated approaches through existing community-based practices and enhancement of social capacity		Surjan and Shaw, 2008; Prashar, et.at., 2012; Rotberg, 2010
		Local government-neighborhood group partnership	Promotion of waste reduction	Surjan and Shaw, 2009
Water resources	India and Nepal	Rainwater harvesting by modifying building bye-laws, providing financial assistance or discounts		UNESCO, 2012; Pandey, et.al., 2003
Flood mitigation	Japan	Underground river consisting of a pipe of diameter 10-12.5 meters and few kilometers long		Zevenbergen and Herath, 2008
Non-structural flood control measures	Japan	Land-use zoning Flood proofing Flood risk mapping Onsite run off control strategies Rainfall storage Infiltration facilities Use of public spaces, such as parks for flood retardation		

Table 24-9: Summary of adaptation options.

Observed change / Impact	Country/ Region		References
Poor are disproportionately impacted by	East and South Asia		Kim, 2011
climate related hazards			
Increased migration due to	Mekong region		Warner, 2010;
environmental (e.g. rapid onset			Black et al., 2011
disasters), social and economic reasons			
Leave farming due to repeated droughts	South Asia		Kulkarni and Rao,
			2008
Loss of crops, income and fallows	Cambodia		Nguyen et al., 2009
Projected Impacts	Country/Region	Projection Details	References
Negative impact on rice crop, increase	Asia	GTAP Model,	Hertel et al., 2010
in food price and cost of living,		projections for 2030,	
increased poverty		scenarios: Impacts	
		resulting low, medium	
		and high productivity	
Loss of livelihoods to indigenous people	Tibet/Himalayas	Qualitative	Salick et al., 2009; Xu
from declining alpine biodiversity		observations	et al., 2009

Table 24-10: Summary of observed changes and projected impacts for livelihoods and poverty.

Aspect/	Country/	Recommended/ Potential	Benefits/ Co-Benefits	References
Issues?	Region	Adaptation strategies		
Delay and shortfall in rainfall	Indonesia	Access to credit and public works project	Able to protect food expenditure in the face of weather shocks	Skoufias et al., 2011
General (droughts, floods etc)	General	Weather index insurance, cattle insurance, seed banks, credit facilities, assisted migration, cash for work	Poverty cantered adaptation, creation of assets and access to resources	Barret et al., 2007; Tanner and Mitchel, 2008; Jarvis et al., 2011
General	General	Assisted migration	Build financial, social and human capital	Barnett and Webber, 2010
General	Vietnam	Yield growth and improving agriculture labour productivity	Rural poverty reduction, livelihood diversification	Janvry and Sadoulet, 2010
Droughts and floods	Philippines	Bundlingofimprovedvarietiesandagronomicpracticesandcombinationproductionandmarket	Economic benefits and social learning	Acosta-Michlik and Espaldon, 2008
General	Asia	Community based adaptation	Capture information at the grassroots, help integrating disaster risk reduction, development, and climate change adaptation, connect local communities and outsiders, and addresses the location specific nature of adaptation.	Aalst et al., 2008; Heltberg et al., 2010; Rosegrant, 2011
General	Asia	Forest management	Resilient livelihoods, buffer from shocks	Chhatre and Agrawal, 2009
General	Asia	Securing rights to resources, community forest tenure rights	Resilient livelihood benefits to the poor indigenous and traditional people	Macchi et al., 2008; Angelsen, 2009
Biodiversity loss	Tibet	Greater involvement of traditional and indeginous people in climate change adaptation decision making	Indigenous knowledge from the years of living in close harmony with nature	Byg and Salick, 2009; Salick et al., 2009

Table 24-11: Summary of adaptation options for securing livelihoods in Asia.

Table 24-12: Location and major characteristics of central Asia glaciations.

	Alatai-Sayan mountains									
Geo-coordinates	Total glacier area in 2009 (km ²)	Quantity of glaciers	ELA, ave. (km, a.s.l.) in 2009	Distribution area (km, a.s.l.)	Glacier thickness, ave. (km)					
45°-54°N; 84°– 103°E	1,562	2,340	2.8	2.1-4.5	0.057					
		Pamir m	ountains							
36°-40°N; 66°-76°E	13,424	11,671	4.6	3.4-7.7	no data					
Tien Shan mountains										
39°-46°N; 69°-95°E	13,196	10,925	4.4	2.8-7.4	no data					



Figure 24-1: The land and territories of 51 countries/regions.



Hazard mortality risk (floods, tropical cyclones and precipitation-triggered landslides)

Source: Developed by the GAR team at UNISDR.

Figure 24-2: Hazard mortality risk. Source: *The United Nations World Water Development Report 4: Managing Water under Uncertainty and Risk*. WWAP (World Water Assessment Programme), UNESCO, Paris. Available at: http://www.unesco.org/new/fileadmin/MULTIMEDIA/HQ/SC/pdf/WWDR4%20Volume%202-Knowledge%20Base.pdf Page 117.



Figure 24-3: Map of Lower Mekong Basin from Mekong River Commission Technical Paper No. 24, 2009 (MRC, 2009).



Figure 24-4: The difference in losses of glacier area in Altai-Sayan, Pamir and Tien Shan determined by location of the mountain ridges in relation to major atmospheric moisture flow and by elevation a.s.l. Remote sensing data analysis from 1960s (Corona) through 2009 (Landsat, ASTER and Alos Prism).



Figure 24-5: The MODIS-Terra satellite image of the Aral Sea on 18 August 2008. Image courtesy by D.M. Soloviev, Marine Hydrophysical Institute, Sevastopol, Ukraine, basing on the data provided by the LAADS Web, NASA-Goddard Space Flight Center (http://ladsweb.nascom.nasa.gov/). Red line shows the Aral Sea coastline in 1960. Yellow line shows the border between Kazakhstan and Uzbekistan.